

Topic # 13

Atmospheric Circulation

How the Energy Balance Drives It
and How It Results in
Global Climate Patterns

p 77-83 in Class Notes

but first . . . Some review!

The Earth

[as viewed from space]

... has the organized, self-contained look of a live creature, full of information, marvelously skilled in handling the sun.

~ Lewis Thomas



OBJECTIVES:

To understand how the
Radiation/Energy Balance
helps to drive **global weather**
& climate

And learn a bit about how and why
climates vary in different parts of the
world.

RADIATION / ENERGY BALANCE

$$R_{\text{NET}} = \begin{array}{c} \text{SW} \\ \downarrow \end{array} + \begin{array}{c} \text{SW} \\ \vdots \downarrow \end{array} - \begin{array}{c} \text{SW} \\ \nearrow \end{array} - \begin{array}{c} \uparrow \\ \text{LW} \end{array} + \begin{array}{c} \text{LW} \\ \downarrow \end{array} = H + LE + G$$

“Radiation Balance” part

$$R_{\text{NET}} = \begin{array}{c} \text{SW} \\ \downarrow \end{array} + \begin{array}{c} \text{SW} \\ \vdots \downarrow \end{array} - \begin{array}{c} \text{SW} \\ \nearrow \end{array} - \begin{array}{c} \uparrow \\ \text{LW} \end{array} + \begin{array}{c} \text{LW} \\ \downarrow \end{array}$$

All components are referring to electromagnetic radiation

All components are referring to modes of heat energy transfer or heat energy storage involving matter

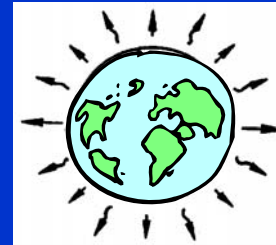
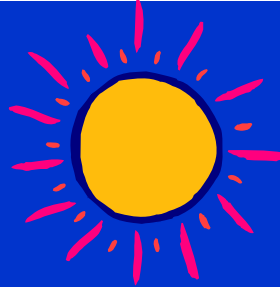
“Energy Balance” part

$$R_{\text{NET}} = H + LE + G$$

Start out here,
with energy
from the SUN
radiated to
Earth and so
forth . . .

“Radiation Balance” part

$$\begin{array}{ccccccc} \text{SW} & & \text{SW} & & \text{SW} & & \text{LW} \\ \downarrow & + & \downarrow & - & \nearrow & - & \downarrow \\ & & & & & & \text{LW} \end{array} = R_{\text{NET}}$$



The R_{NET} is then able
to be used in thermal
energy “heat transfer”
processes which
manifest themselves
as weather & climate!

“Energy Balance” part

$$R_{\text{NET}} = H + LE + G$$

Thermal Energy Review

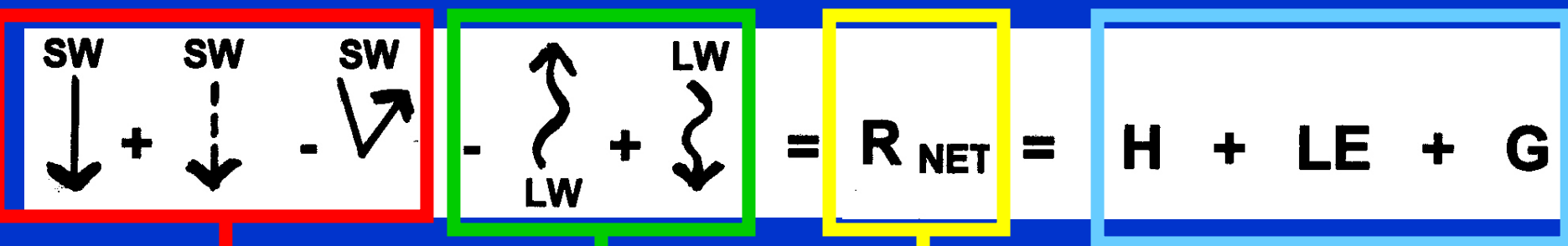
Heat (def) = the thermal energy that is transferred from one body to another because of a temperature difference.

- Sensible Heat transfer (H)
- Latent Heat transfer (LE)

plus (after transfer) thermal energy can be **STORED (G)**

$$H + LE + G$$

ENERGY IN THE EARTH-ATMOSPHERE SYSTEM



Ultimate source
of energy is the
SUN (SW)

After
absorption of
SW, LW energy
is radiated in &
out by EARTH
& Atmosphere

Any
NET
(leftover)
energy

Goes into
the HEAT
TRANSFER
processes that
drive climate

LINKING THE ENERGY BALANCE TO ATMOSPHERIC CIRCULATION . . .

**We'll start with the SUN
(SOLAR INSOLATION)**

IN – SOL- ATION =

**Amount of incoming solar energy
received by a point on Earth's surface**

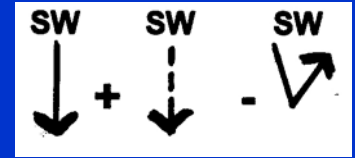
Recall quote on p 77 . . .



How does the Earth

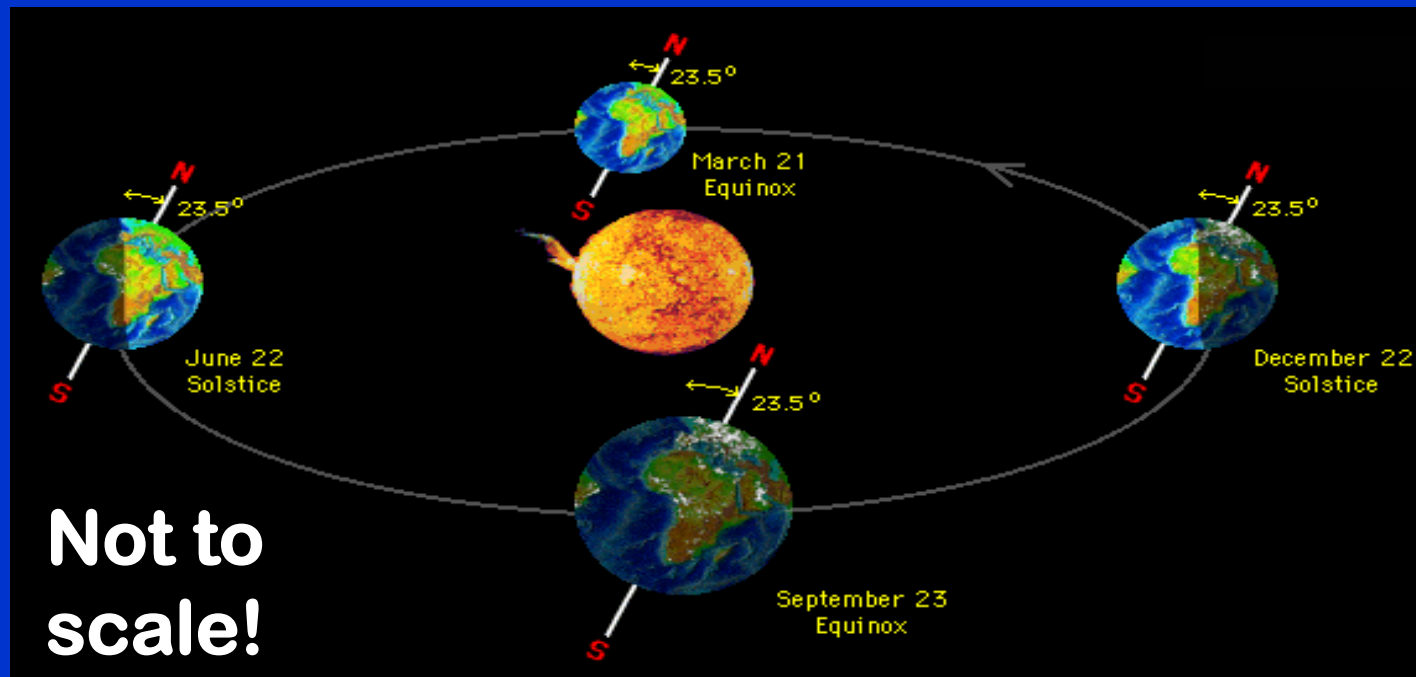
“HANDLE” THE SUN ?????

To drive the circulation, the initial source of energy is from the Sun:



Seasonal & latitudinal variations of solar insolation: 3 Principles →

http://mesoscale.agron.iastate.edu/agron206/animations/01_EarthSun.html



**Not to
scale!**

Seasonal & latitudinal variations of solar insolation:

3 Principles of EARTH-SUN RELATIONSHIPS

*(They define the SEASONS in
different latitudes!)*

#1 OBLIQUITY OF EARTH'S AXIS

#2 ECCENTRICITY OF EARTH'S ORBIT

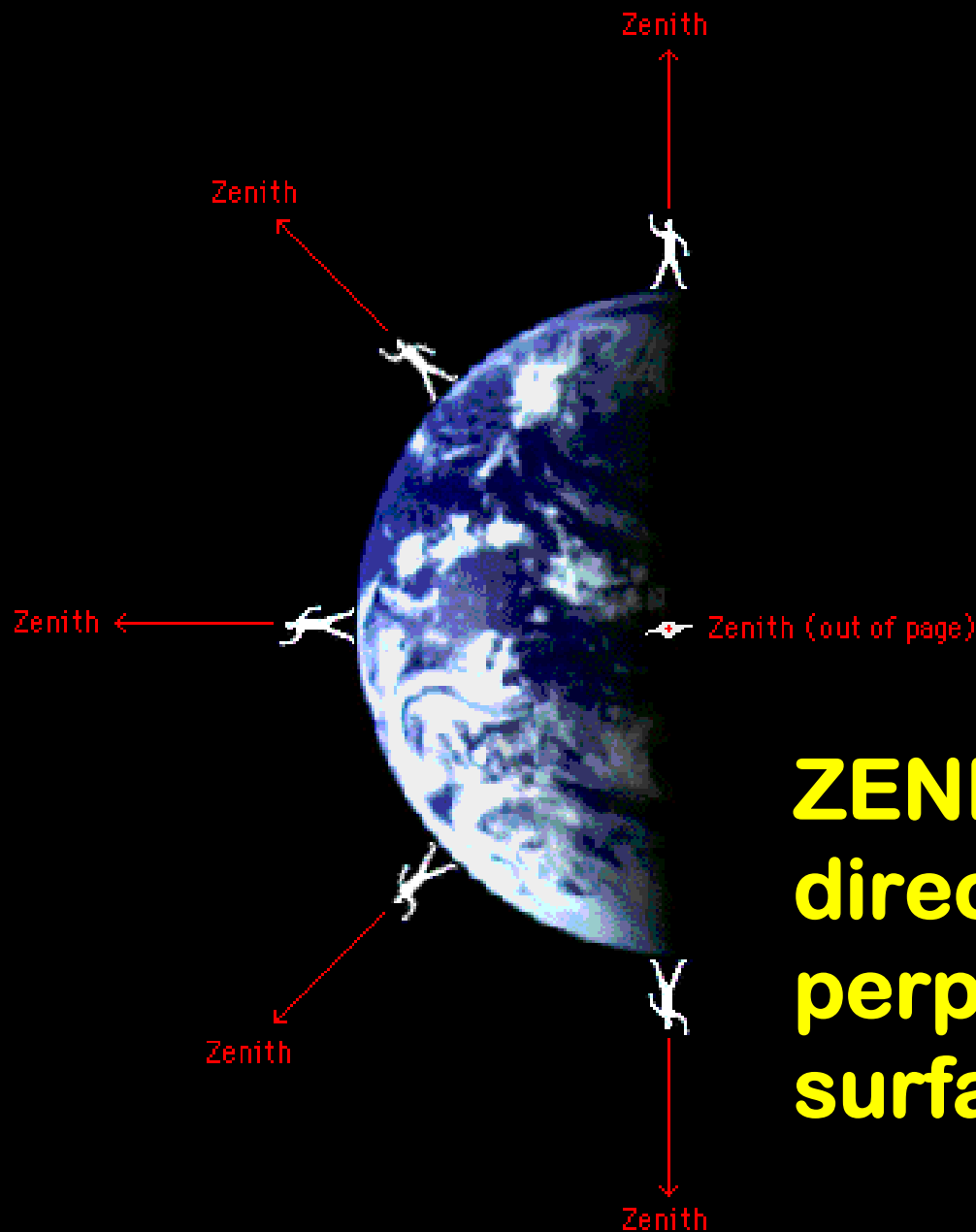
3 Timing of Seasons in Relation to Orbit:

