

# NATURAL CLIMATIC FORCING

Earth-Sun orbital relationships, changing land-sea distribution (due to plate tectonics), solar variability & **VOLCANIC ERUPTIONS**

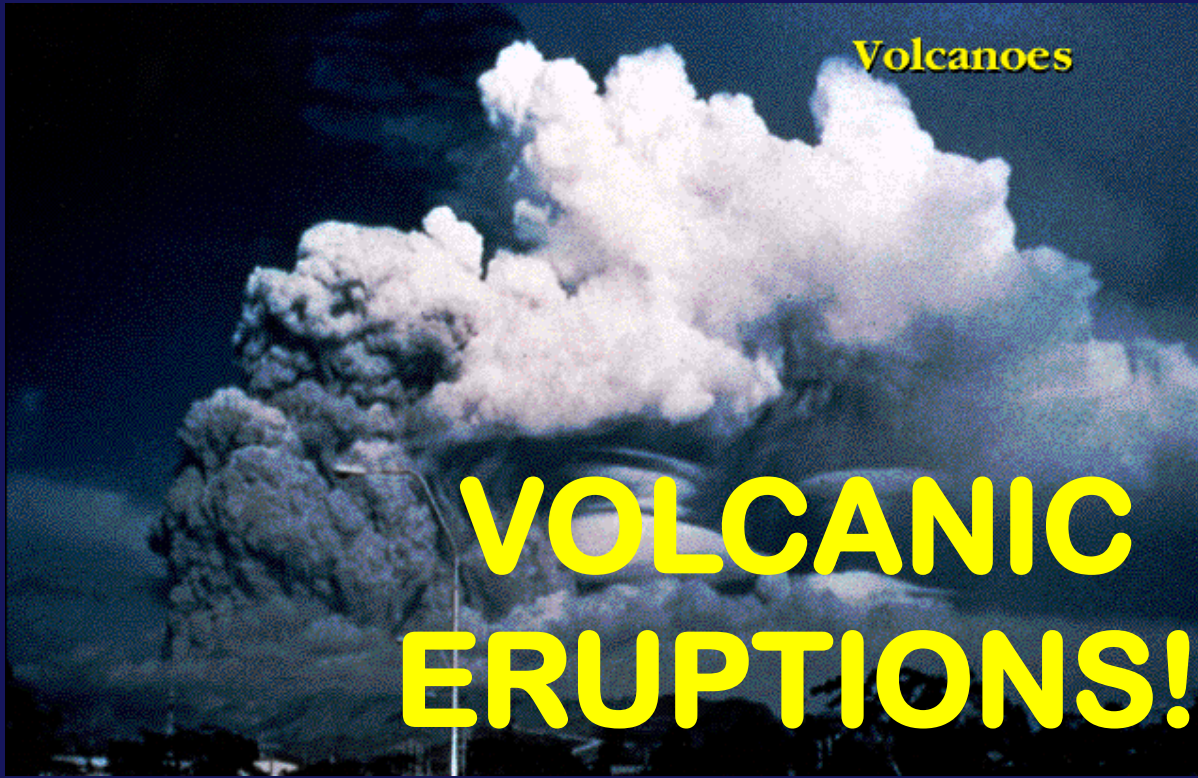
vs.

# ANTHROPOGENIC FORCING

Human-Enhanced GH Effect,  
contribution of catalysts for **OZONE DEPLETION**

Volcanoes

# VOLCANIC ERUPTIONS!



**Volcanoes are one way the  
Earth gives birth to itself.**

**~Robert Gross**

**Volcanic eruptions  
contribute to the natural  
Greenhouse Effect by  
adding CO<sub>2</sub> into the  
atmosphere:**

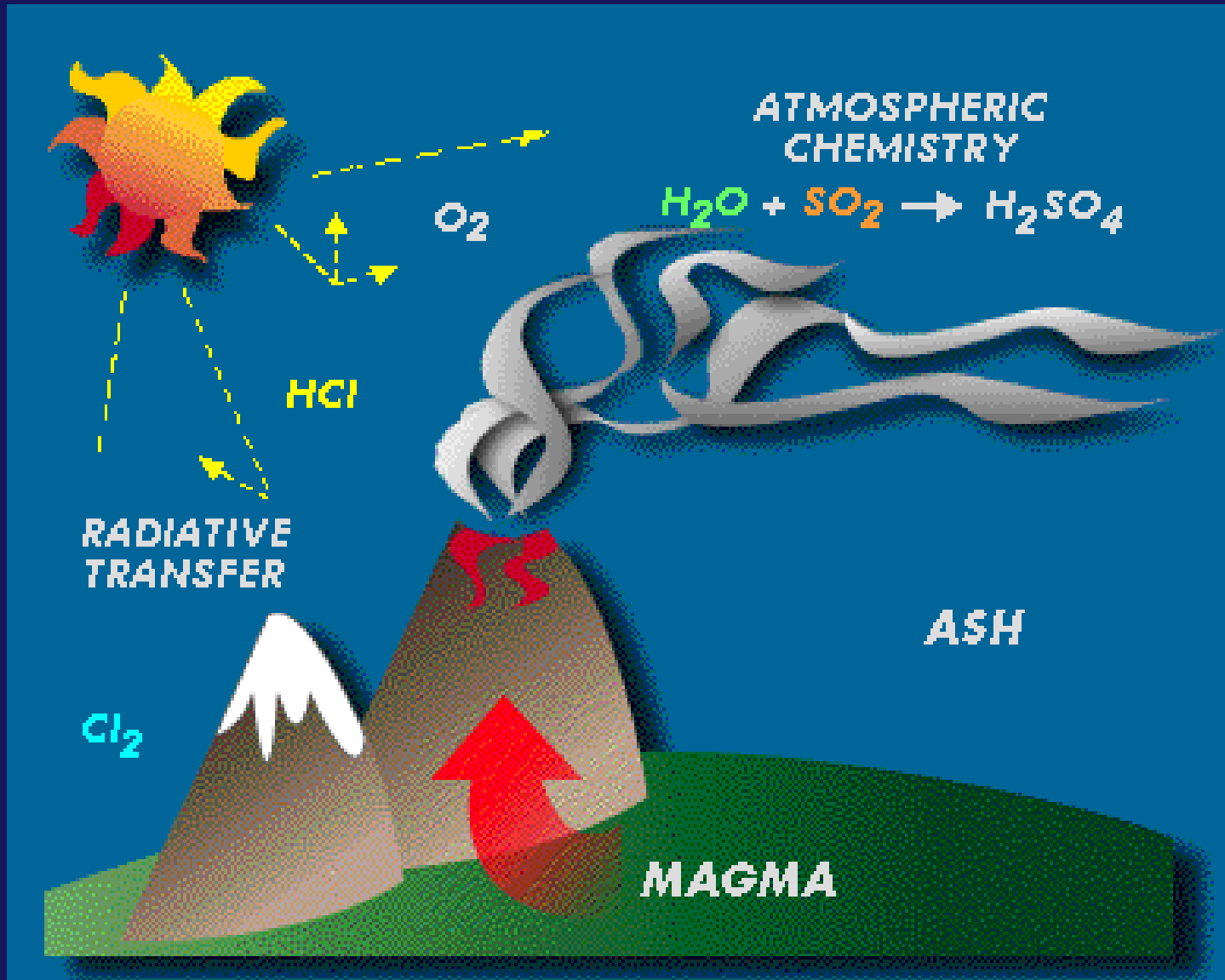
**Volcanic outgassing  
of CO<sub>2</sub>  
into atmosphere**

**0.06 Gtons**

**This carbon flux  
is more or less  
in balance over  
time**

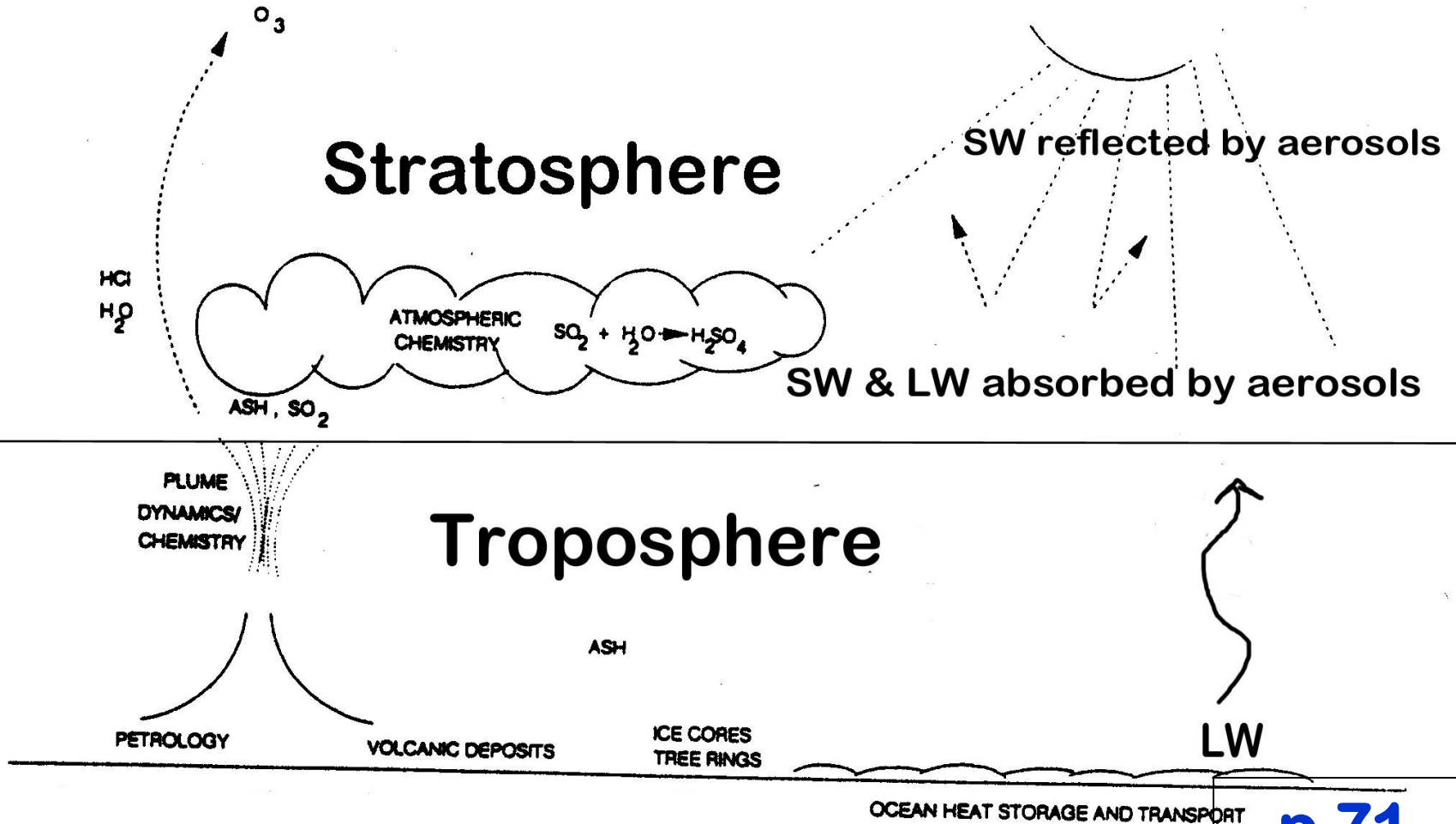


Eruptions can also have a more direct climatic effect under certain conditions:



# How the Climatic Effect Occurs . . . . through **the ENERGY BALANCE** of course!

ozone destruction hastened by chemical reactions on aerosol surfaces



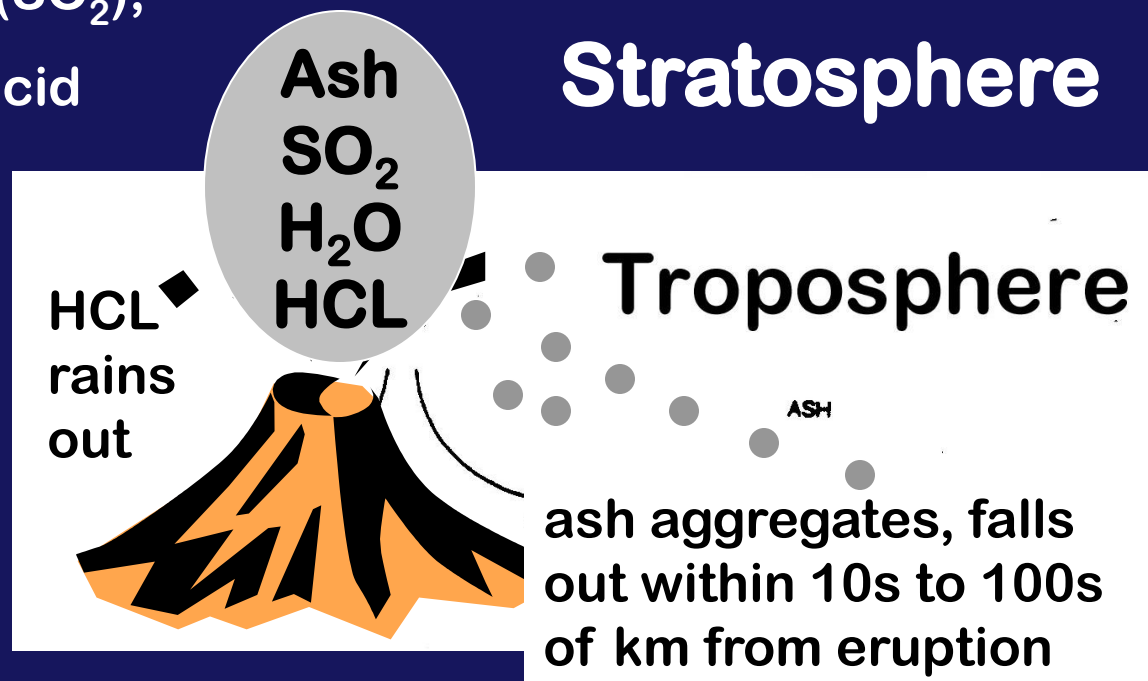
# Large volcanic eruptions inject sulfur gases, water vapor, HCL into the stratosphere:

water vapor ( $\text{H}_2\text{O}$ )