Seasonal & latitudinal variations of solar insolation:

PLUS -- 2 factors that determine the
AMOUNT OF SOLAR INSOLATION
as seasons progress:

- (1) **INTENSITY** of sun's rays
(perpendicular = more intense)
- (2) **DURATION** of daily insolation
(longer day length = more insolation)

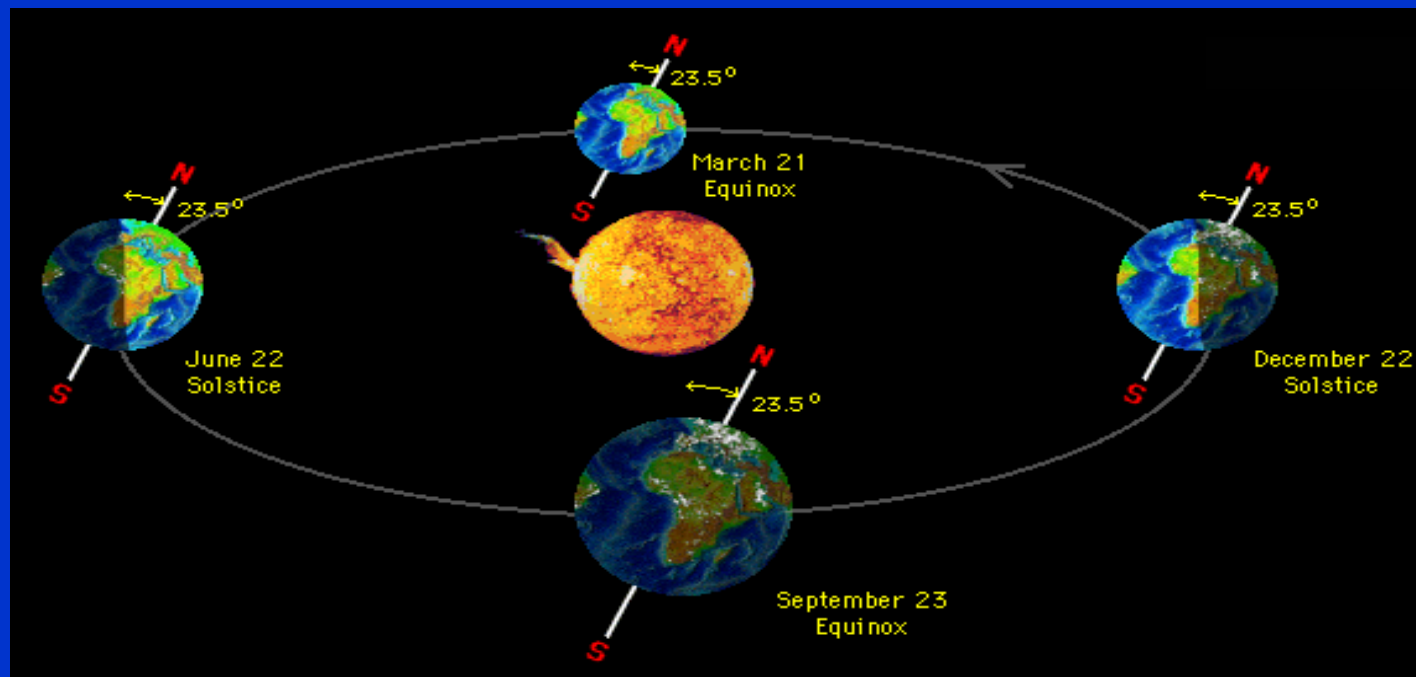
Since the zenith at any place on the earth is the point directly overhead (90° from all horizons), at each place on the earth it points in a different direction into space.



ZENITH – the point directly overhead & perpendicular to the surface

EARTH-SUN RELATIONSHIPS & The SEASONS:

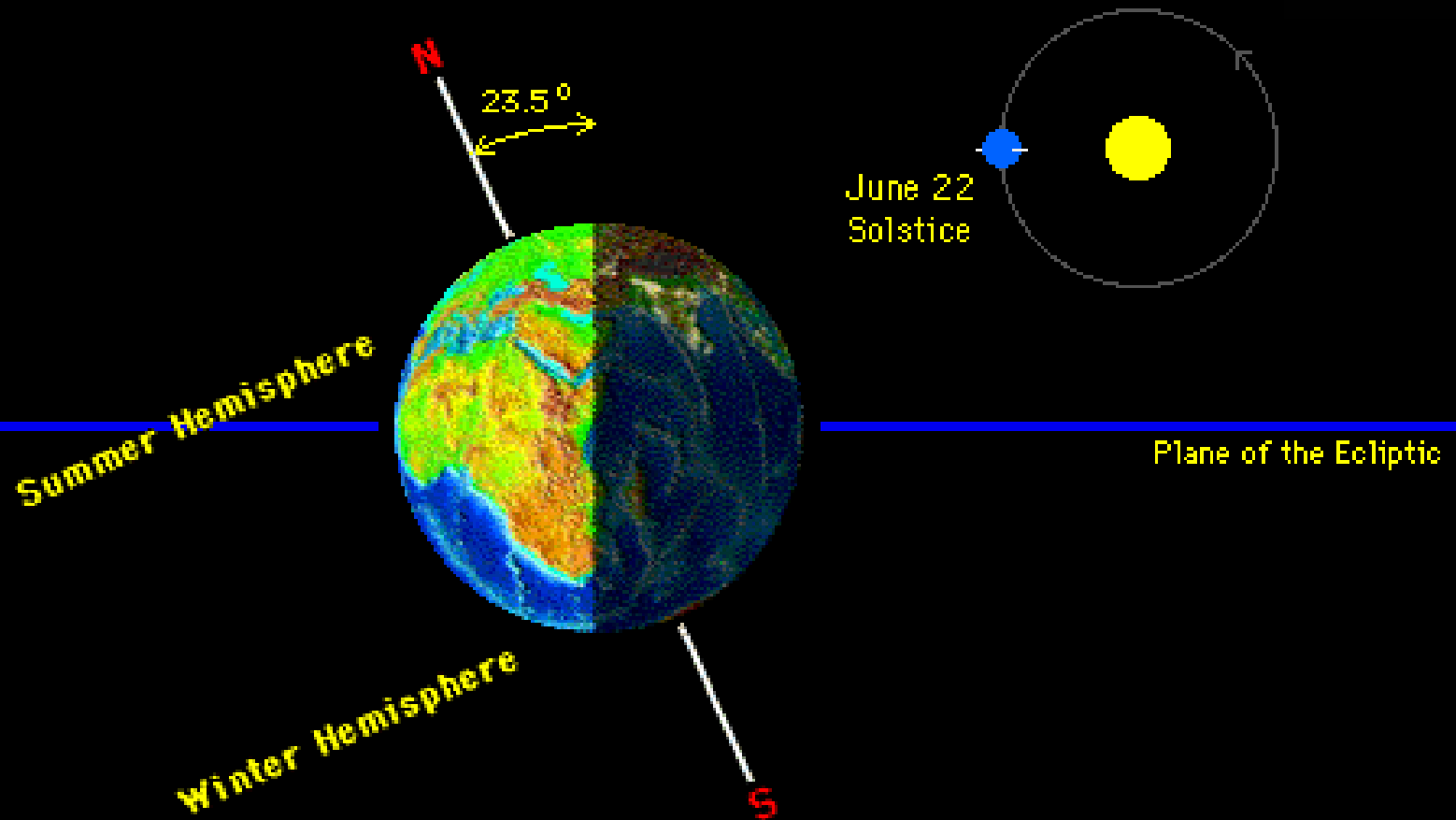
http://mesoscale.agron.iastate.edu/agron206/animations/01_EarthSun.html



#1 OBLIQUITY OF EARTH'S AXIS

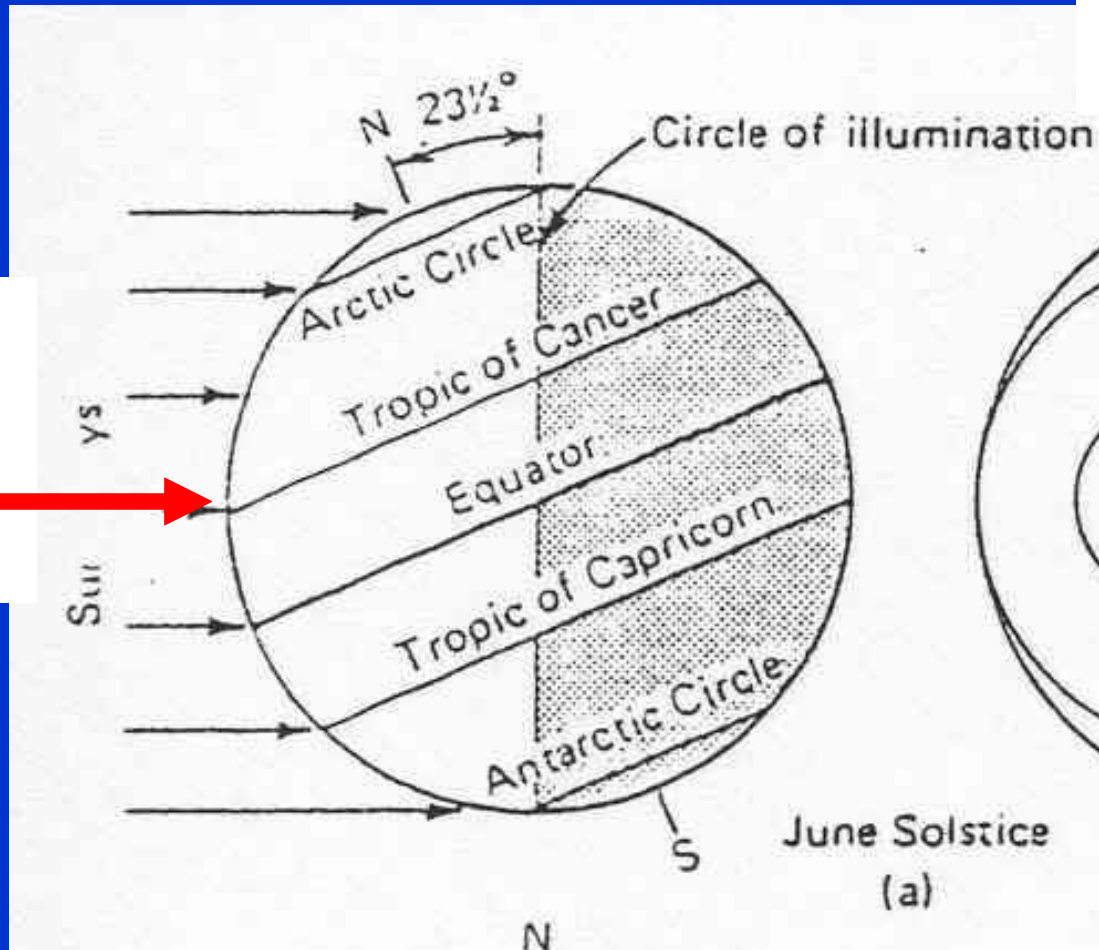
(axis "tilts" 23.5 degrees from plane of ecliptic)