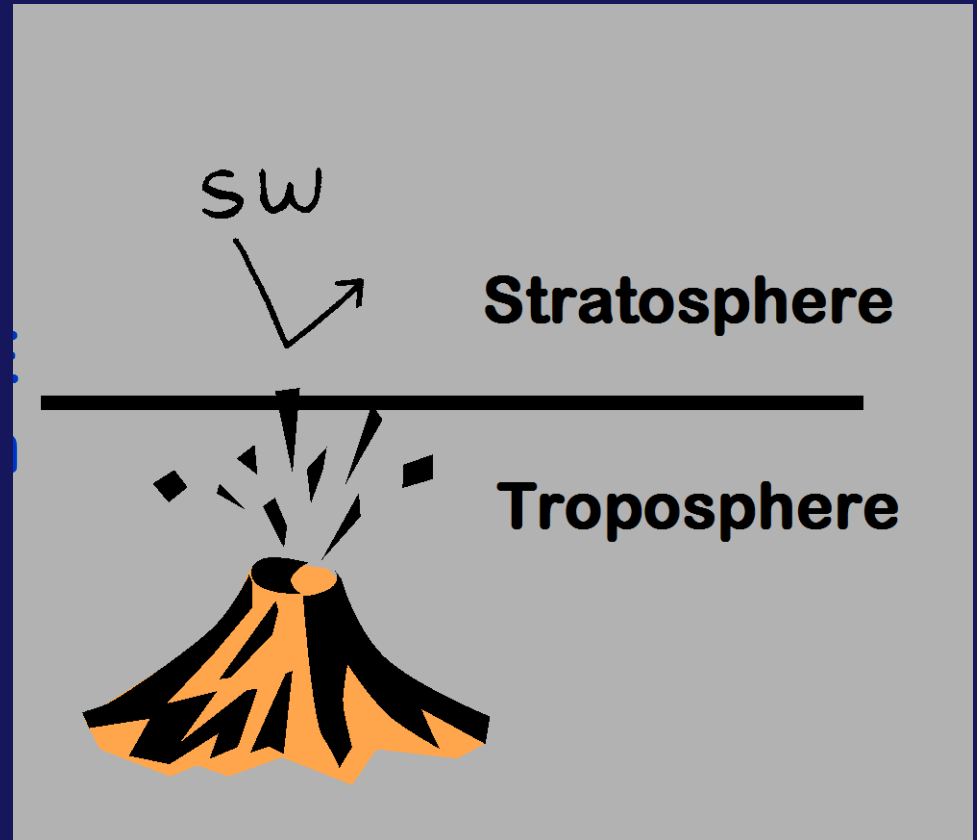
sulfur dioxide ( $\text{SO}_2$ ),

hydrochloric acid  
(HCl)

mineral ash  
into the  
stratosphere



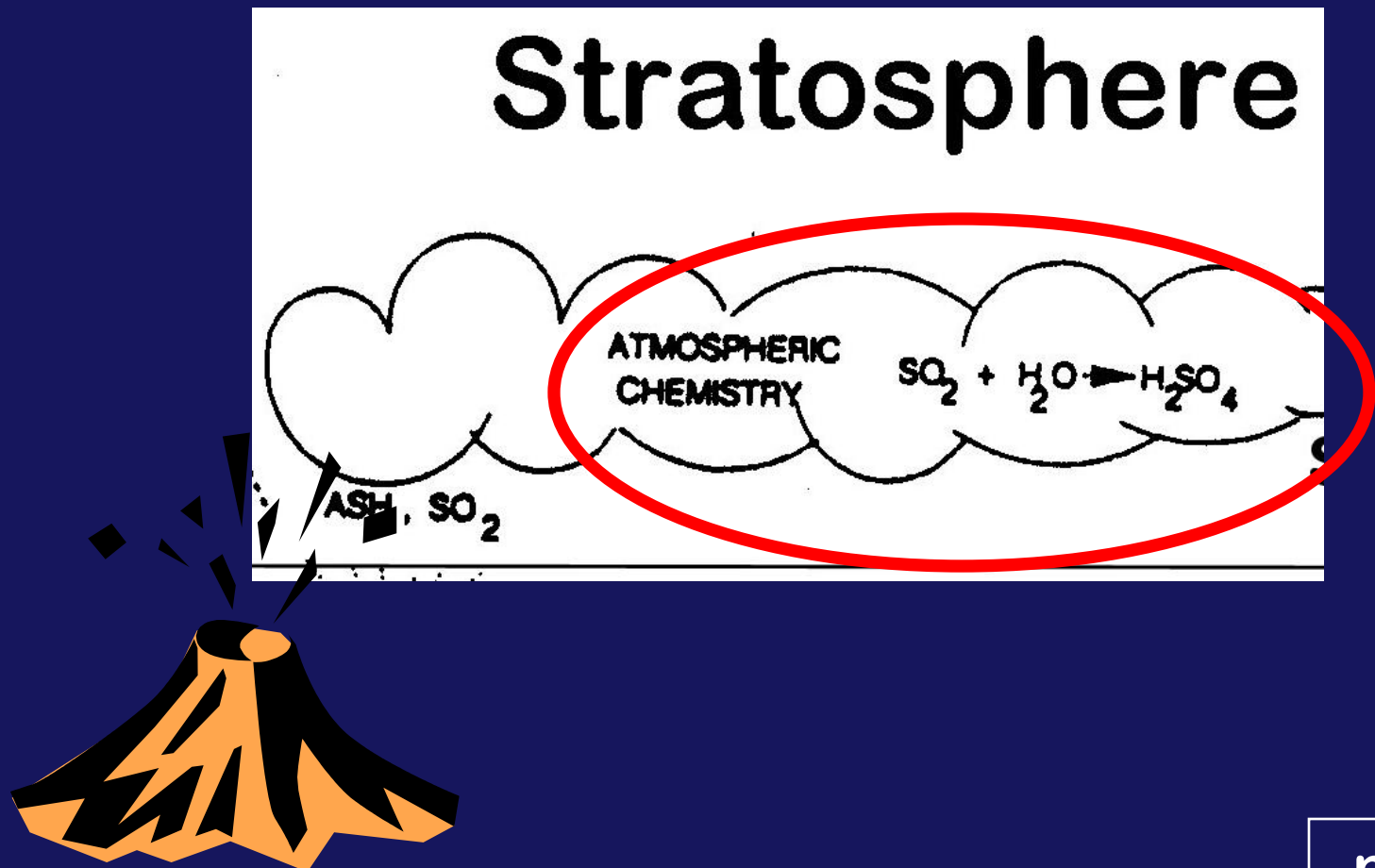
Albedo of ejected  
ASH in the  
**STRATOSPHERE**  
is *not* the reason  
for cooling after  
an eruption!  
(most ash falls out  
early)



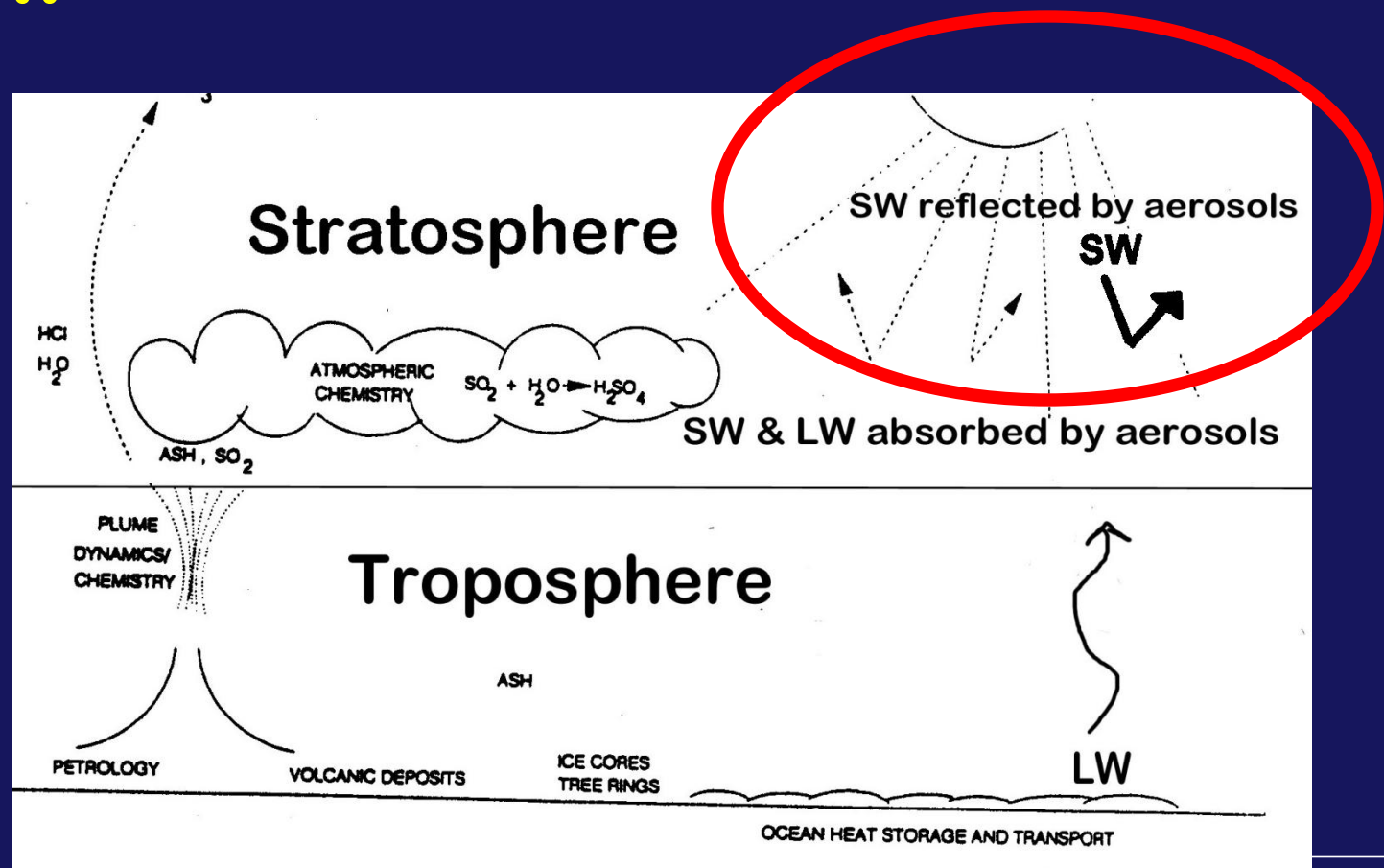
What ***DOES*** reflect the incoming  
shortwave radiation after an  
eruption?



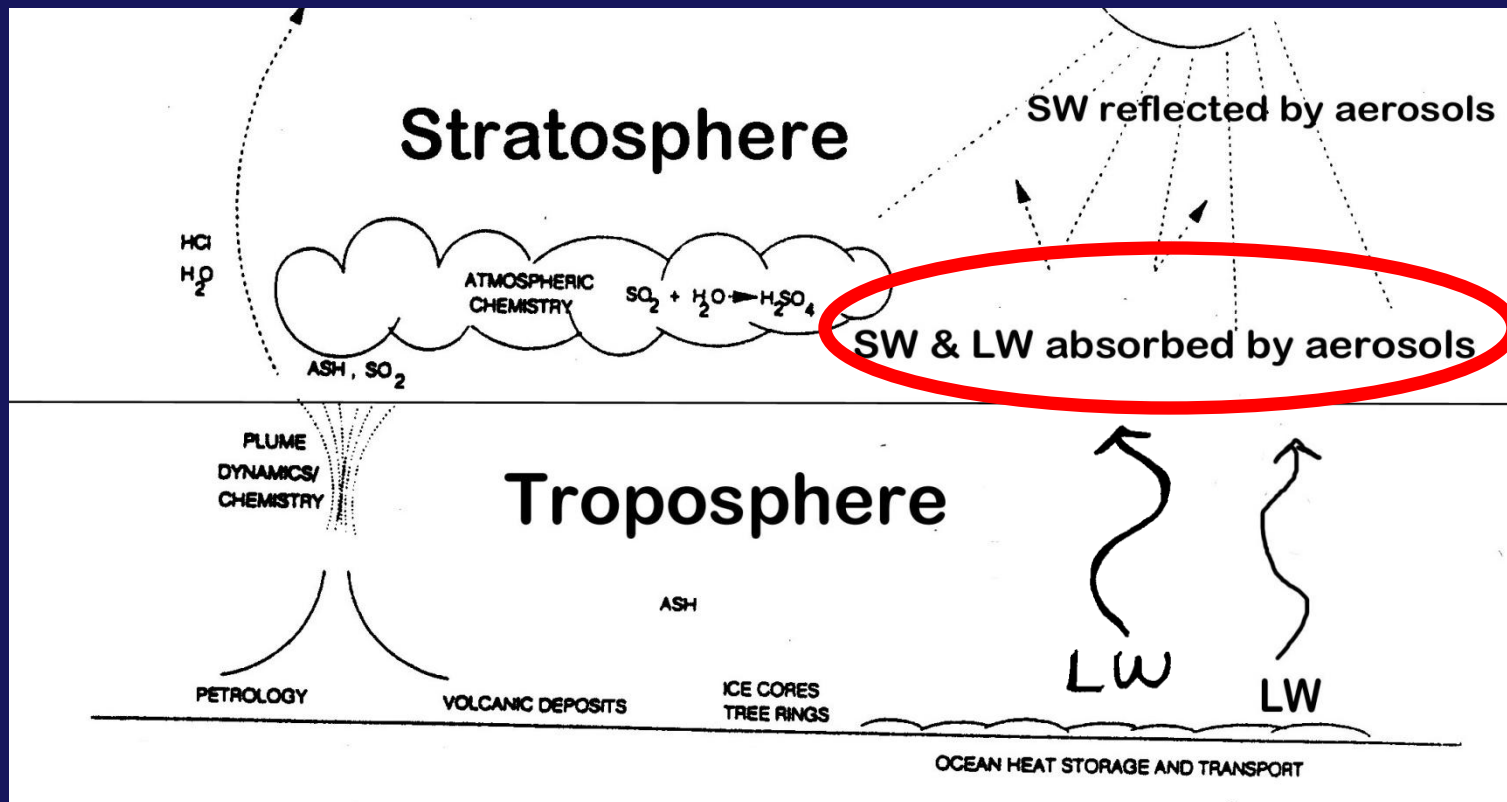
**SO<sub>2</sub>** remains gaseous and is eventually converted to **sulfuric acid** (H<sub>2</sub>SO<sub>4</sub>) which condenses in a mist of fine particles called **sulfate aerosols**.



the sulfate aerosols *reflect* some of the incoming solar SW radiation back to space, **cooling the troposphere below**