JUNE SOLSTICE

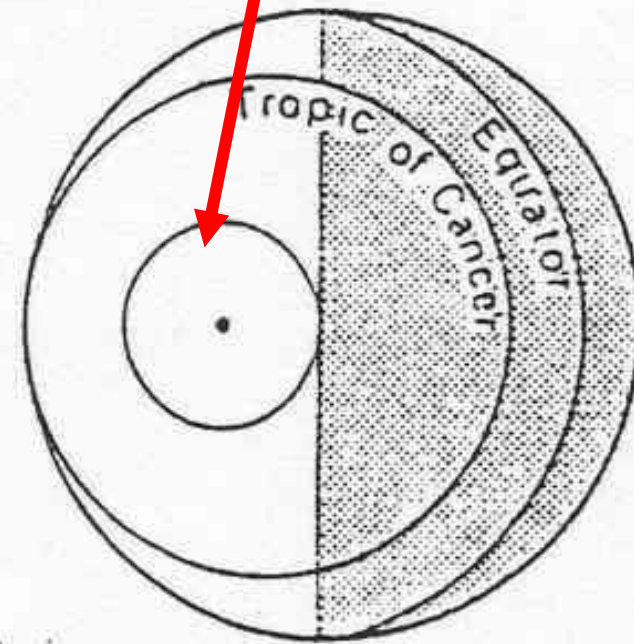


JUNE SOLSTICE

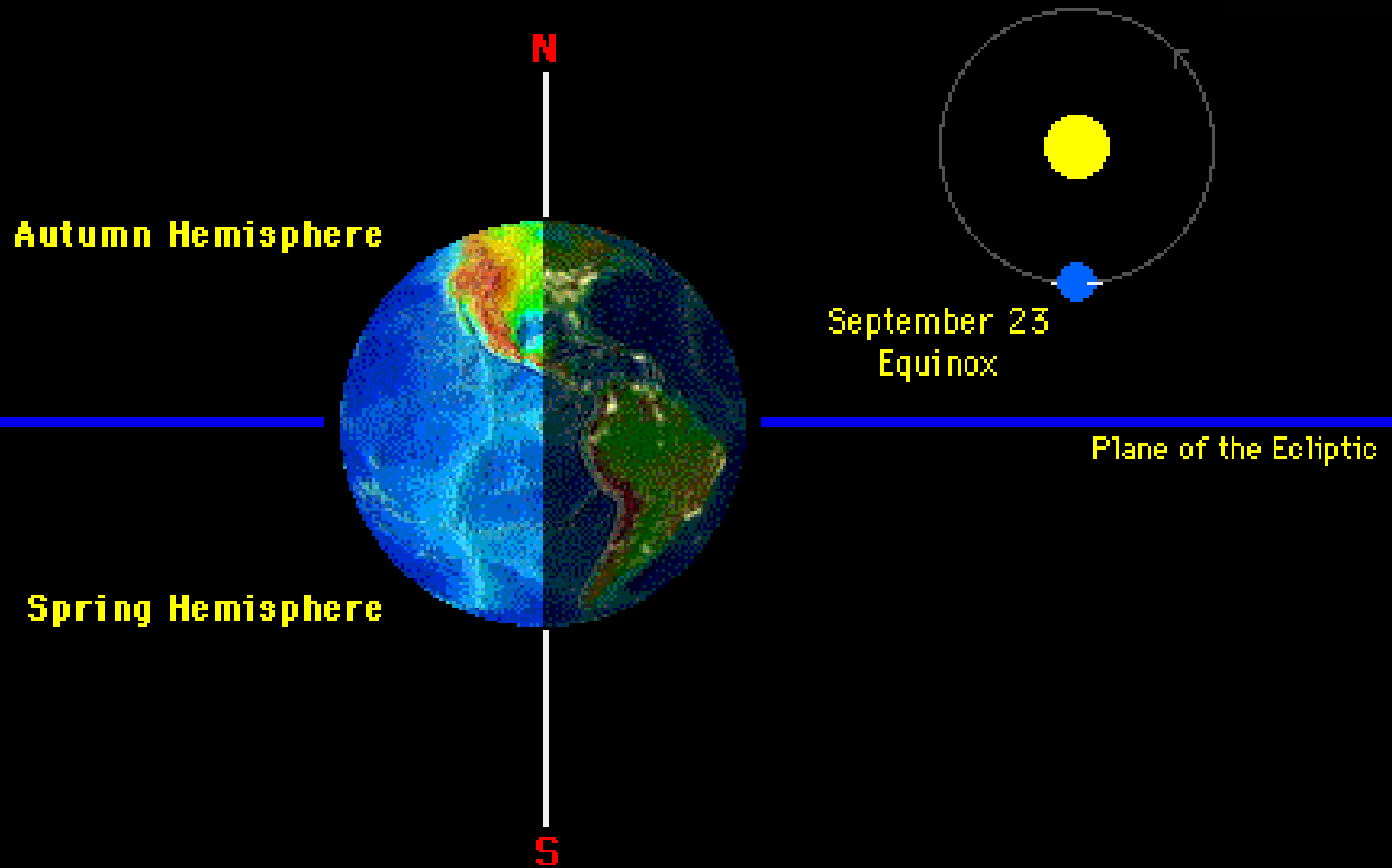
Most
intense
solar
radiation



24 hours
of sunlight

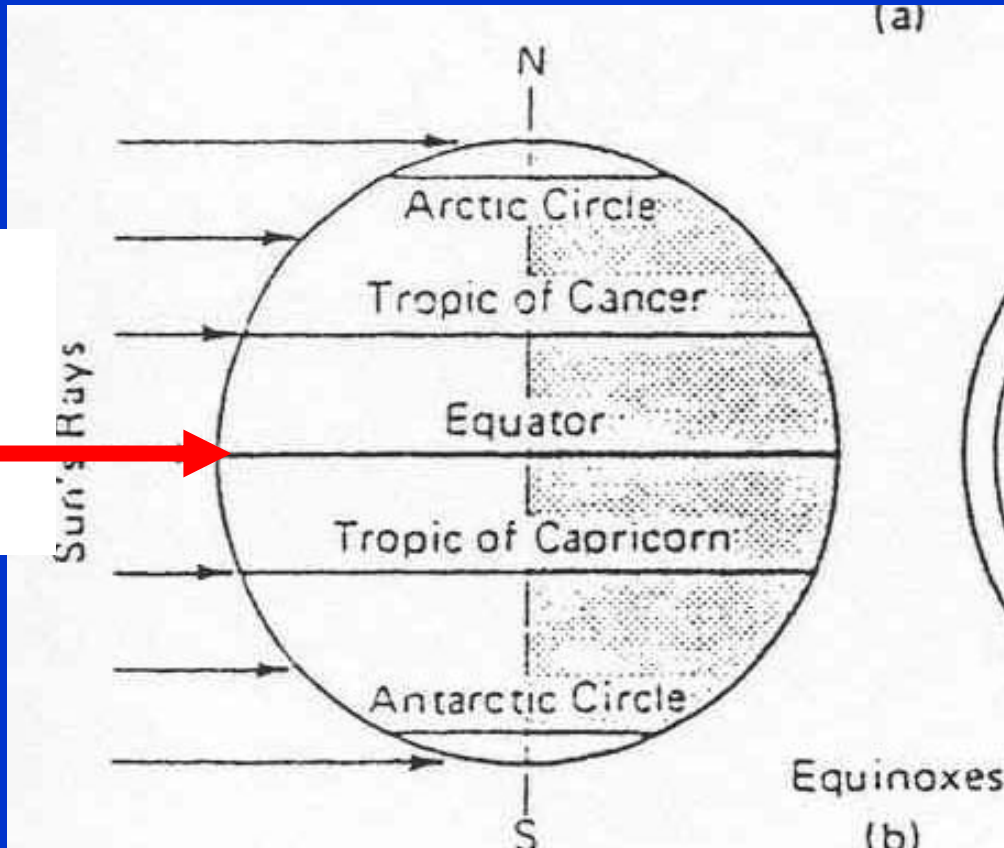


March or September Equinox

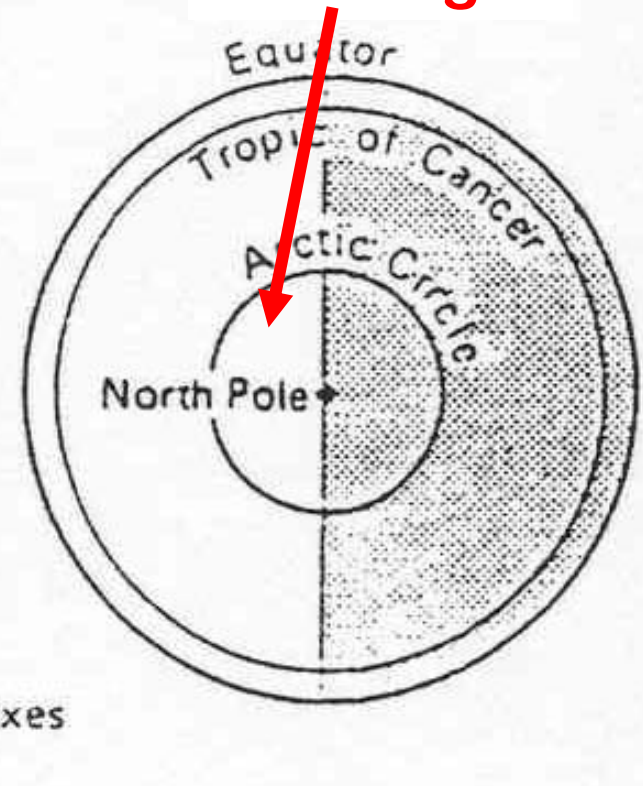


MARCH & SEPTEMBER EQUINOXES

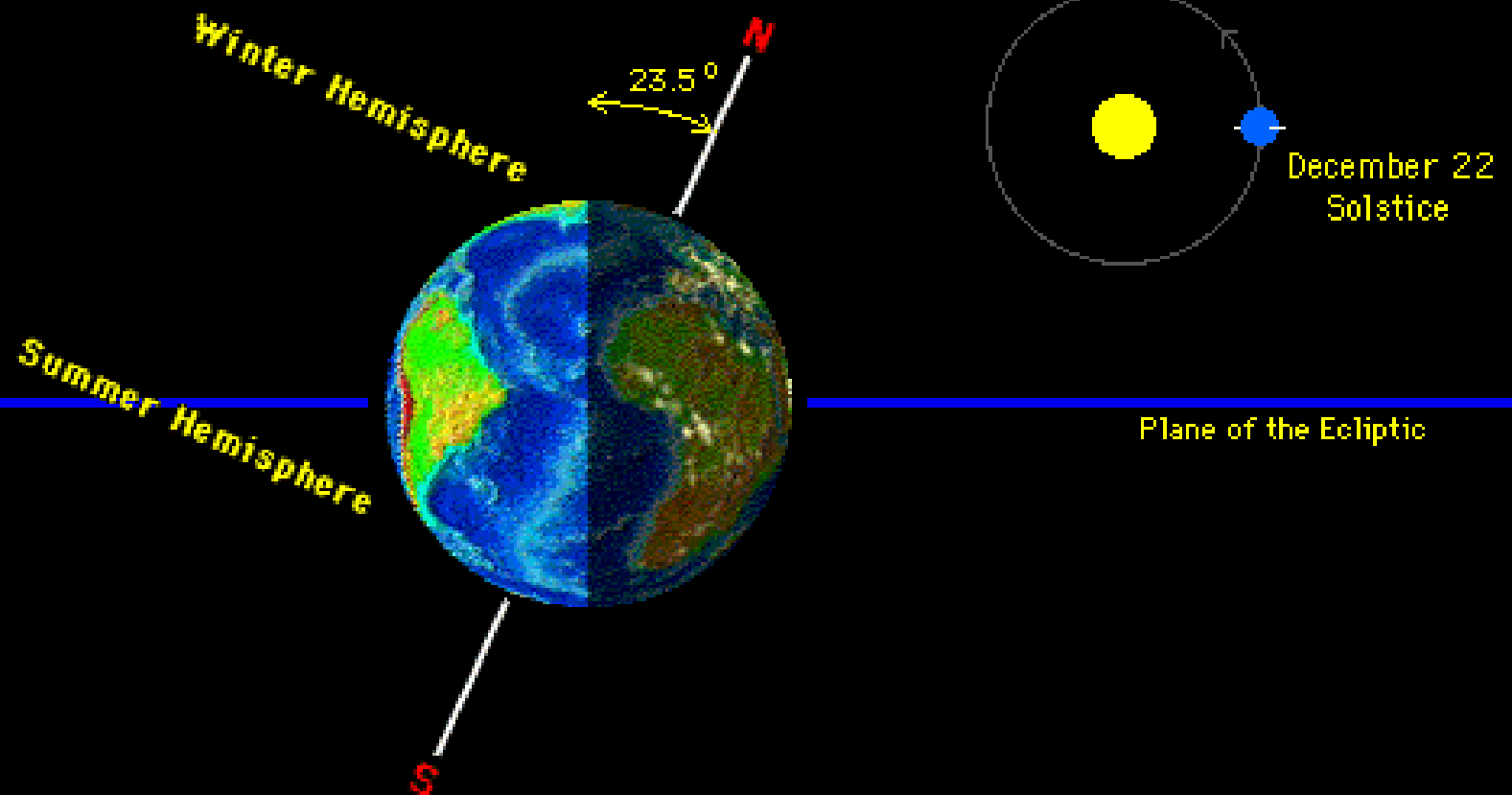