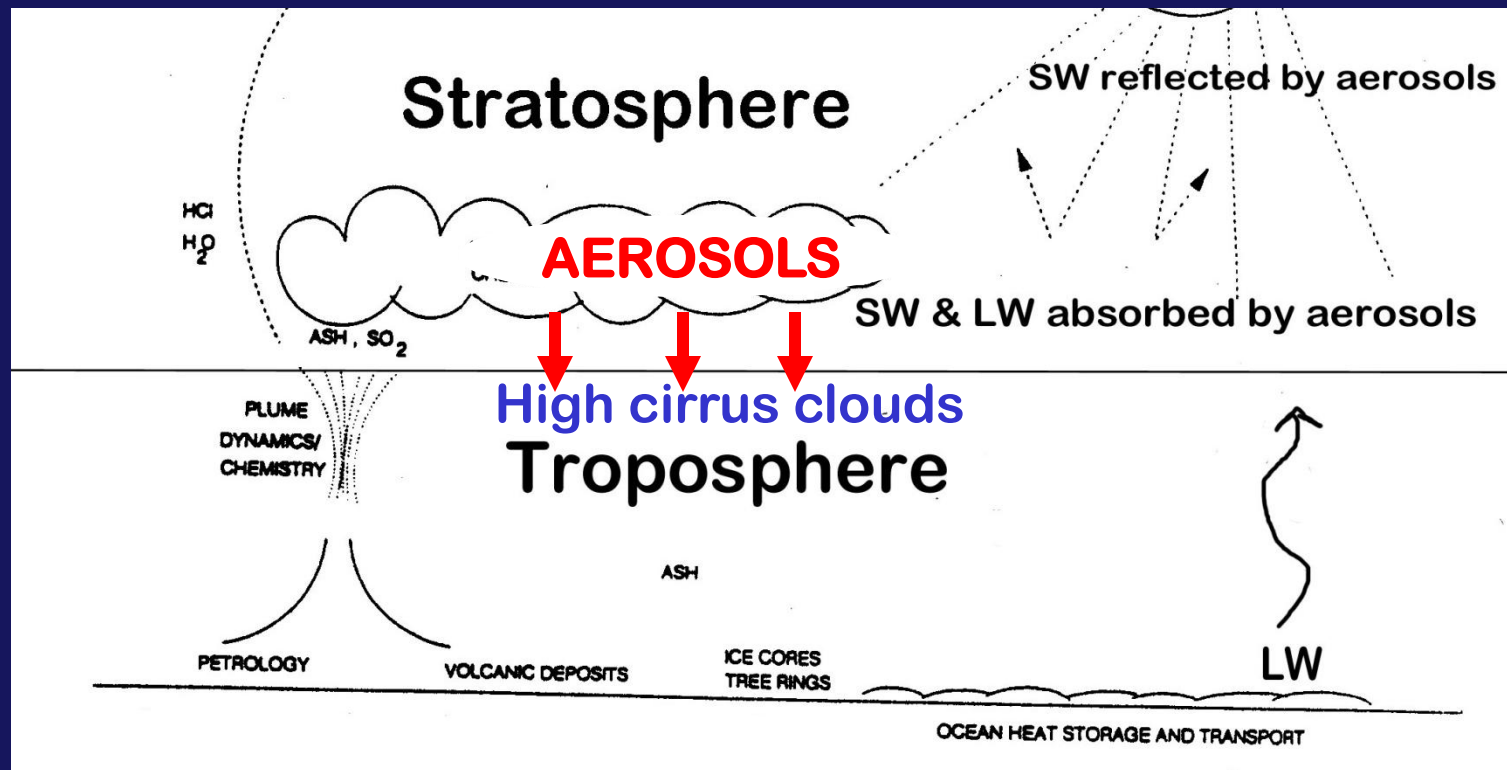


BUT - the aerosols also **ABSORB** certain wavelengths of the incoming SW radiation and some of the Earth's outgoing LW radiation, this **warms the stratosphere** (not the troposphere)

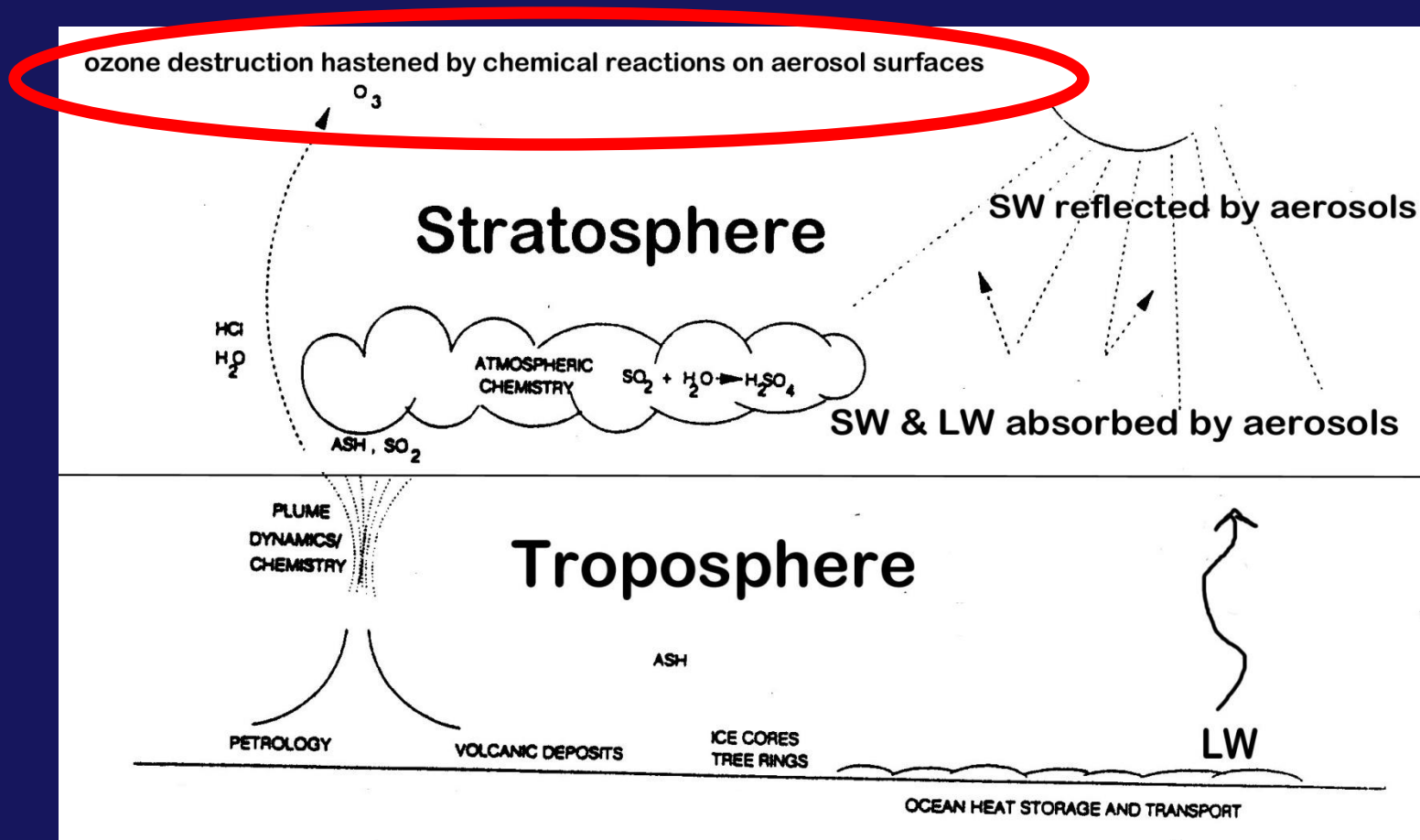


Then, as the **aerosols settle** into the upper troposphere, they may serve as nuclei for **cirrus (high) clouds**, further affecting the Earth's radiation balance \*

\* either absorbing or reflecting, depending on the cloud's albedo and other factors

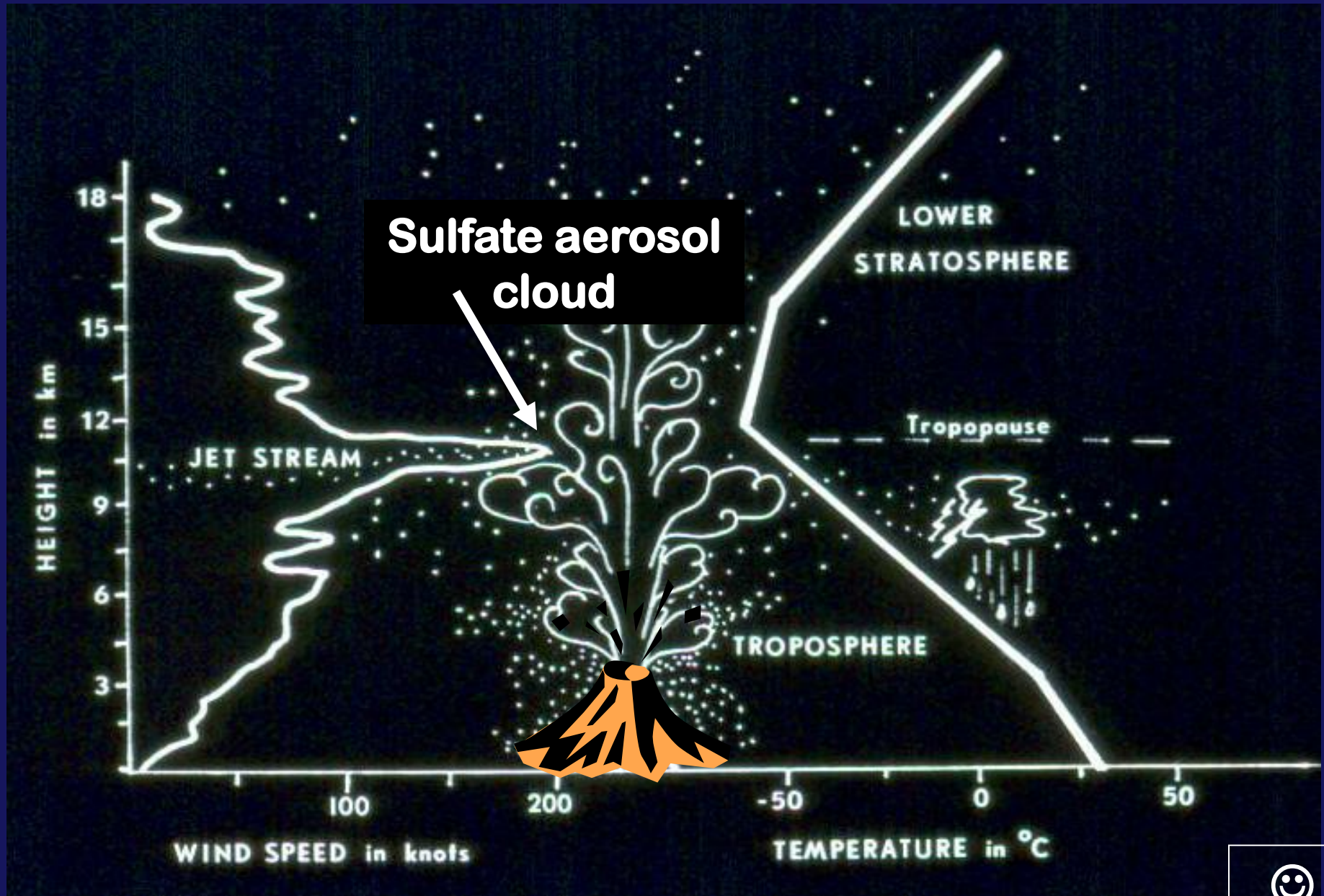


# Chemical effects of the sulfate aerosol cloud can also produce responses in the climate system through **OZONE destruction** (Topic #13)





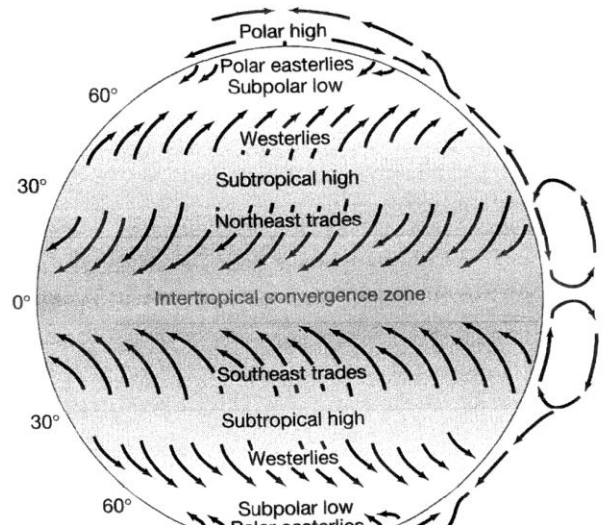
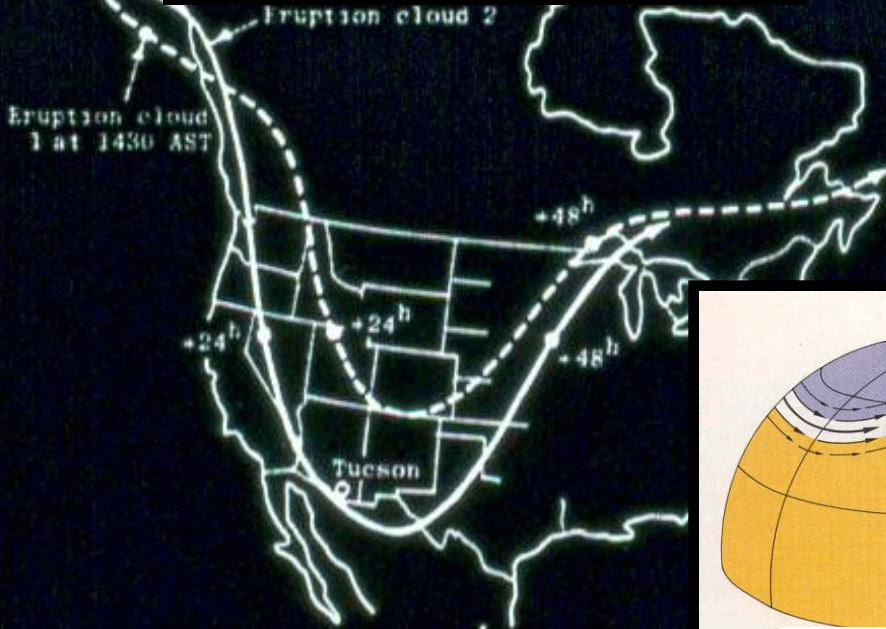
# How do eruption effects become GLOBAL??



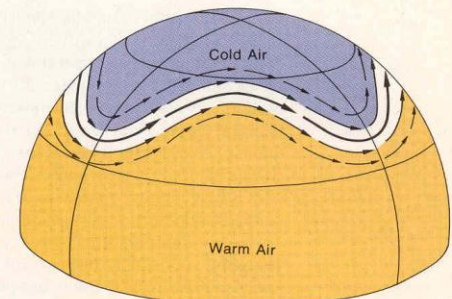
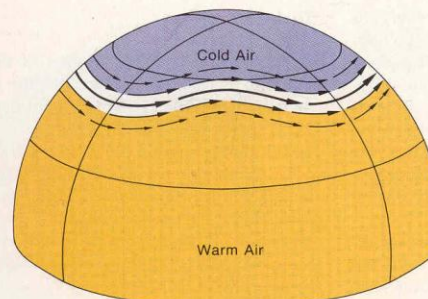


# DUST TRAJECTORIES JAN. 1976

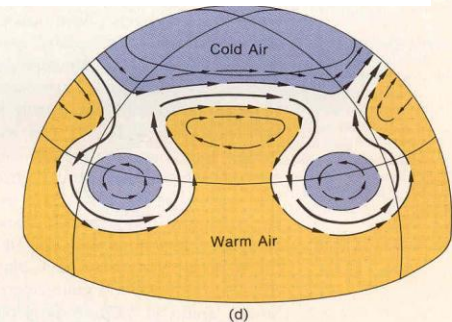
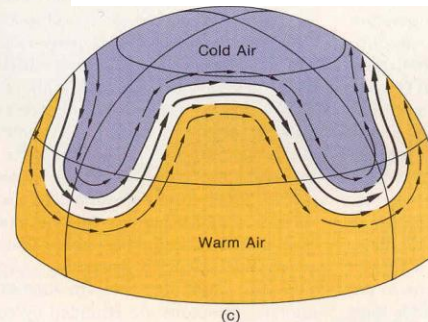
**Mt. St. Augustine eruption**



**Surface wind circulation**



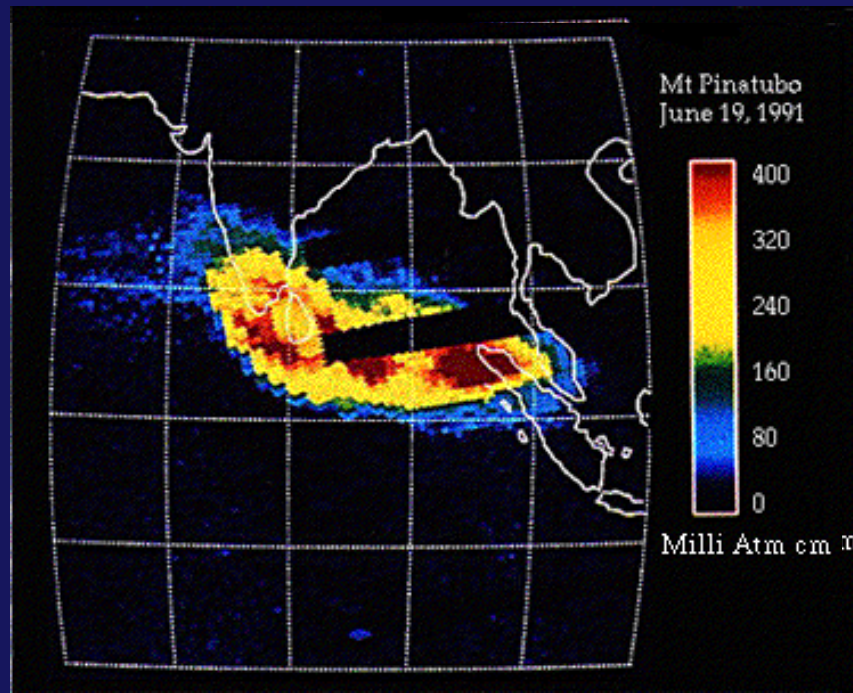
**Upper level wind circulation**



**Through the  
atmospheric  
circulation!**



# Mt Pinatubo Eruption in the Philippines, June, 1991

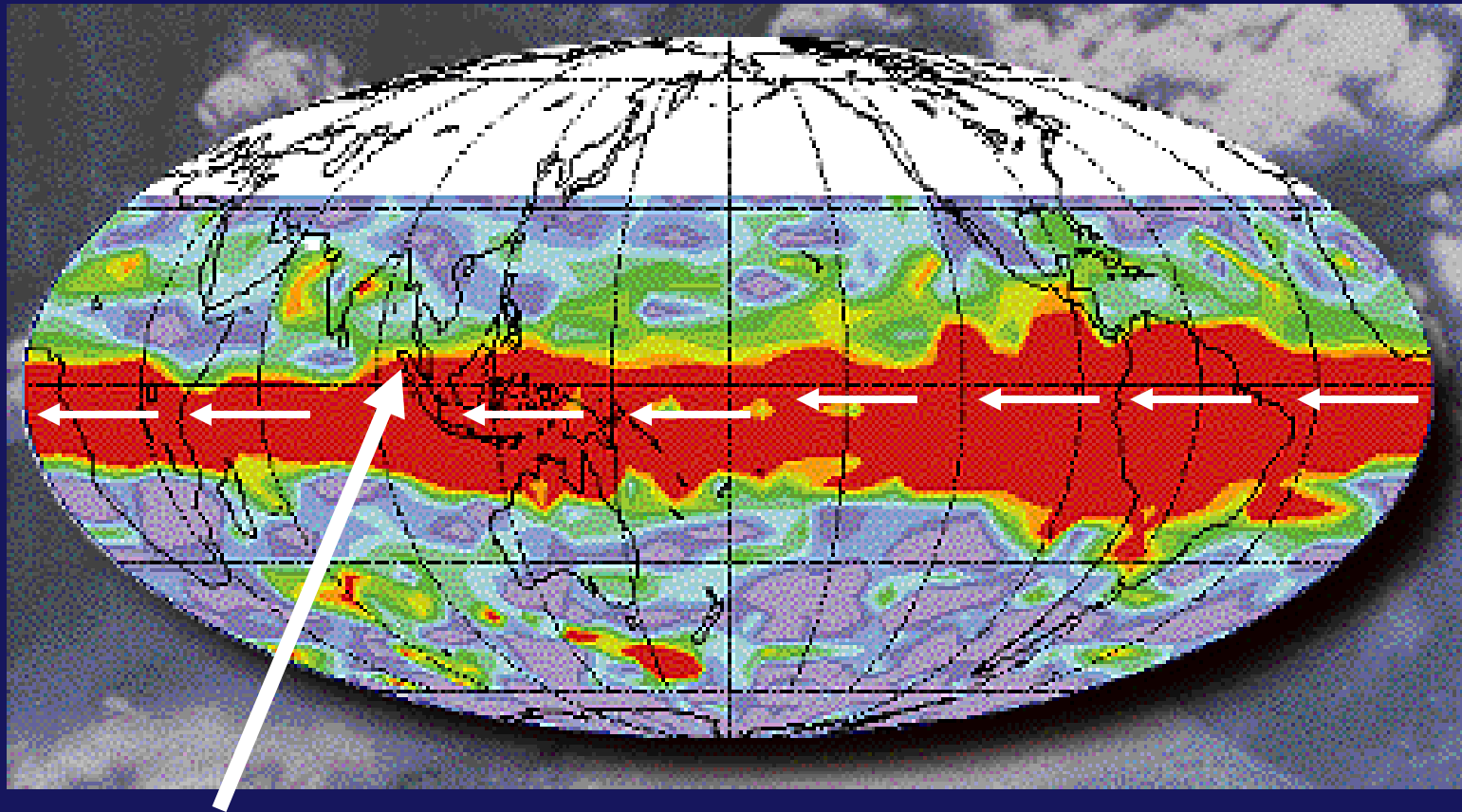


Satellite-derived image of  
sulfur dioxide thickness in the atmosphere  
red = higher thickness





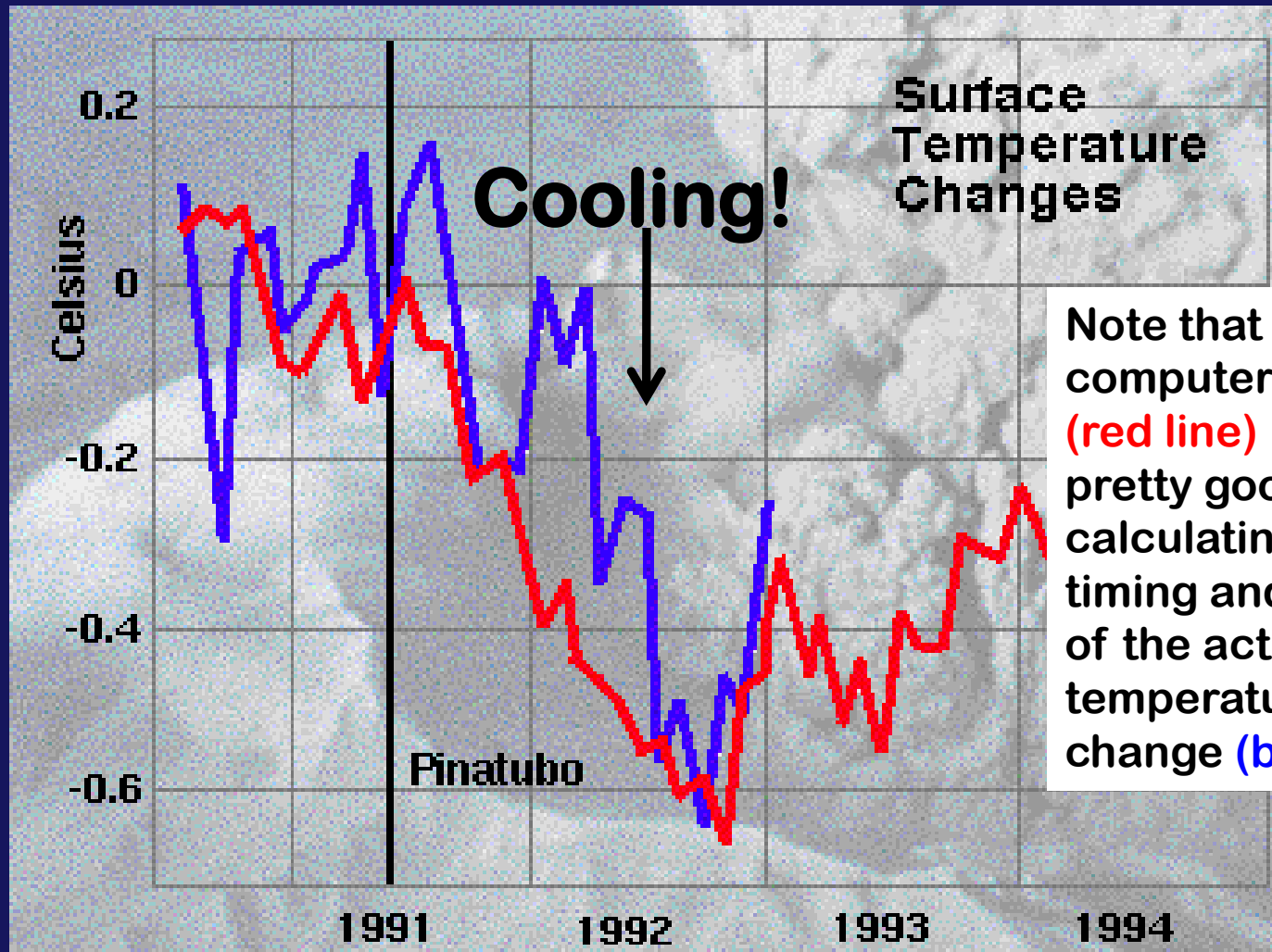
**By Sept 21, 1991 increased levels of sulfur dioxide had dispersed worldwide**



**Mt Pinatubo**



# Mt Pinatubo eruption June 1991

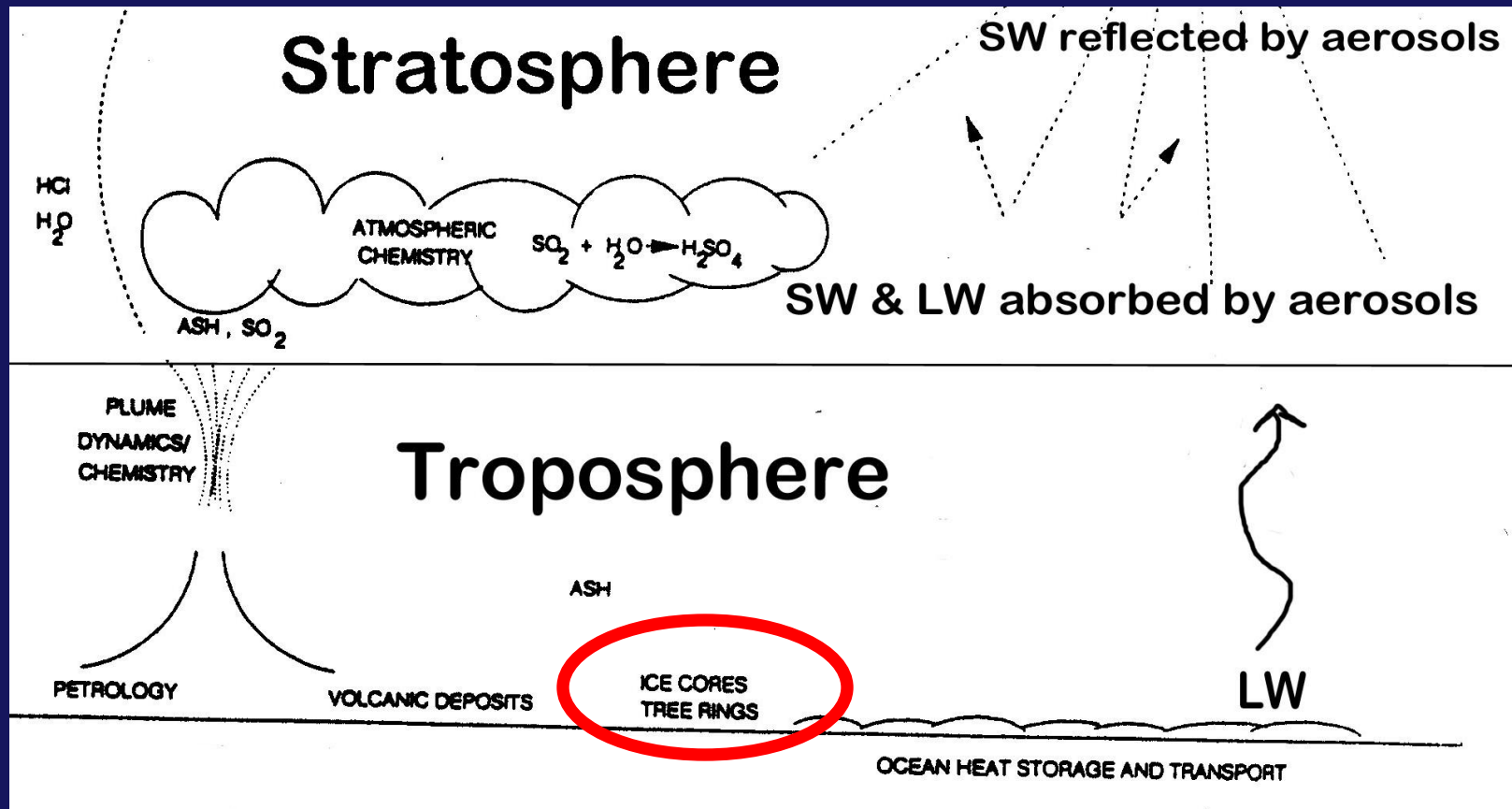


Blue line = observed temperature change after eruption

Red line = modeled temperature change after eruption



Major volcanic eruptions are infrequent events, but their climatic effects can be recorded over long time periods in **ICE CORES & TREE RINGS!**