Most intense solar radiation



12 hours of sunlight

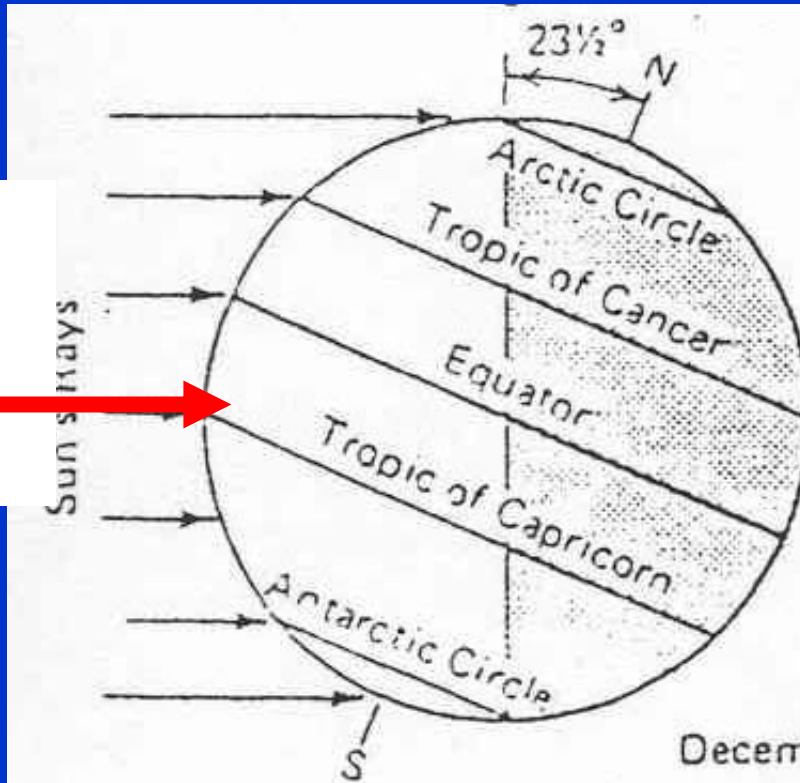


DECEMBER SOLSTICE

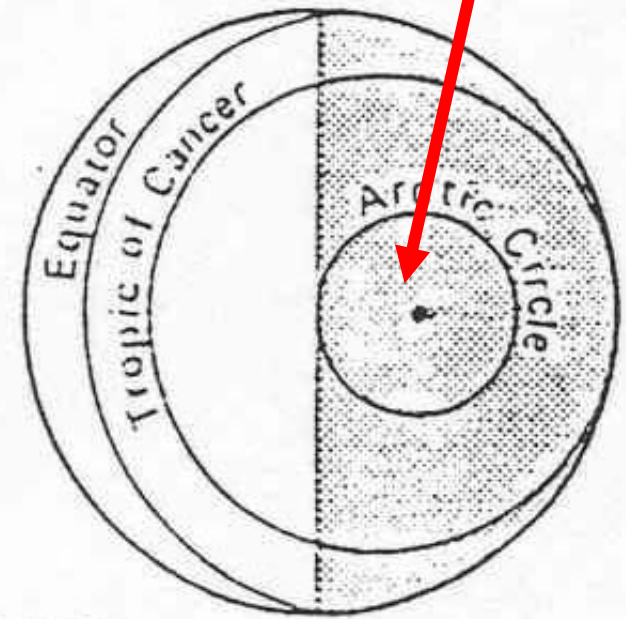


DECEMBER SOLSTICE

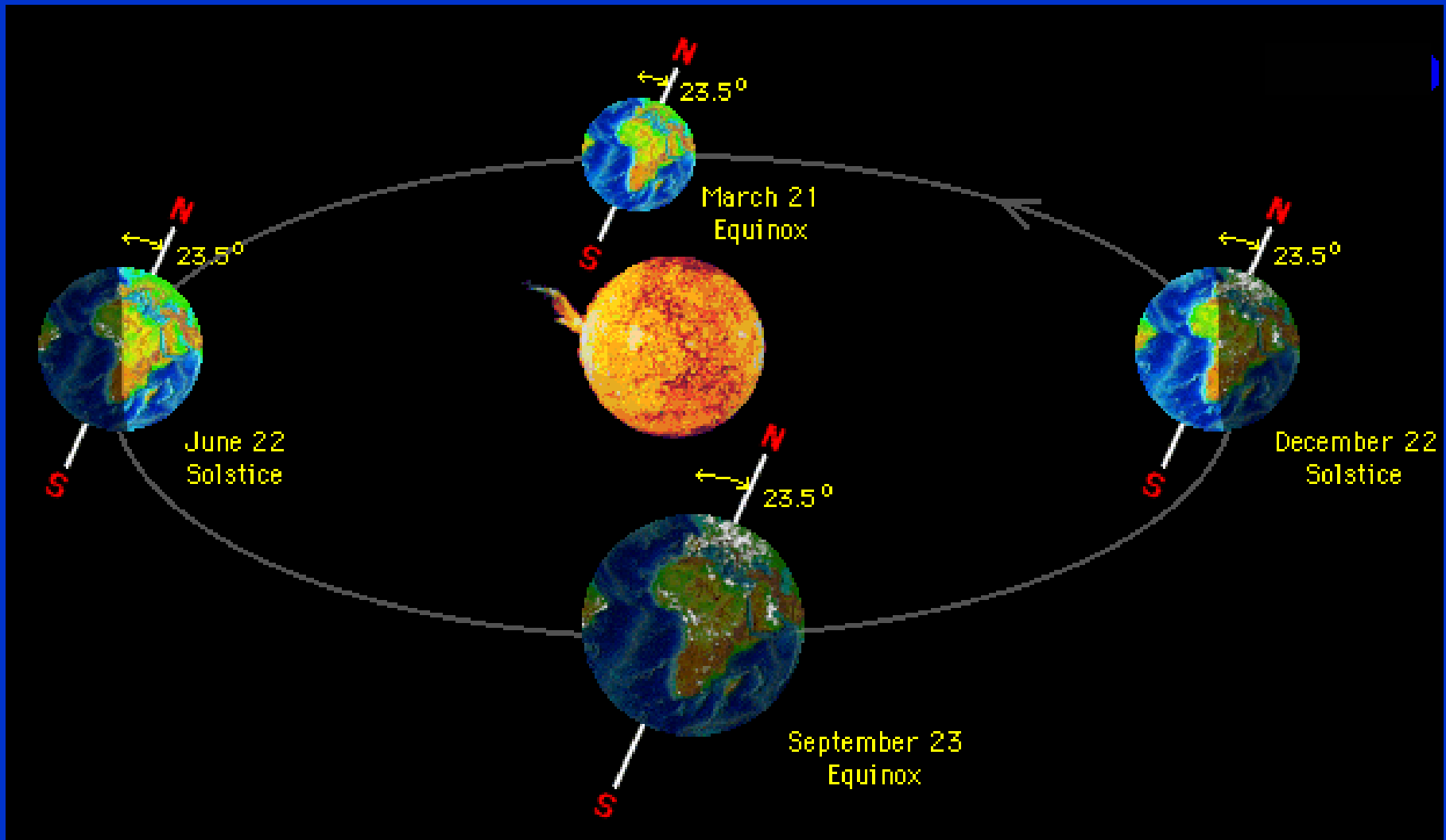
Most
intense
solar
radiation



24 hours of
darkness

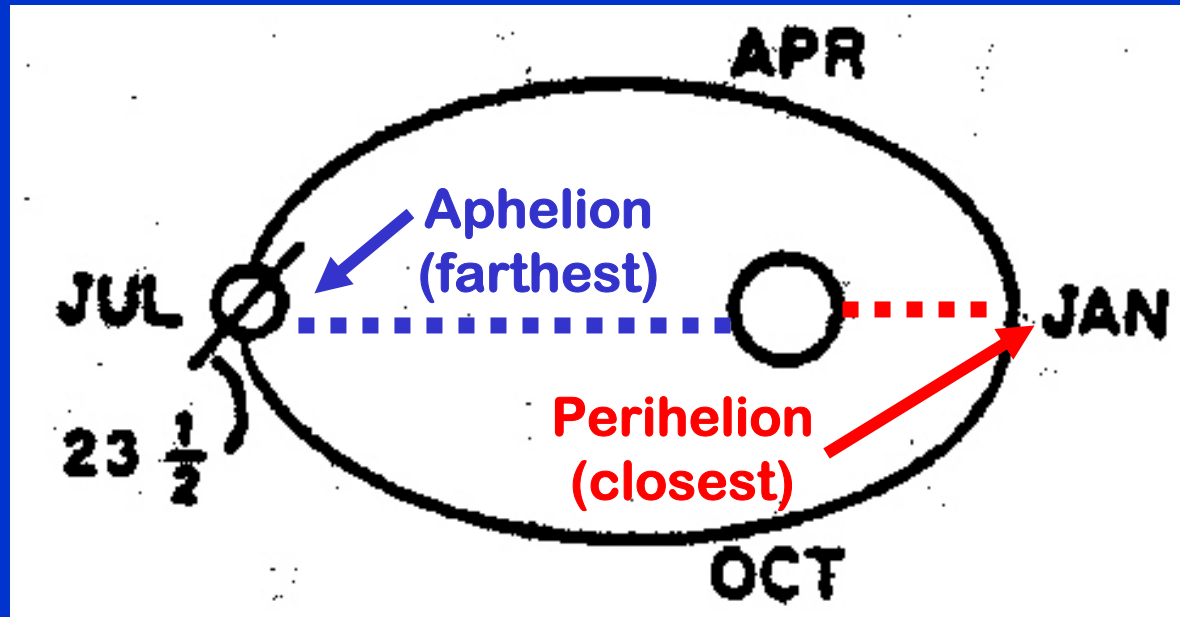


December Solstice
(c)