**Field conditions**

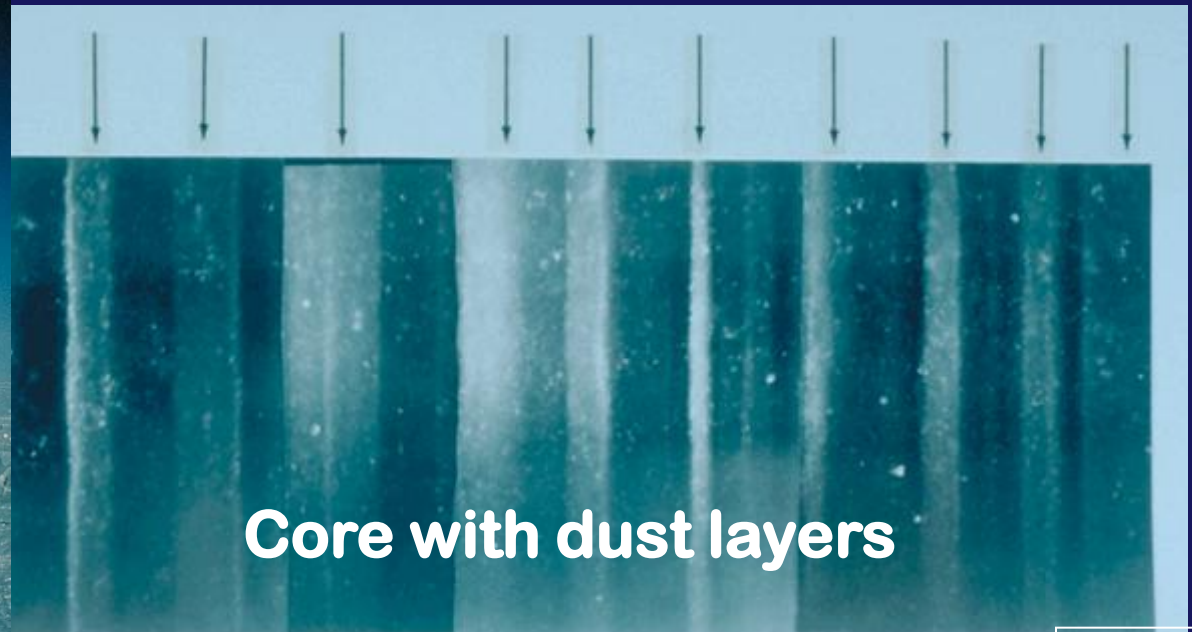
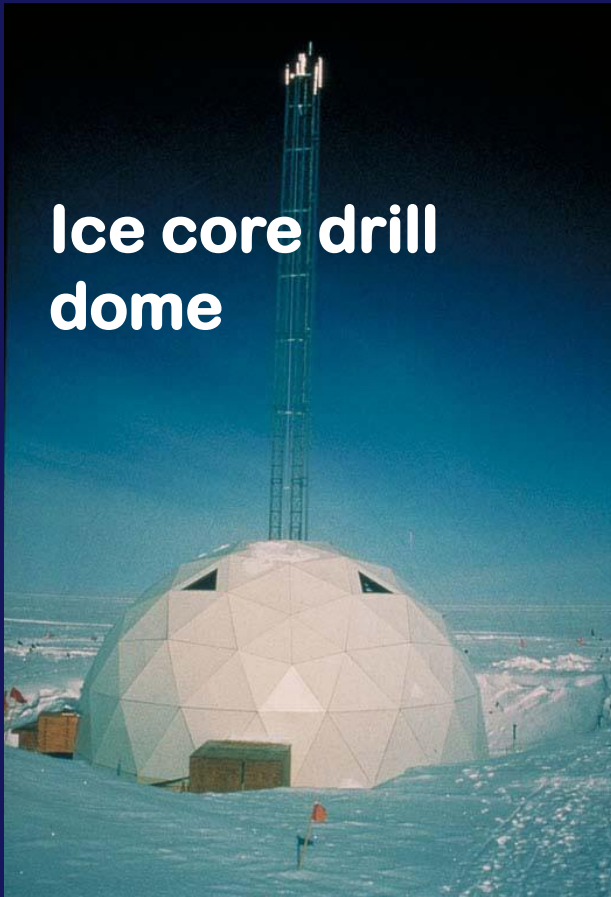


# ICE CORES



**Examining core**

**Ice core drill dome**

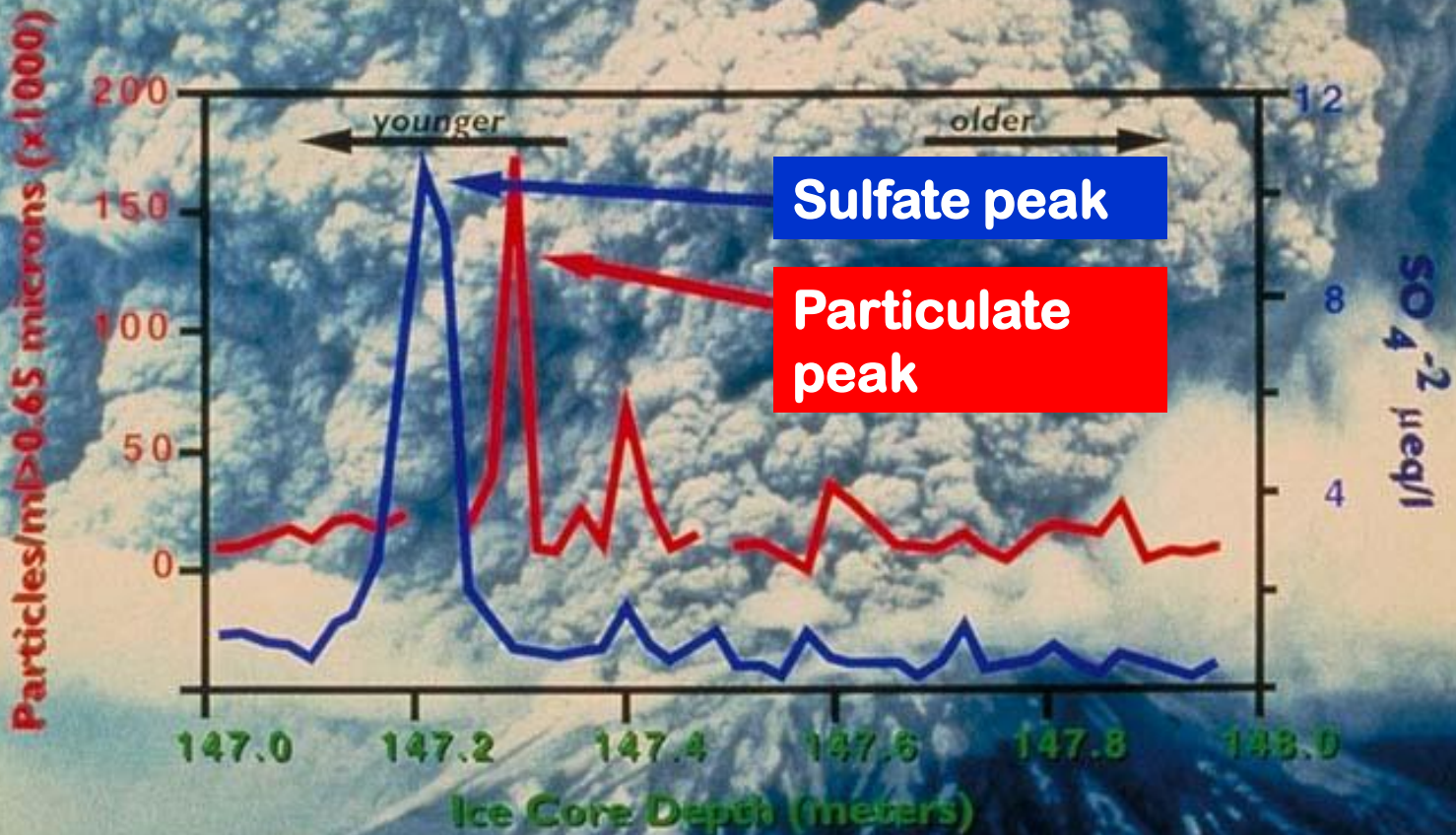


**Core with dust layers**





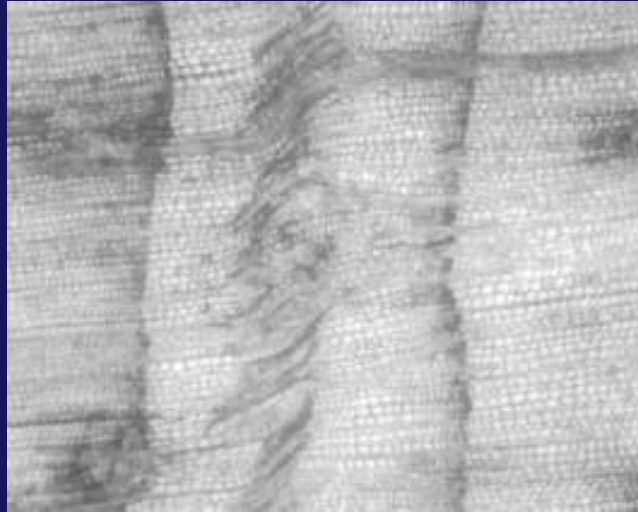
The 1479 A.D. eruption of Mount St. Helens appears as a peak in particle concentration & sulfate records in ice cores from Greenland



Data from Fiacco et al. ( 1993). Photo from Dept. of Natural Resources, Wash. State