#2 ECCENTRICITY OF ORBIT

Earth's orbit around sun is not symmetrical:
“ellipse”



3 Timing of Seasons in Relation to Orbit: PRECESSION OF THE EQUINOXES

(Earth currently is closest to the Sun in January
– just after the Dec Solstice; but this has changed in the past)

Take Notes on p 76

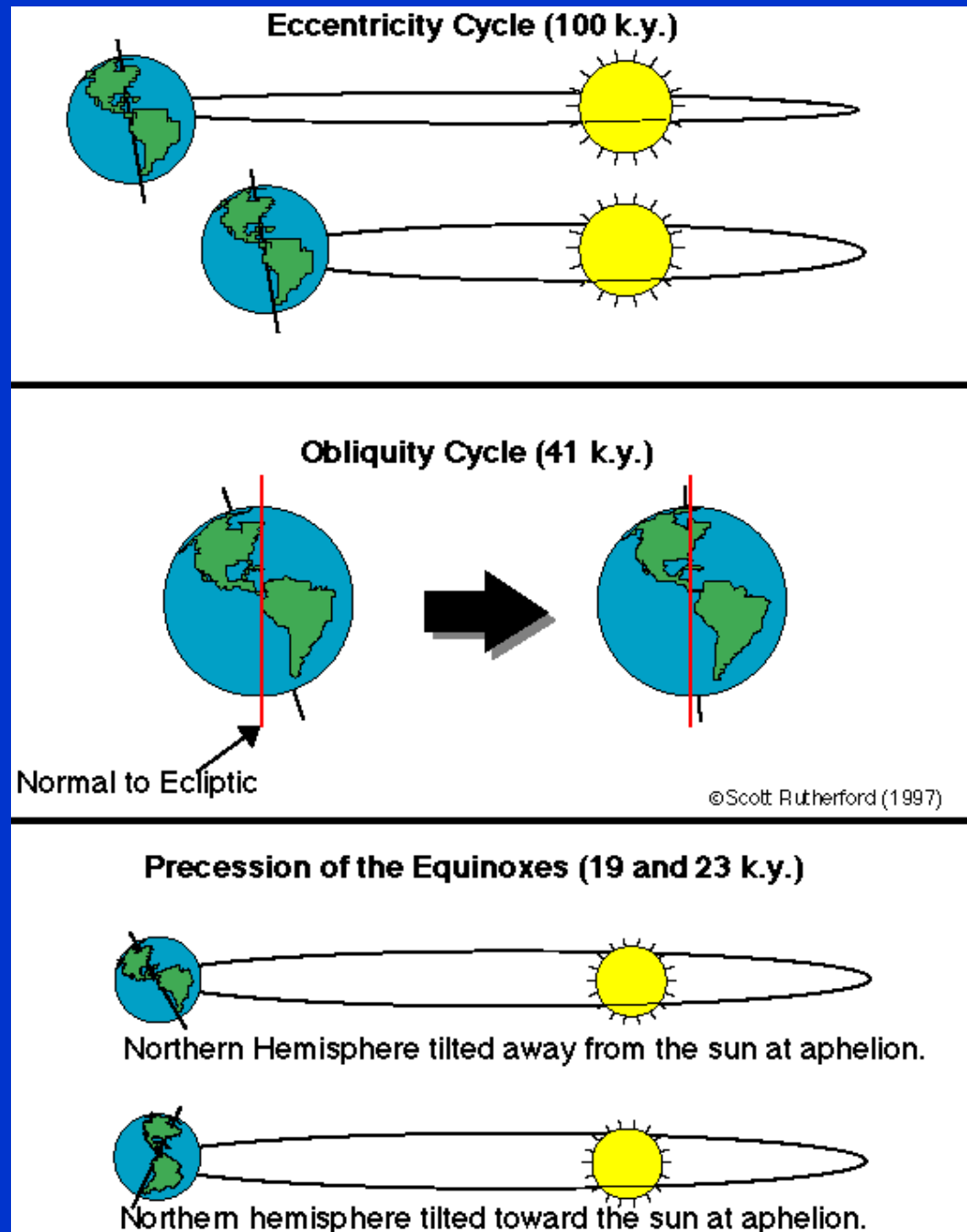
Obliquity

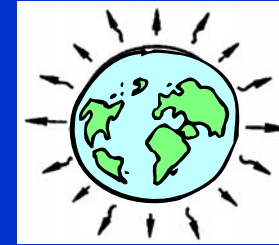
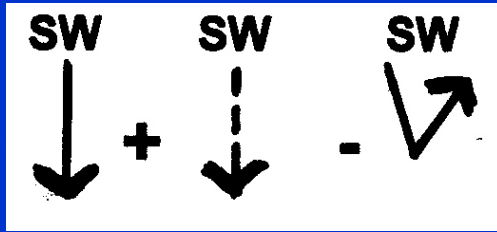
Eccentricity

Precession

.. have all
CHANGED over
the past 100,000
years!!

IGC CD MODULE
SELECT
→ “Weather in
Motion” – Ch 16
Orbital Variations

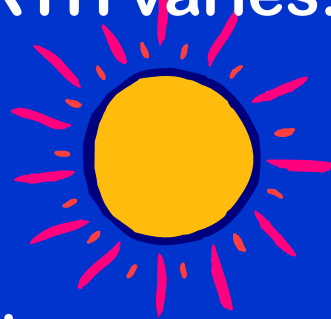




KEY CONCEPT:

The amount of SW absorbed by EARTH varies:

- by LATITUDE
- by SEASON

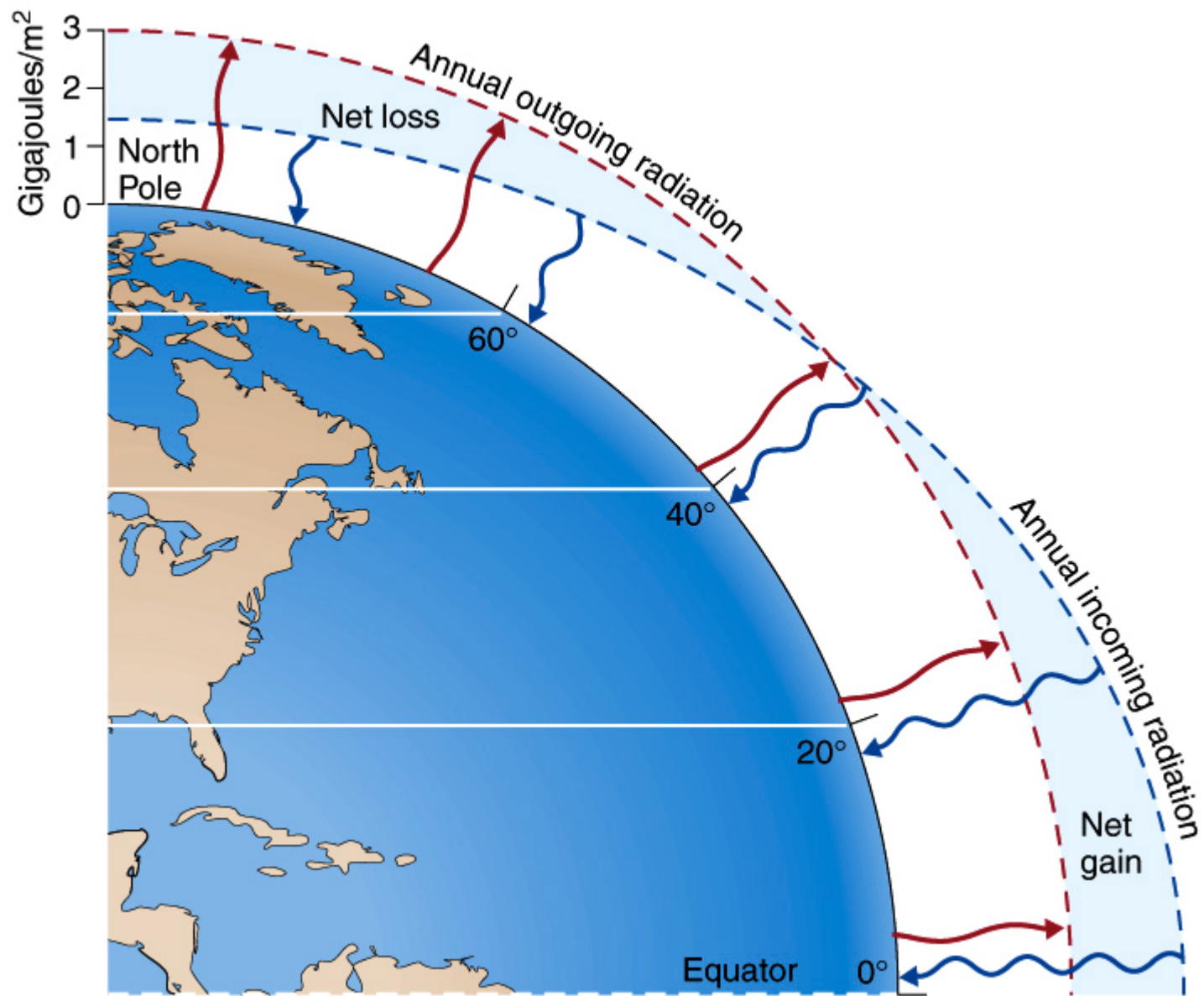


Annually, MORE is absorbed in the LOW LATITUDES (near Equator) than in the HIGH LATITUDES (near Poles)