# TREE RINGS

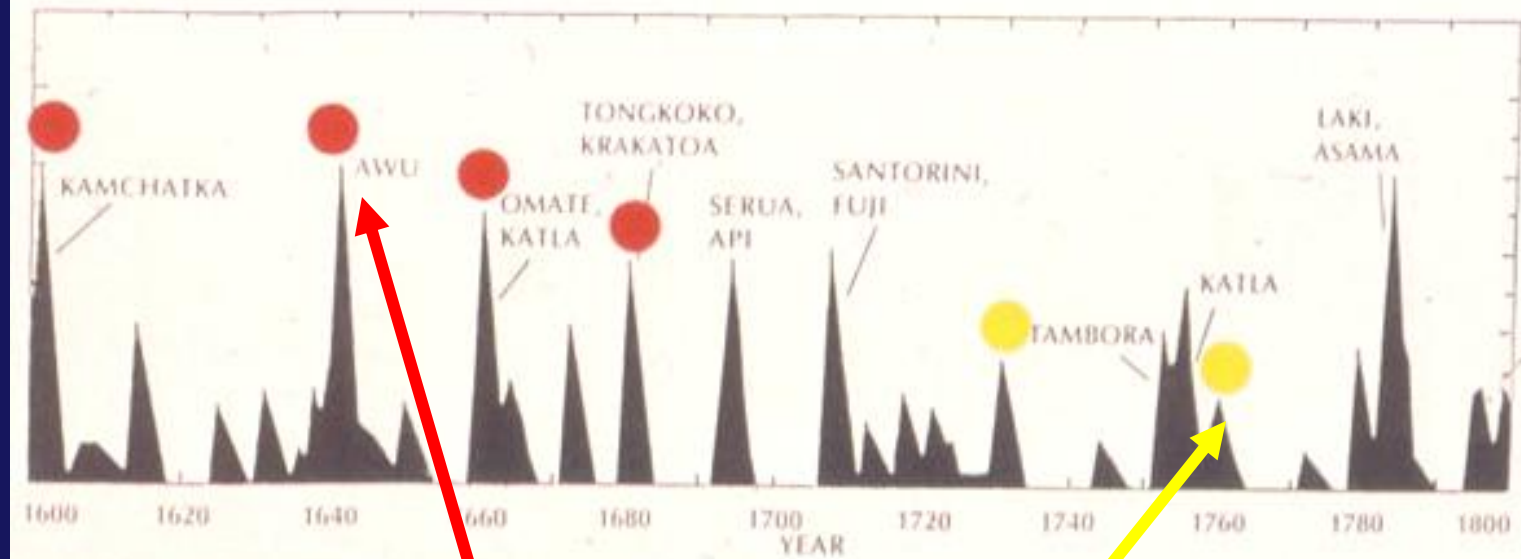


**Eruption – Tree Ring  
Connection via  
FROST RINGS**



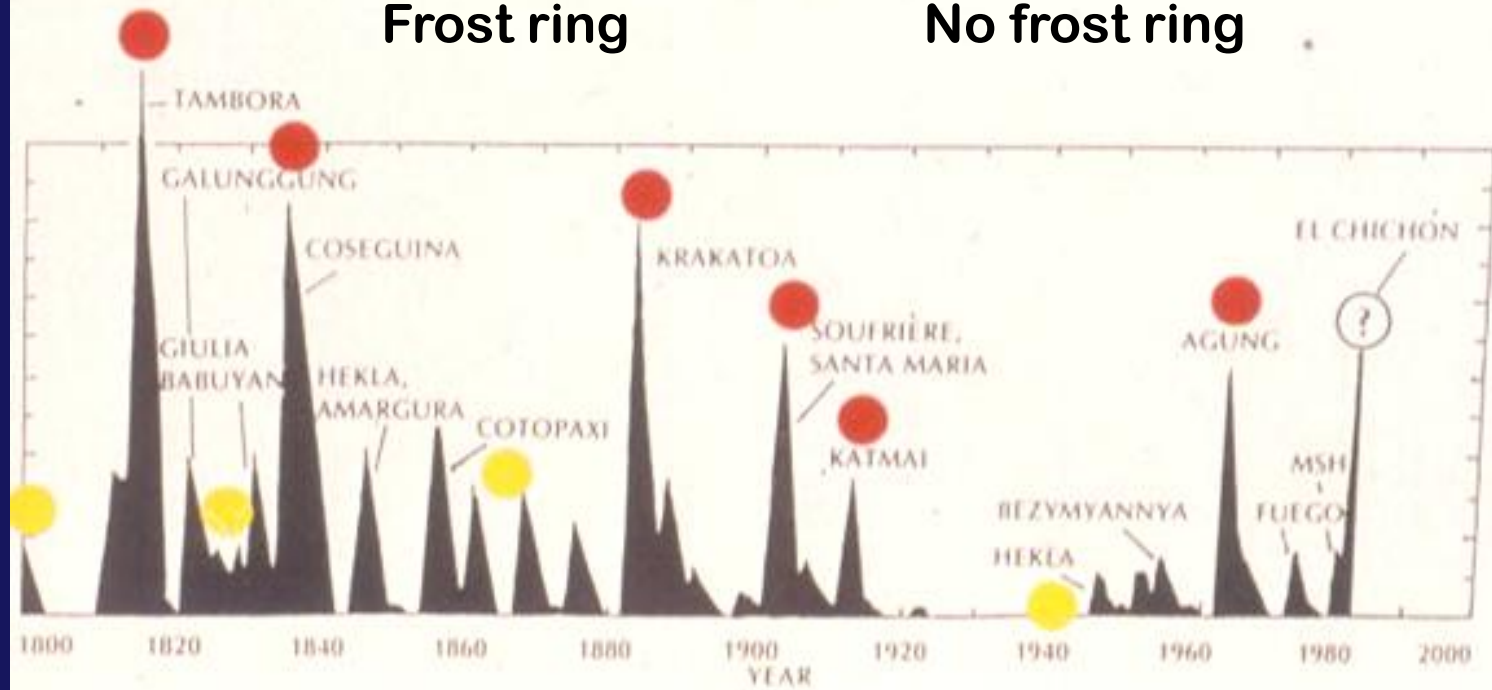


Estimate of Eruption "Dust Veil" →

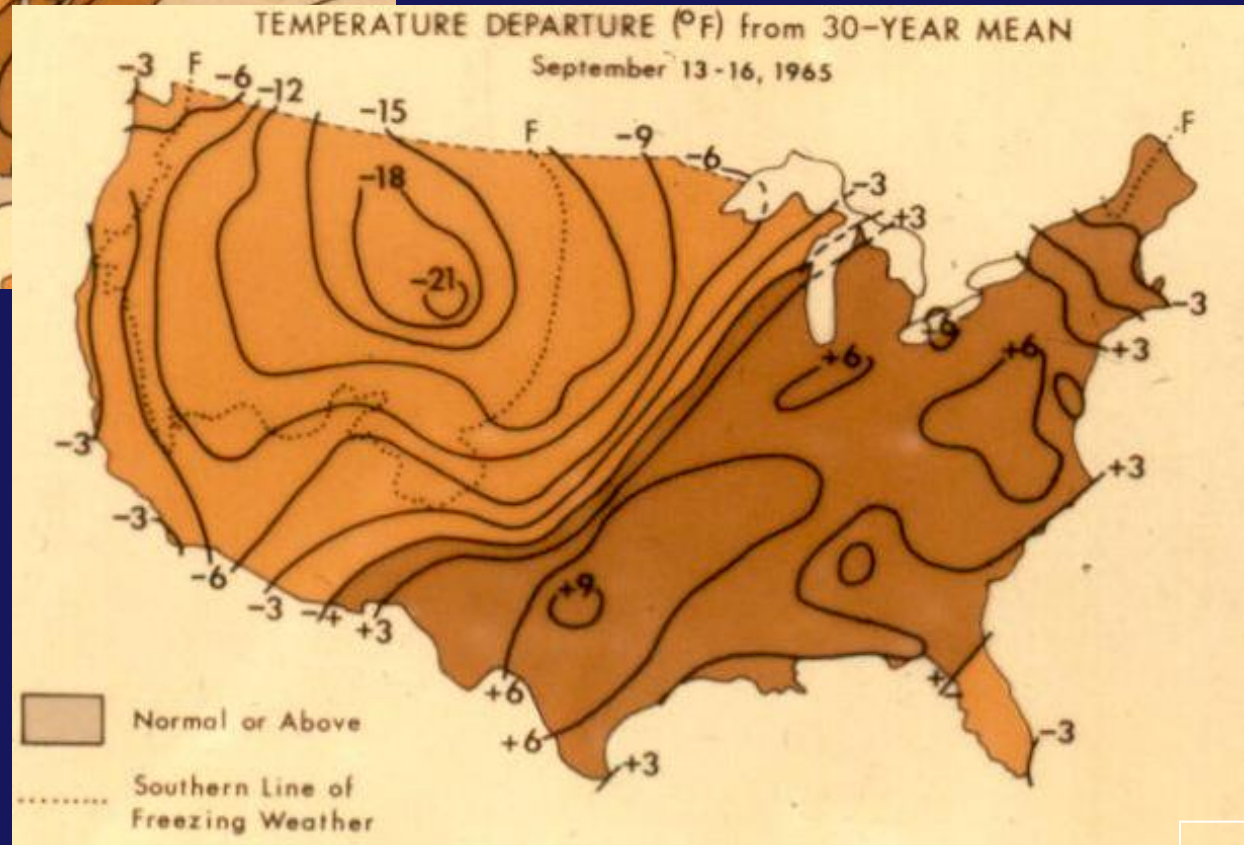
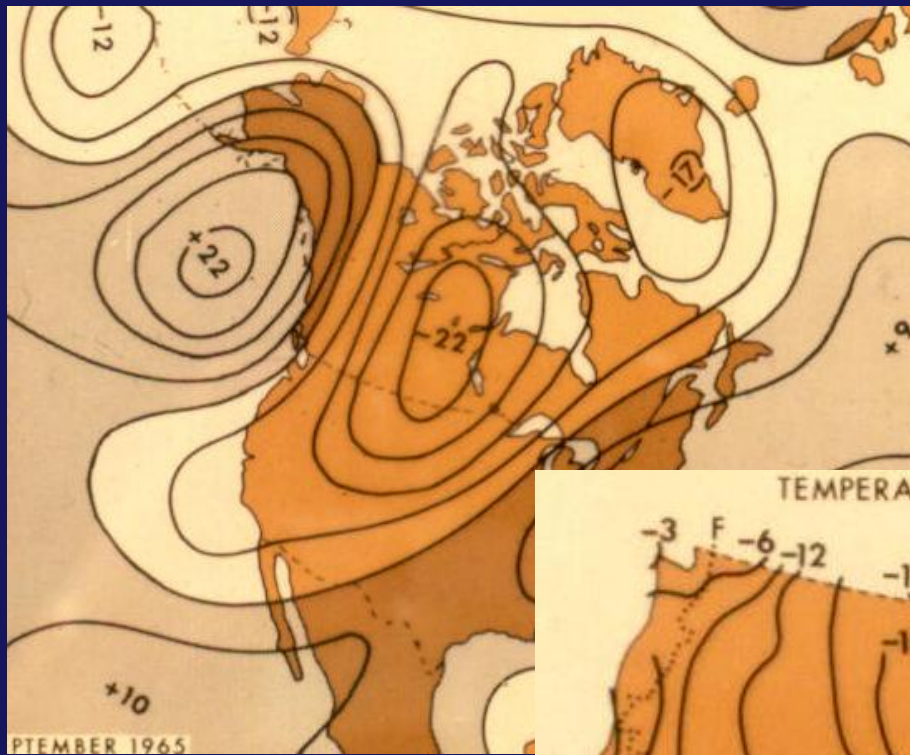


Frost ring

No frost ring

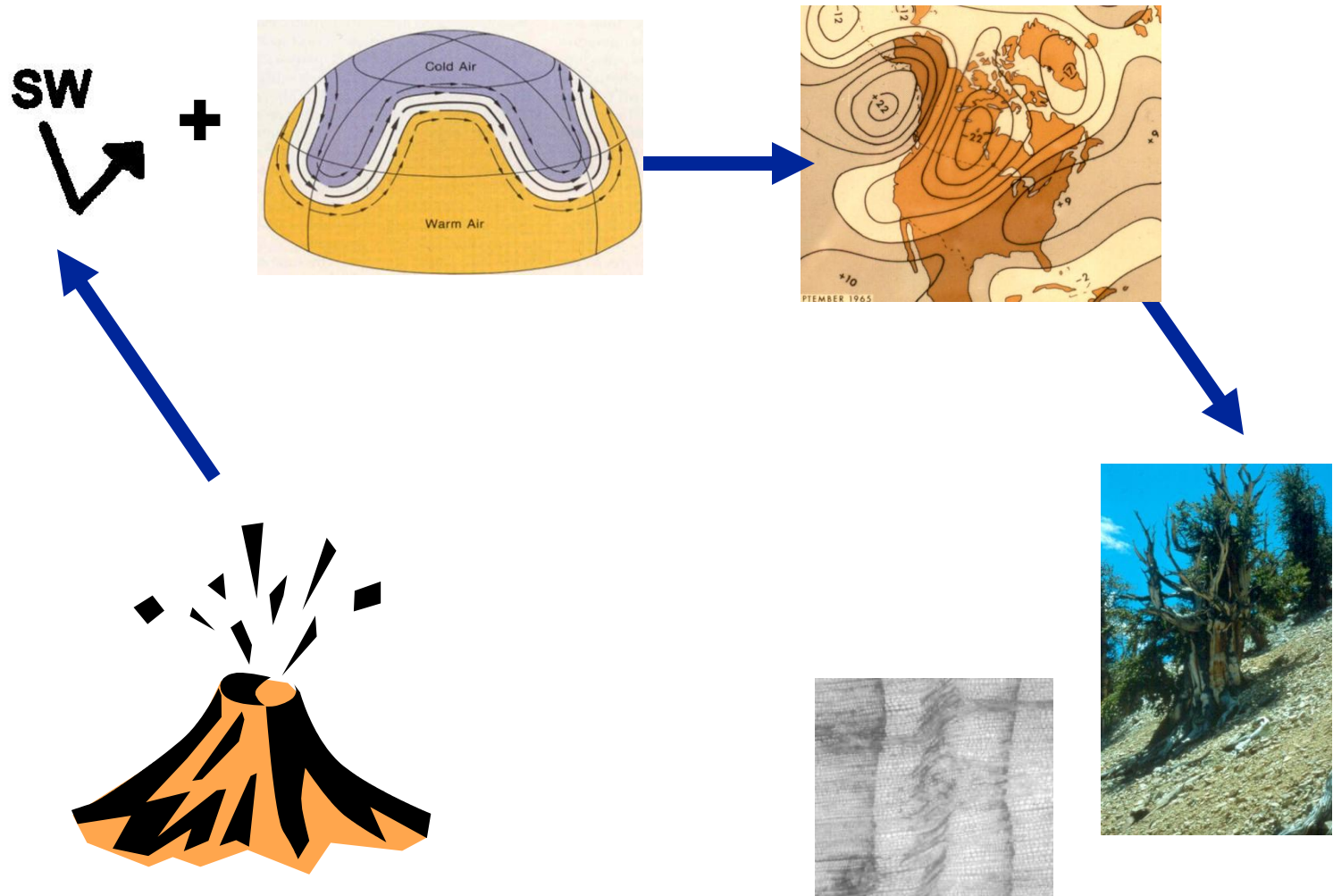


# Agung eruption 1963 → frost ring response in 1965





# Energy Balance Effects & Global Atmospheric Circulation



# WHICH ERUPTIONS ARE THE MOST CLIMATICALLY EFFECTIVE?

- **EXPLOSIVE**
- **high SULFUR content in magma**
- **whose eruption clouds inject into the STRATOSPHERE**
- **Low Latitude Eruptions** ←

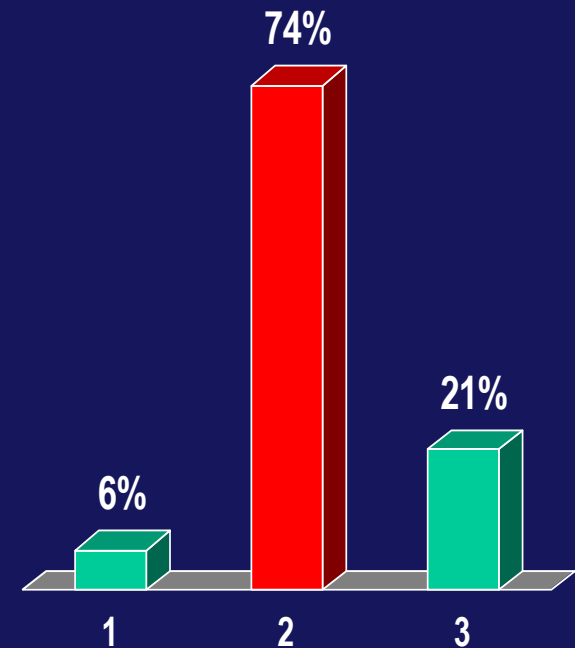
### **Q3 Why do you think Low Latitude eruptions are more climatically effective and have more of a effect?**

- 1. Because the temperature is warmer in tropical latitudes and hot air rises.**
- 2. Because the Hadley Cell circulation can distribute the volcanic aerosols into both hemispheres if the eruption occurs near the equator.**
- 3. Because the tropopause is lower over Low Latitudes and hence its easier for aerosols to get injected into the stratosphere where they will not be rained out.**

### Q3 Why do you think Low Latitude eruptions are more climatically effective and have more of an effect?

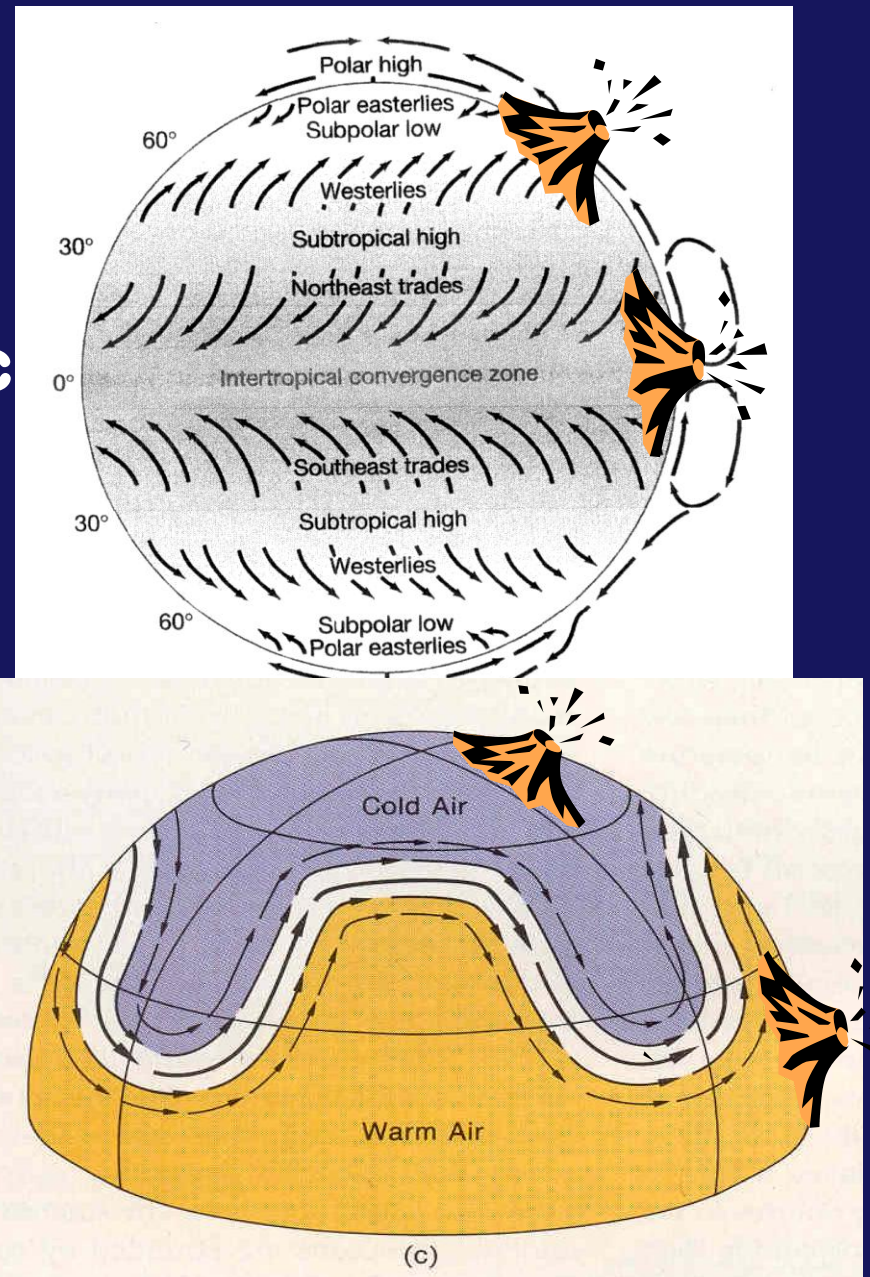
**BEGIN ANSWERING NOW!**

1. Because the temperature is warmer in tropical latitudes and hot air rises.
2. Because the Hadley Cell circulation can distribute the volcanic aerosols into both hemispheres if the eruption occurs near the equator.
3. Because the tropopause is lower over Low Latitudes and hence its easier for aerosols to get injected into the stratosphere where they will not be rained out.



- The **GEOGRAPHIC LOCATION** of the erupting volcano influences the climatic effectiveness of an eruption because of the General Circulation of the Atmosphere.

- **Low latitude eruption** clouds get circulated more broadly & in both hemispheres



# HOW DO REGIONAL CLIMATES RESPOND TO AN EXPLOSIVE ERUPTION?

In general, explosive eruptions warm the stratosphere and cool the troposphere, especially during the summer season.

## Major tropical eruption:

- Stratospheric heating is larger in the tropics → enhanced pole-to-equator temperature gradient, esp. in winter.

N.H. winter → enhanced gradient produces a stronger polar vortex → stationary wave pattern of tropospheric circulation resulting in winter warming of NH continents.

# HOW MUCH TROPOSPHERIC COOLING CAN OCCUR AND HOW LONG DOES IT LAST?

- Individual large eruptions can result in a 1-to-3 year cooling of average surface temperatures of 0.3 to 0.7° C.

Tambora in 1815

Krakatau in 1883

Agung in 1963

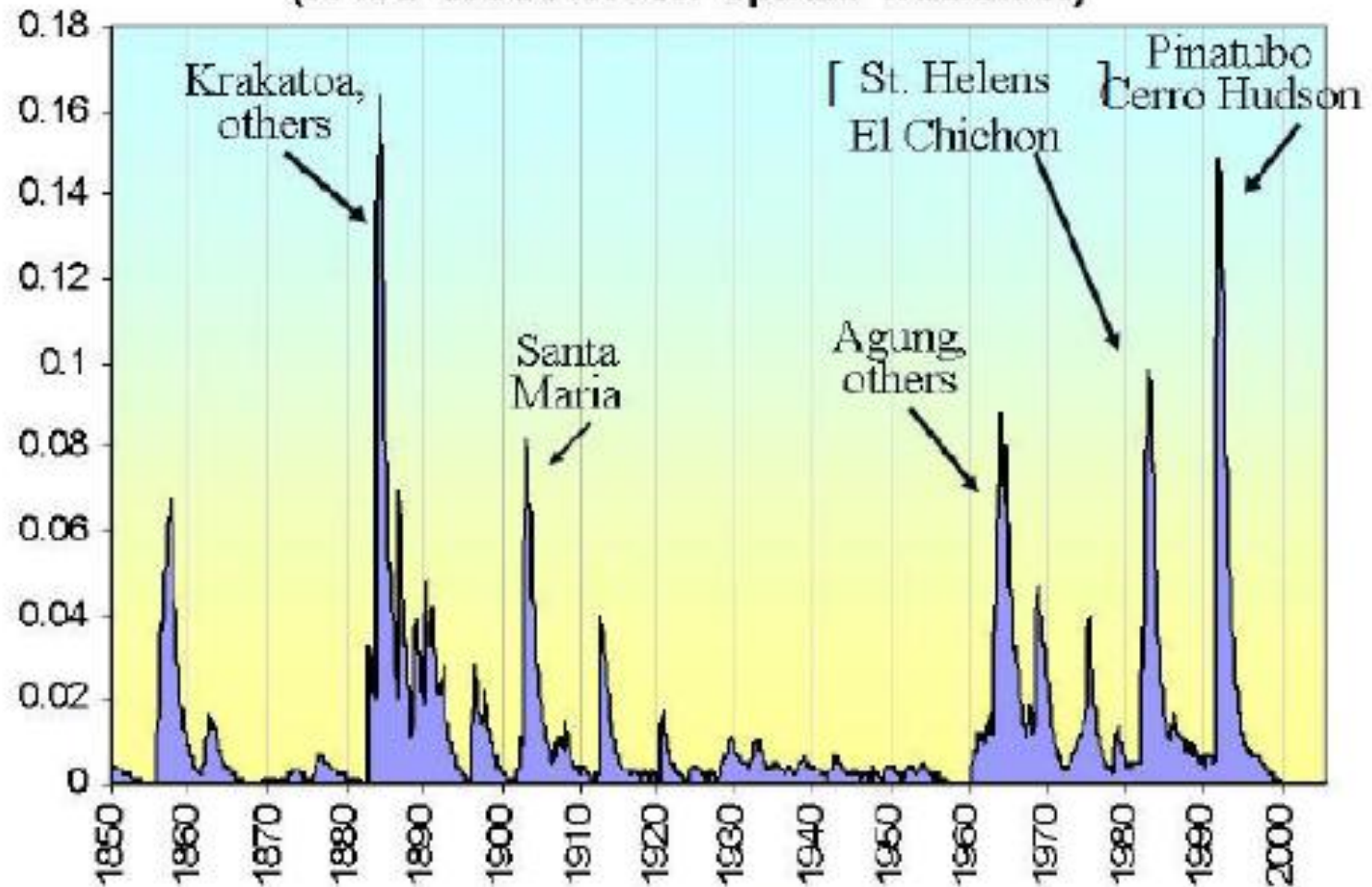
El Chichon in 1982



# HOW IMPORTANT IS EXPLOSIVE VOLCANISM AS A FORCING MECHANISM FOR PAST AND FUTURE CLIMATE CHANGES?

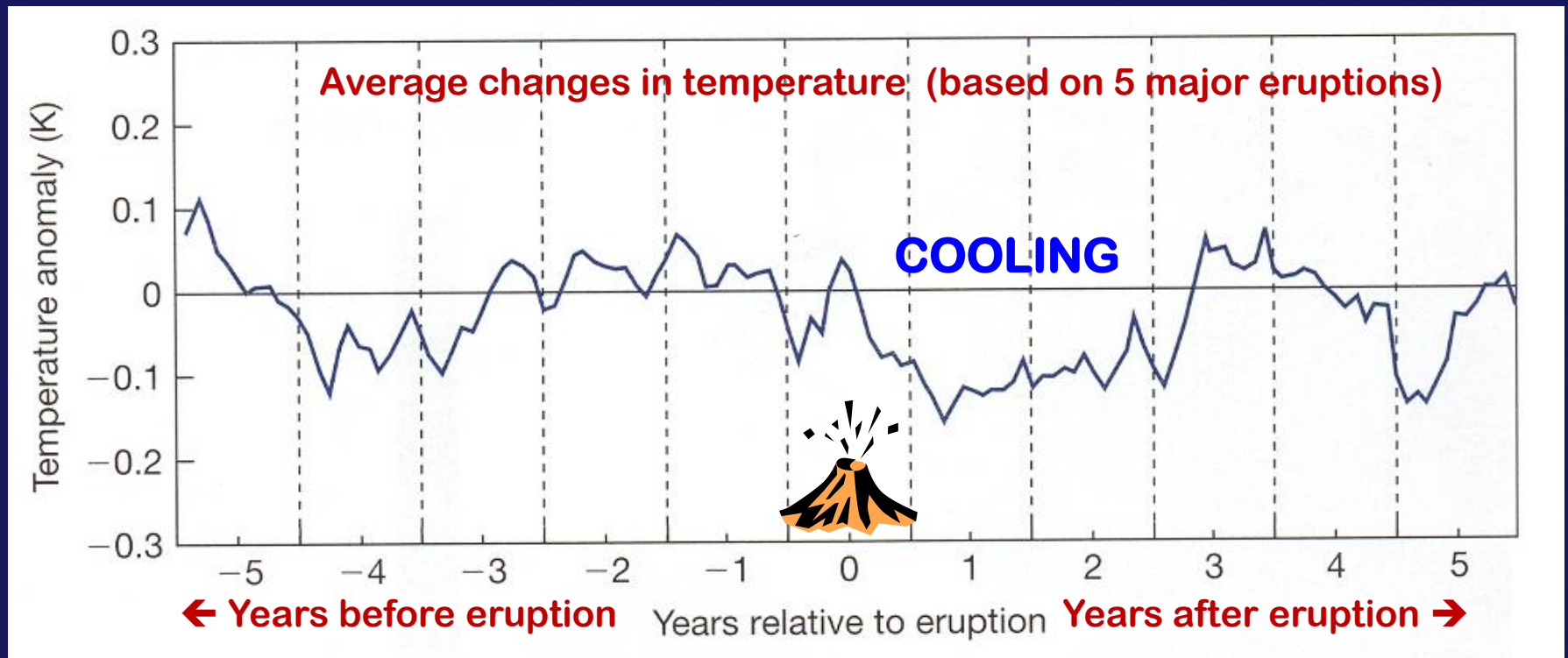
- interdecadal climate change (“Little Ice Age”)
  - Individual years, such as 1816, the “Year without a Summer” after the eruption of Tambora in 1815
  - Link not always conclusive – e.g., El Nino at same time, etc.

## Stratospheric Volcanic Aerosol (NASS GISS Aerosol Optical Thickness)



***Volcanic aerosols in the high atmosphere block solar radiation and increase cloud cover leading to widespread cooling, especially significant in summer***

# Typical Global Cooling Pattern after a Volcanic Eruption



This graph shows the global mean temperature changes for years before (-) and after a large eruption (at year zero)

# Comparison Table of Eruptions

Estimated N.H.  
temperature  
change °C

		How much magma → how big an eruption	How much aerosol got into each hemisphere		Sulfur-rich if high H <sub>2</sub> SO <sub>4</sub>		
Latitude							
Eruption	Year	Magma Erupted (km <sup>3</sup> )	Stratospheric Aerosol (Mt)		Petrologic Estimate (Mt)		N.H. ΔT
			S.H.	N.H.	H <sub>2</sub> SO <sub>4</sub>	HCl	
Tambora (8°S)	1815	50	150	150	52	220	-0.4 to -0.7
Krakatau (6°S)	1883	10	30-38	55	2.9	3.8	-0.3
Santa Maria (15°N)	1902	9	22	<20	0.6	0.4	-0.4
Katmai (58°N)	1912	15	0	<30	12.0	4.0	-0.2
Agung (8°S)	1963	0.6	30	20	2.8	1.5	-0.3
Mount St. Helens (46°N)	1980	0.35	0	—	0.08	0.04	0 to -0.1
El Chichón (17°N)	1982	0.3-0.35	<8	12	0.07*	—	-0.2
Pinatubo (15°N)	1991	5(±1)	—	20-30	~0.3*	—	-0.5

IMPORTANT: a dash — means that  
**NO INFORMATION IS AVAILABLE,**  
NOT a value of zero!

# **THINK-PAIR –SHARE UNGRADED GROUP ACTIVITY!**

**P.S. This is one of my  
favorite questions to ask  
on the FINAL EXAM!!!!**

**#1. List 4 reasons why Tambora in 1815 resulted in the largest GLOBAL cooling:**

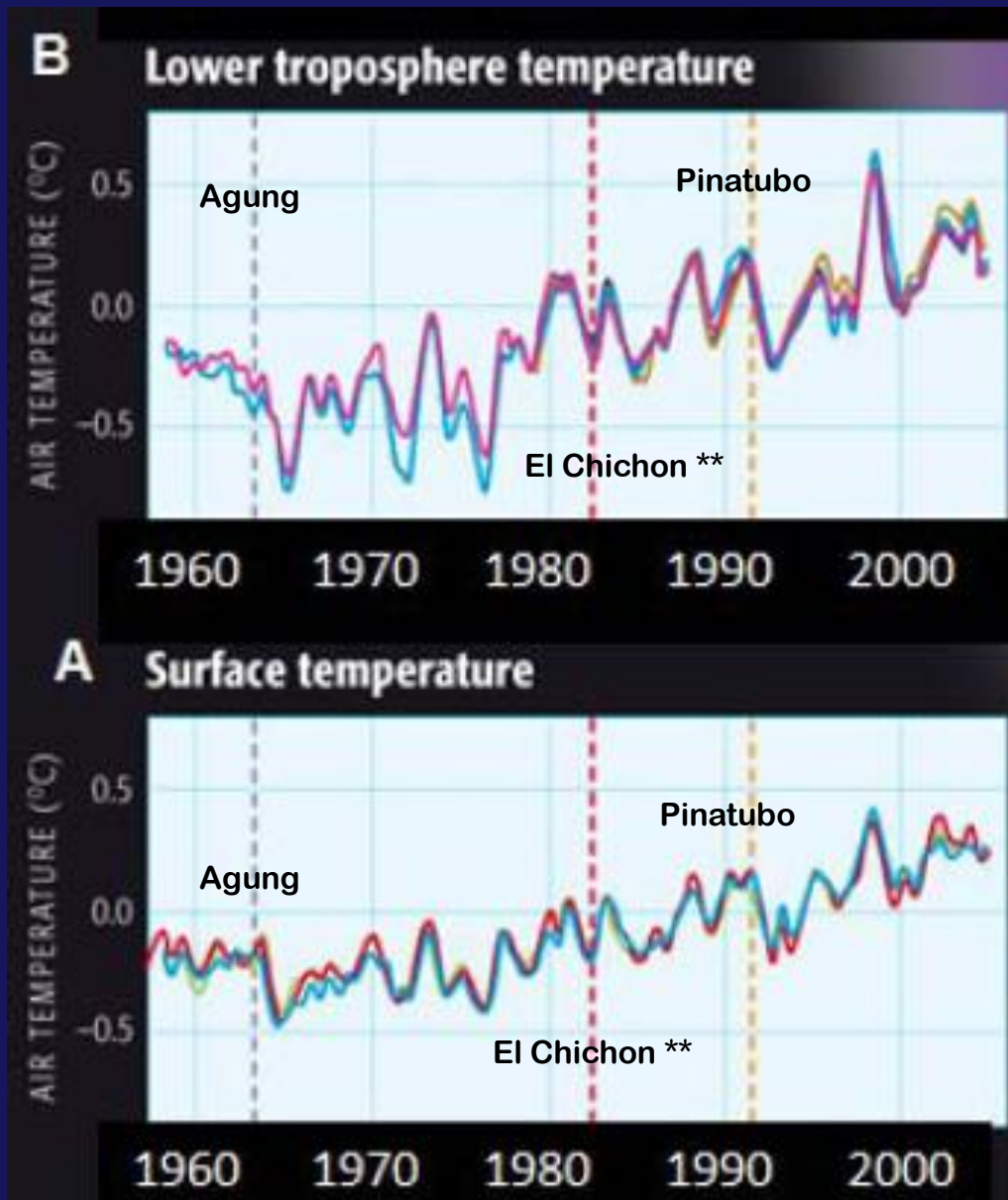
**# 2. Give at least two reasons why the eruption of Mt St. Helens was NOT a very climatically effective eruption:**



Write in the ERUPTIONS at top of page

Agung (1963)

Pinatubo (1991)



**#3. Which levels show a COOLING and which show a WARMING immediately after the eruption?**

**\*\* NOTE:** At the time of the El Chichon eruption, there was warming taking place due to a strong El Nino, hence the temperature change after this eruption shows a different response.