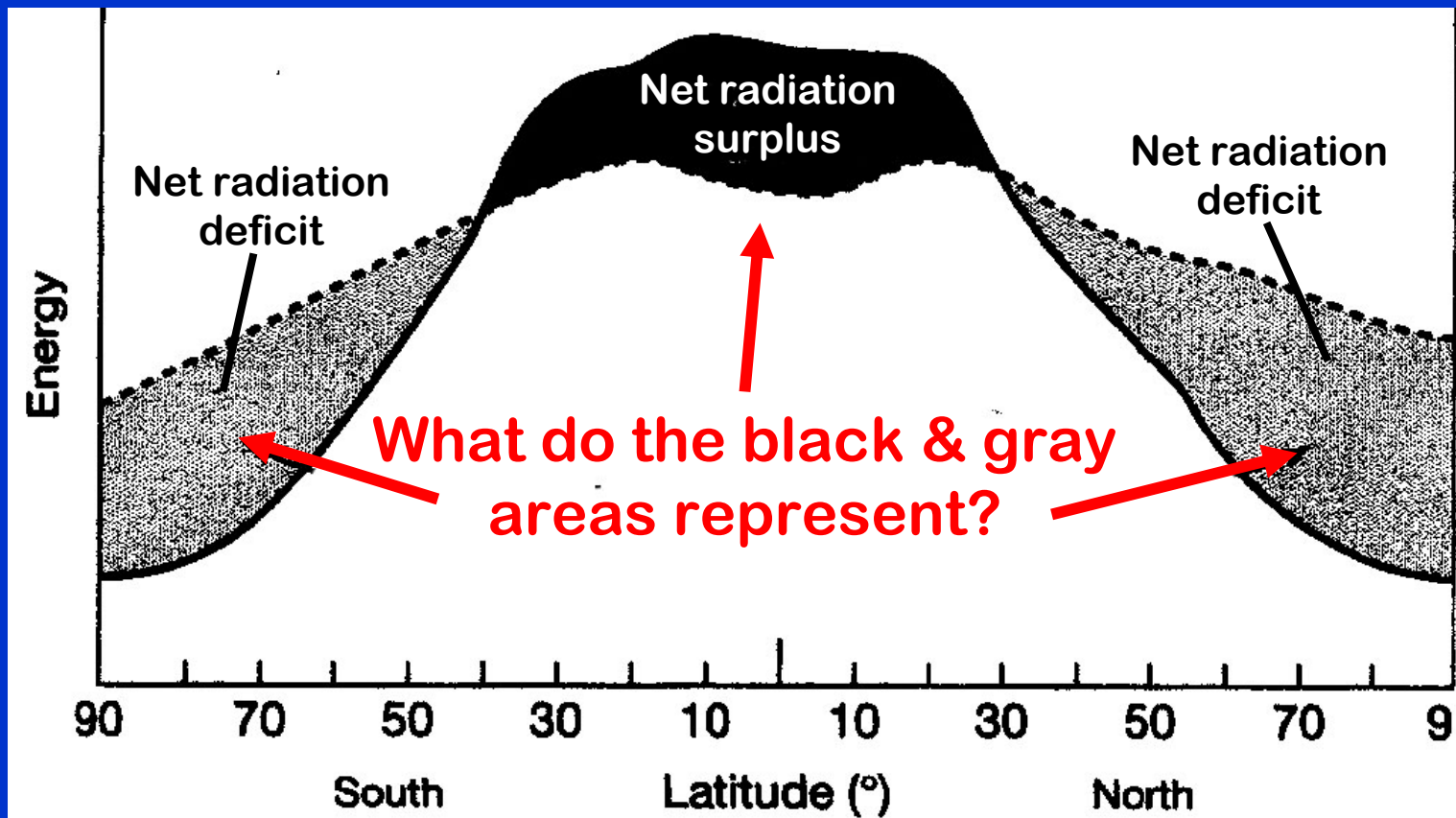
The EARTH radiates out LW fairly evenly from latitude to latitude, but MORE LW energy is radiated out in warm Equatorial latitudes & LESS in cold Polar latitudes

Remember?

$$E = \sigma T^4$$

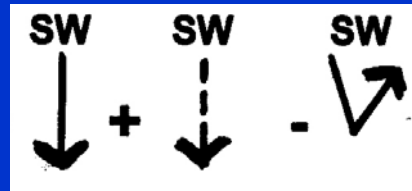


This Figure is on p 80

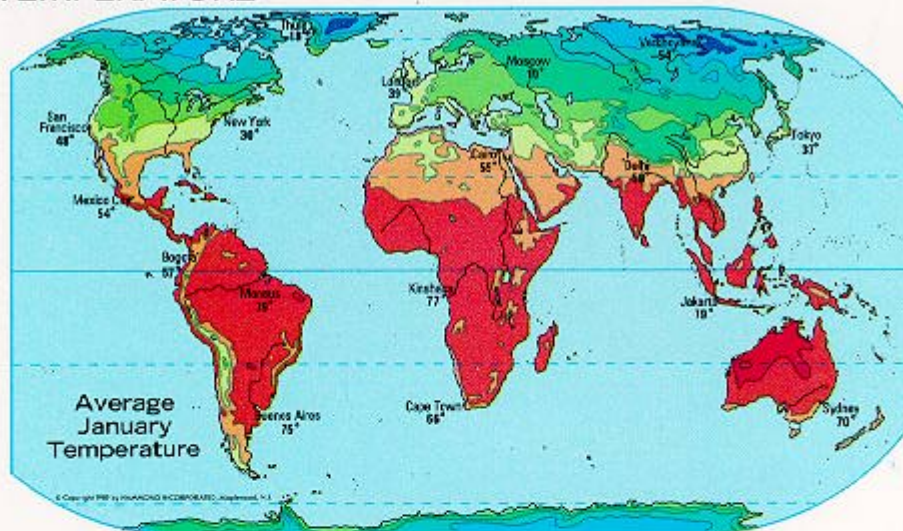


———— Absorbed solar energy

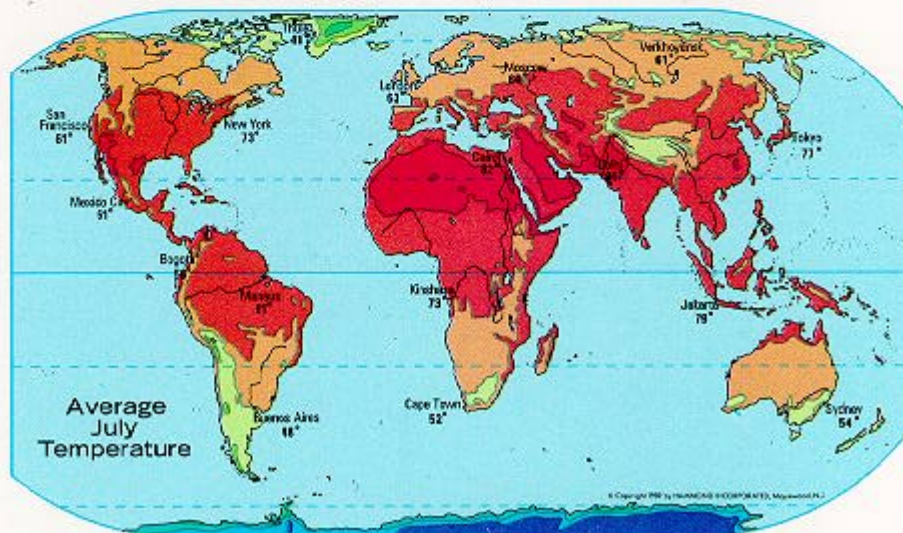
----- Emitted infrared energy
(at top of atmosphere)



WORLD-TEMPERATURE



FAHRENHEIT	CELSIUS	FAHRENHEIT	CELSIUS	FAHRENHEIT	CELSIUS	
Over 85°	Over 30°	32° to 50°	0° to 10°	-22° to -4°	-30° to -20°	London 35°
68° to 85°	20° to 30°	14° to 32°	-10° to 3°	-40° to -22°	-40° to -30°	Average temperature at selected stations
50° to 68°	10° to 20°	-4° to 14°	-20° to -13°	Under -40°	Under -40°	



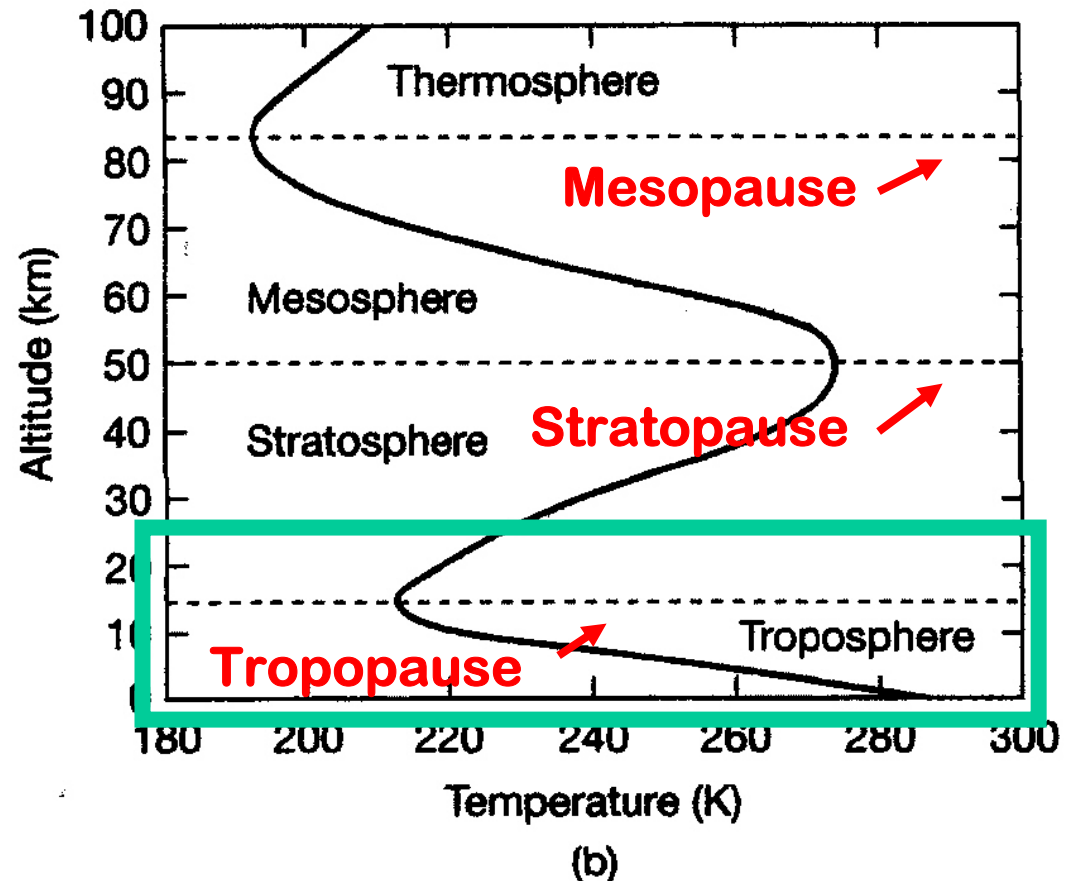
IGC CD MODULE SELECT

→ “Weather in Motion” – Ch 2
January & July
1999 Global Movie of Land & Sea
Surface
Temperatures &
Clouds

Recall: Vertical Structure of the Atmosphere

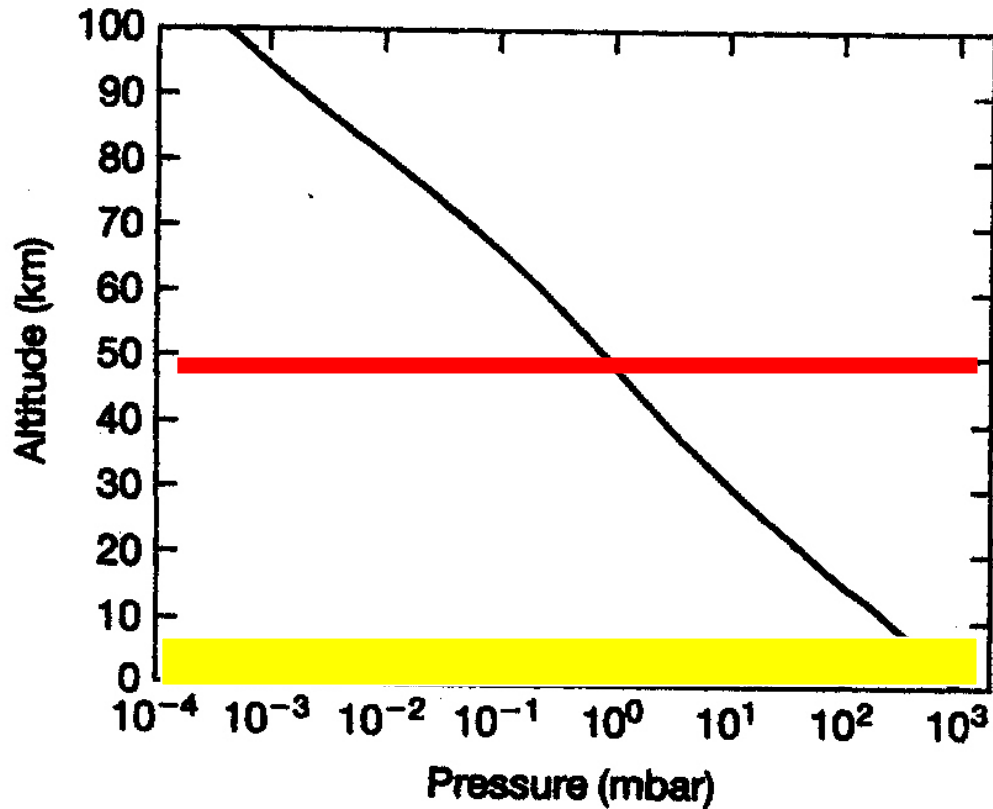
defined by
CHANGES in
TEMPERATURE
with height.

**MOST WEATHER
& CLIMATE
PROCESSES take
place in the
TROPOSPHERE (&
some in
Stratosphere)**



ATMOSPHERIC PRESSURE

Atmospheric Pressure & Mass Vary with Height

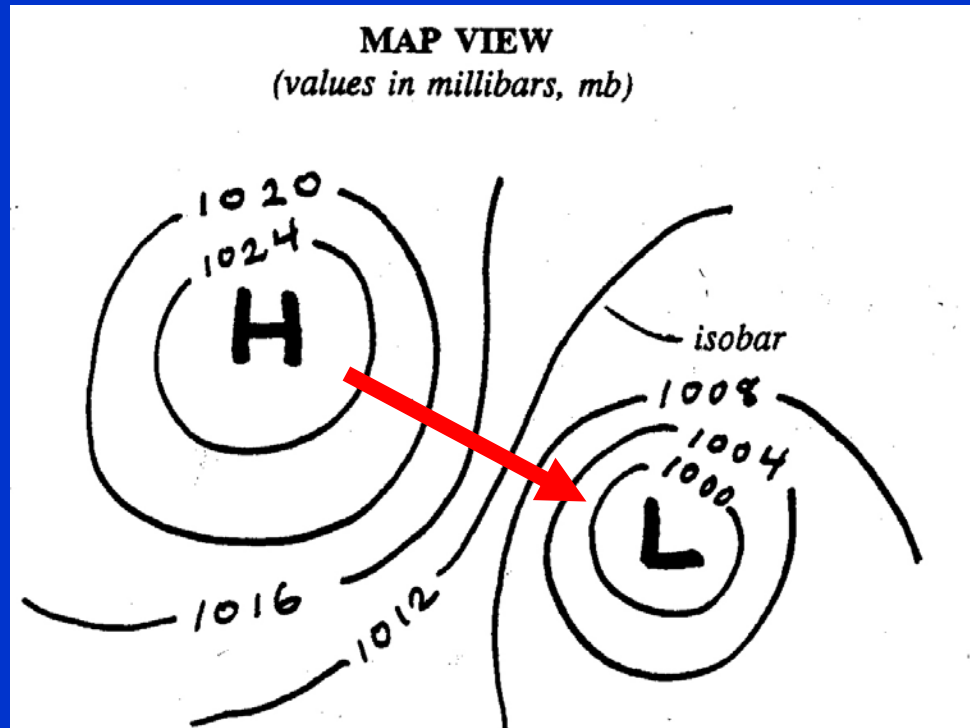


99% of mass lies below ~ 50 km (top of Stratosphere)

50% of mass lies below ~ 6 km (middle Troposphere)

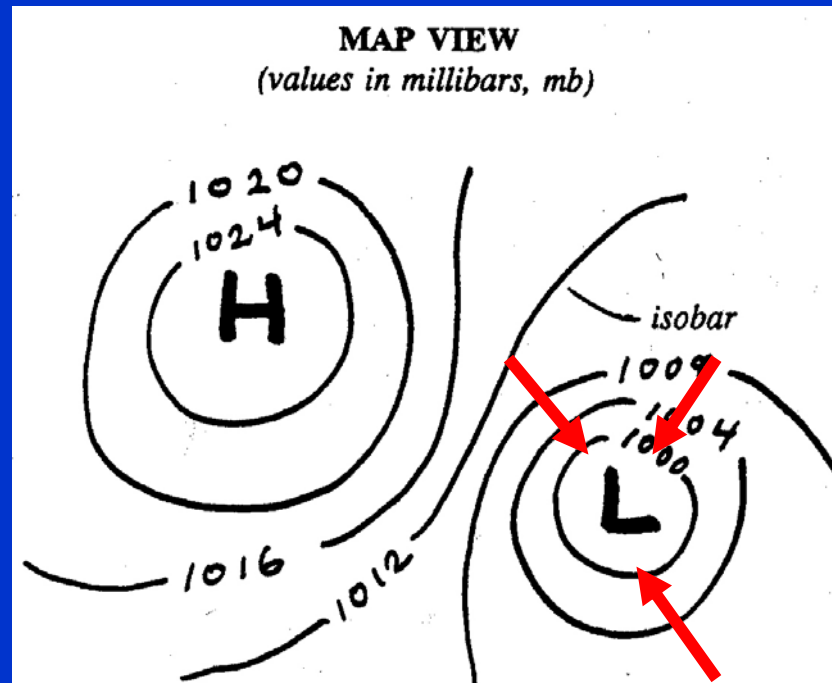
Def: the weight of the air above a given point or level.

In general: Winds tend to flow from
HIGH → LOW Pressure areas

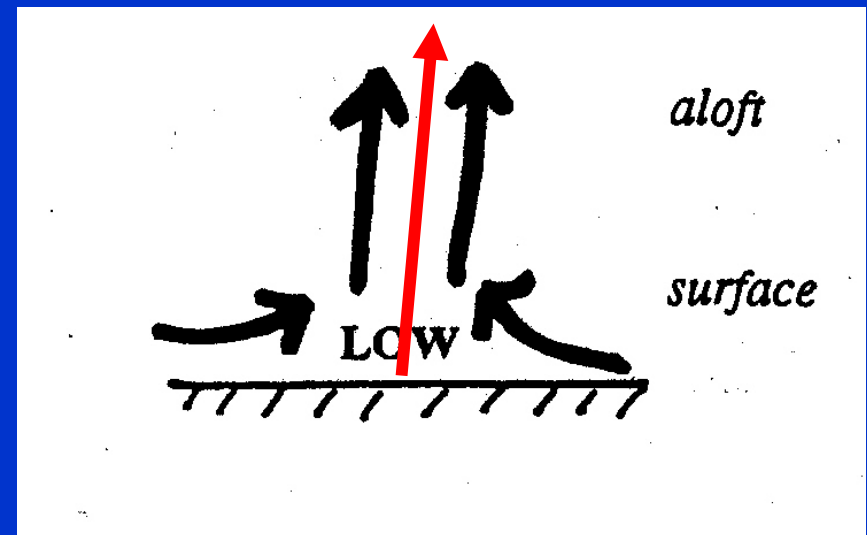


Areas or centers of high and low pressure at the Earth's surface have the following characteristics:

LOWS - air converges into lows and rises in the center of lows

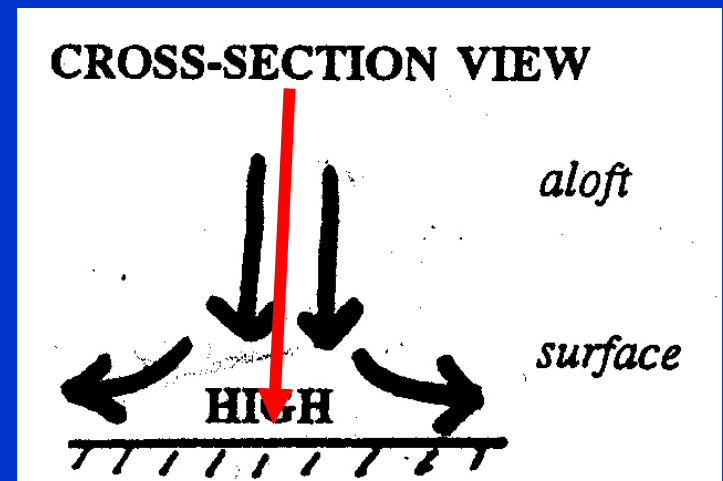
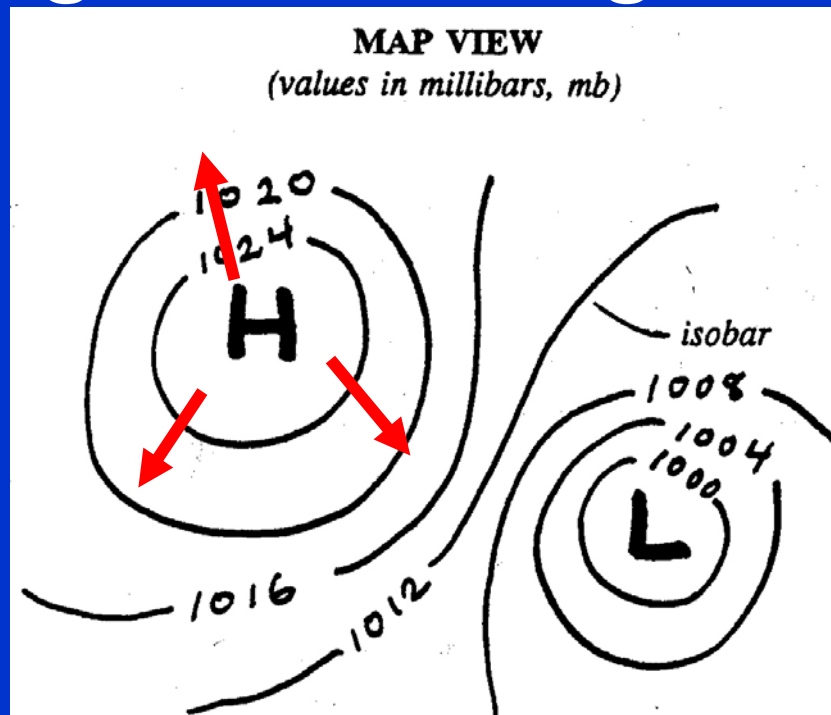


CROSS-SECTION VIEW

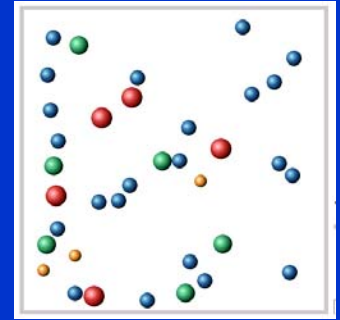


Areas or centers of high and low pressure at the Earth's surface have the following characteristics:

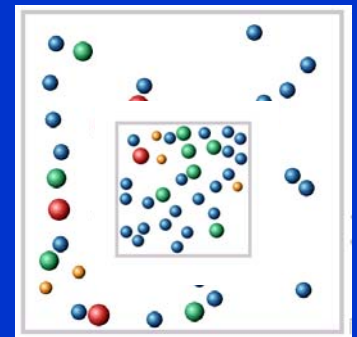
HIGHS - air subsides in the center of highs and diverges out of highs

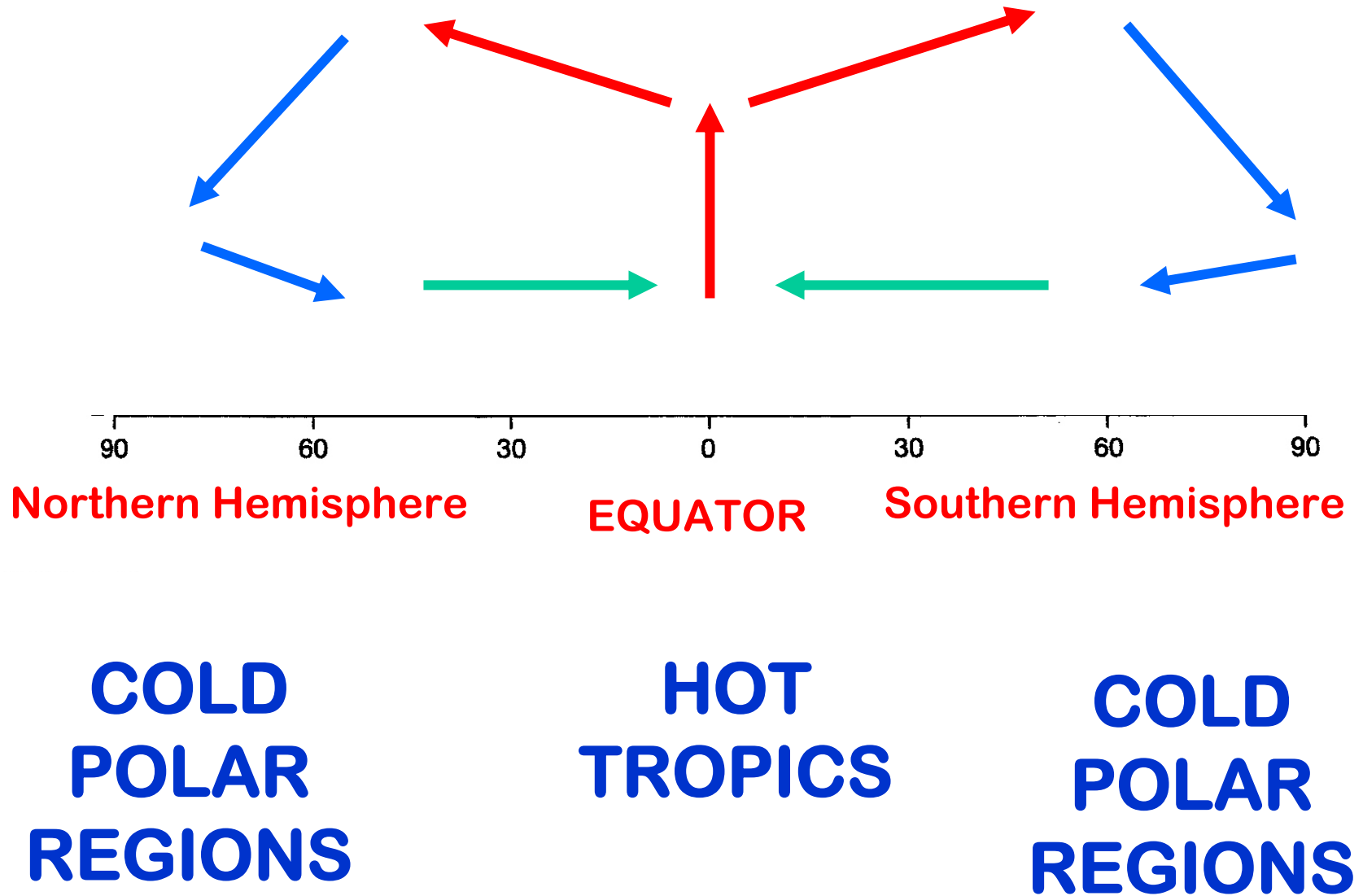


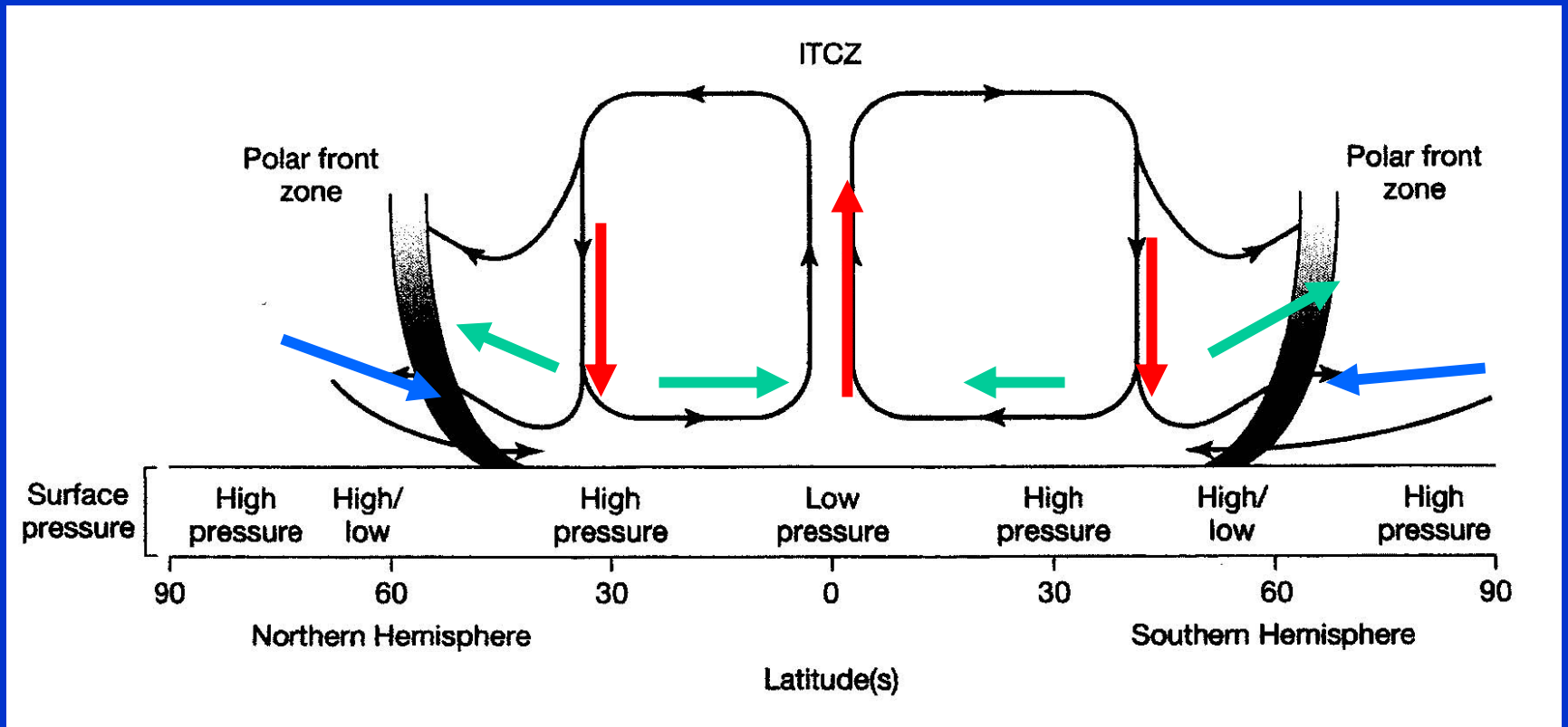
LOWS: (rising leads to **expansion and cooling** of air, and condensation of water vapor ==> clouds and possibly precipitation)



HIGHS: (sinking leads to **contraction and warming of air**, and increased water vapor holding capacity ==> clear skies, dry air)



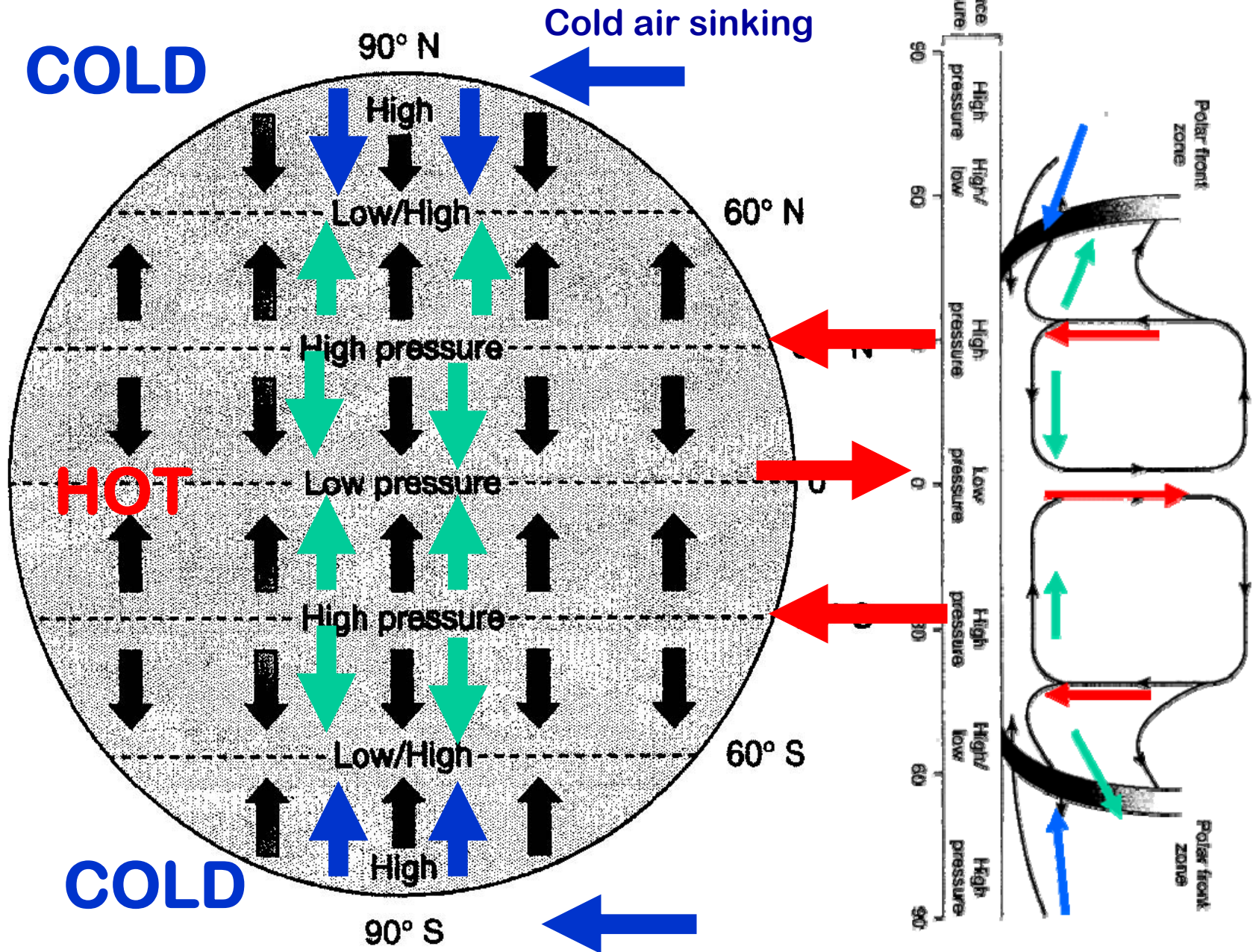




**COLD
POLAR
REGIONS**

**HOT
TROPICS**

**COLD
POLAR
REGIONS**



**Lecture ended here and the
class then got in groups to
finish up the
I-4 BCP Tree-Ring activity.**