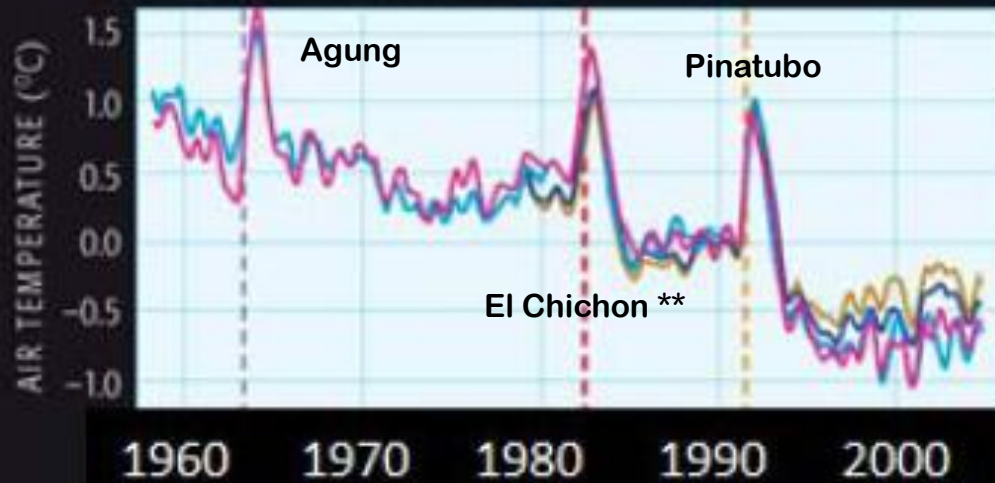
**When ANSWERING # 3 & #4 – focus on Agung & Pinatubo only**

Write in the **ERUPTIONS** at top of page

Agung (1963)

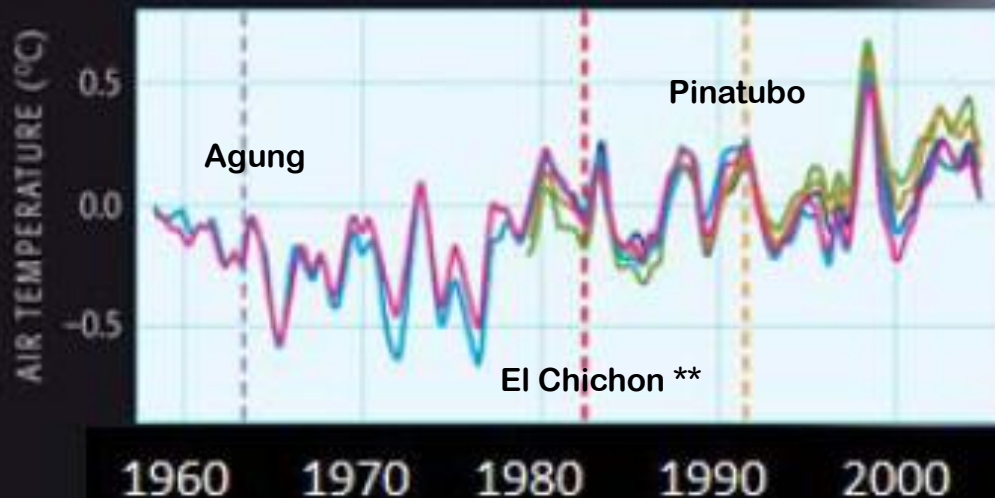
Pinatubo (1991)

**D Lower stratosphere temperature**



**C**

**Mid- to upper troposphere temperature**



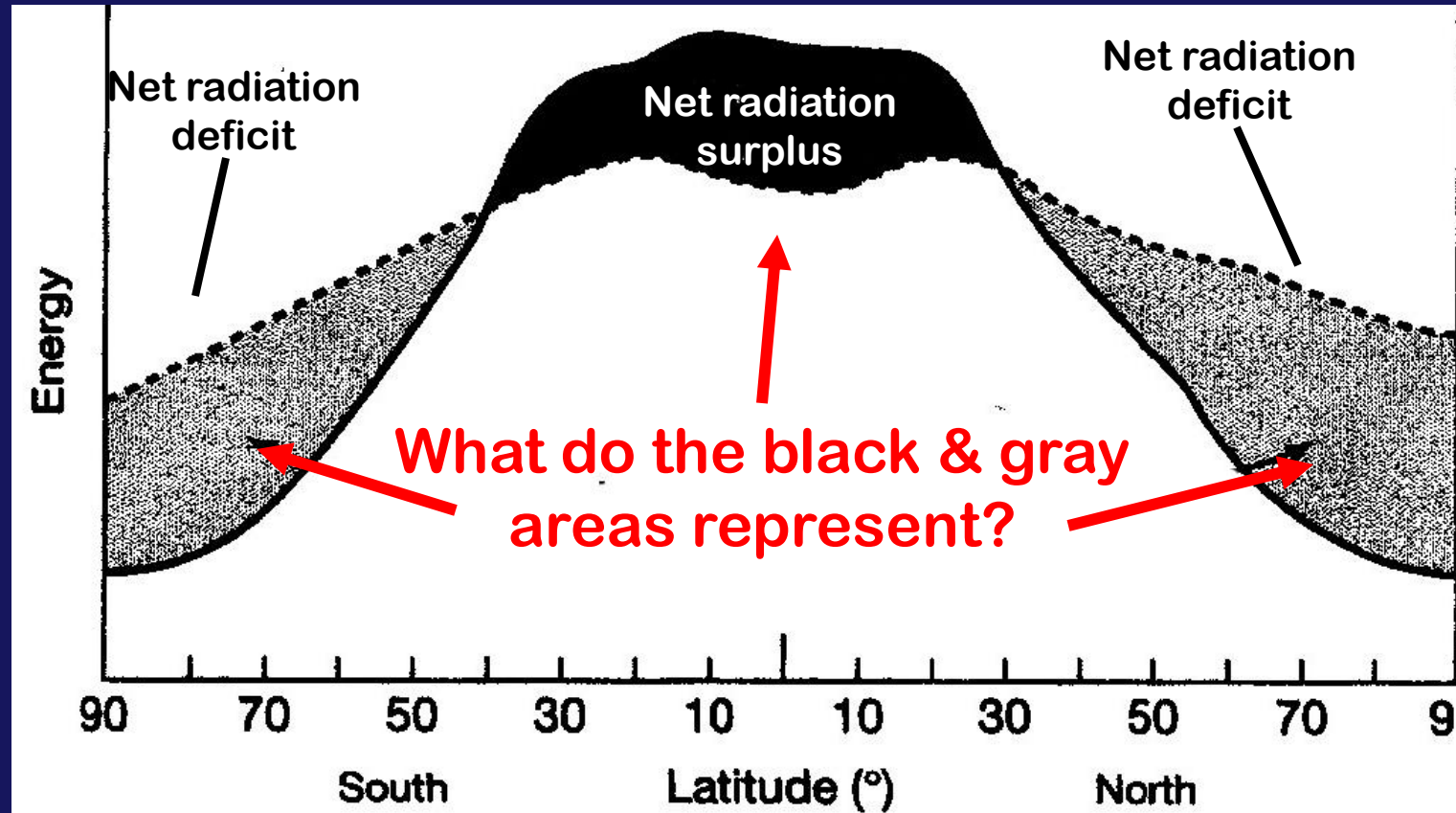
**#4. Explain WHY each level's TEMPERATURE responded as it did to the Agung & Pinatubo eruptions?**

(by referring to the Radiation Balance)

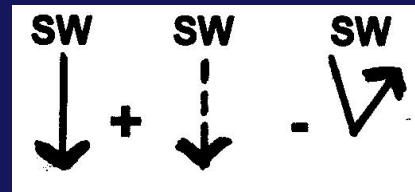
**When ANSWERING # 3 & #4 – focus on Agung & Pinatubo only**



# REMEMBER THIS IMPORTANT GRAPH?



———— Absorbed solar energy

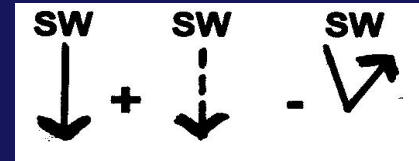


----- Emitted infrared energy  
(at top of atmosphere)



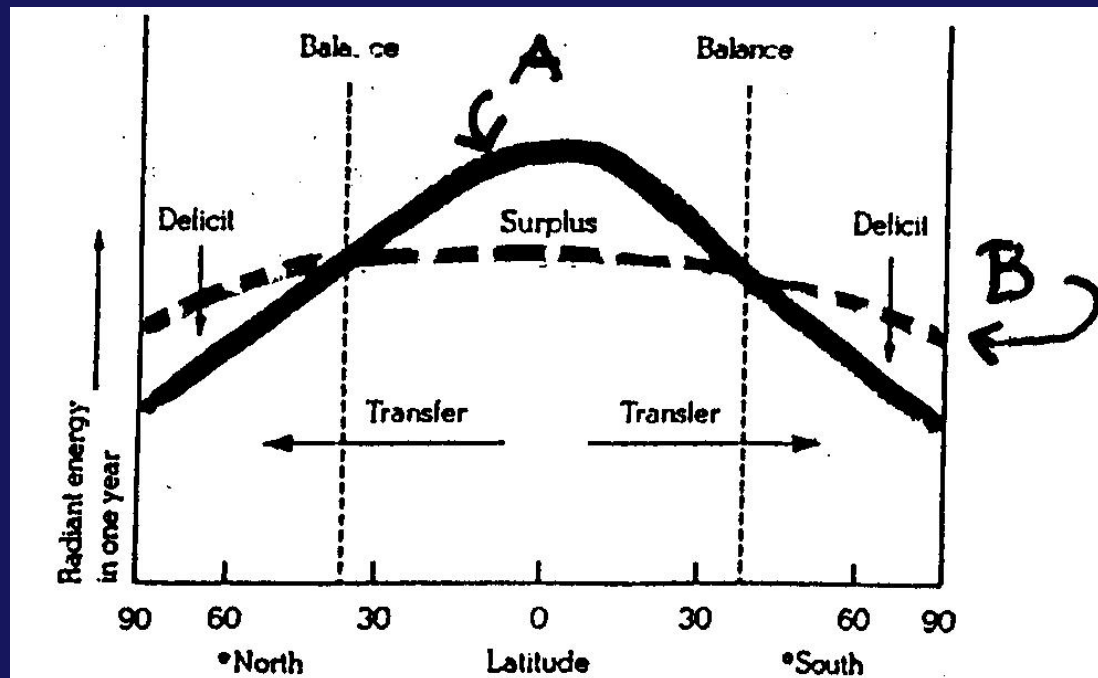
**A = incoming SW (solar) radiation**

———— Absorbed solar energy

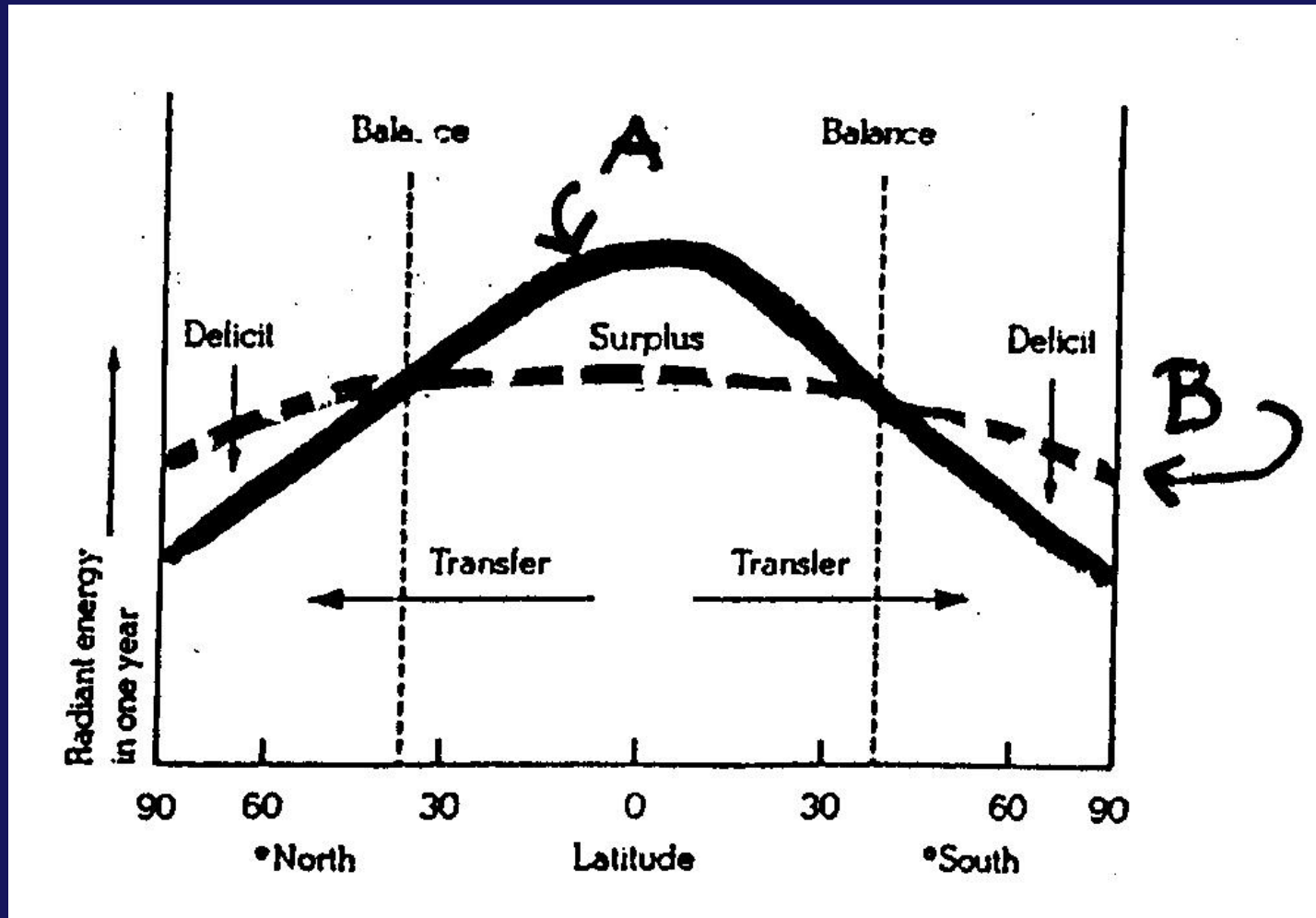


**B = outgoing LW (terrestrial) radiation**

----- Emitted infrared energy  
(at top of atmosphere)



**#5. SKETCH A NEW CURVE A OR NEW CURVE B to show how the energy balance would change if a major volcanic eruption occurred.**



*Assume:*

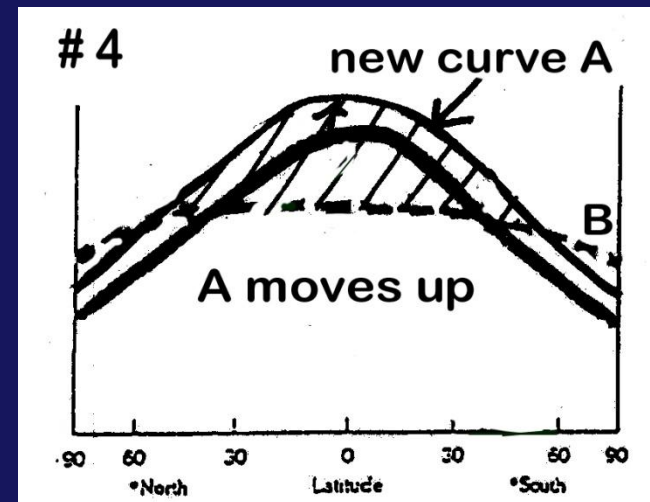
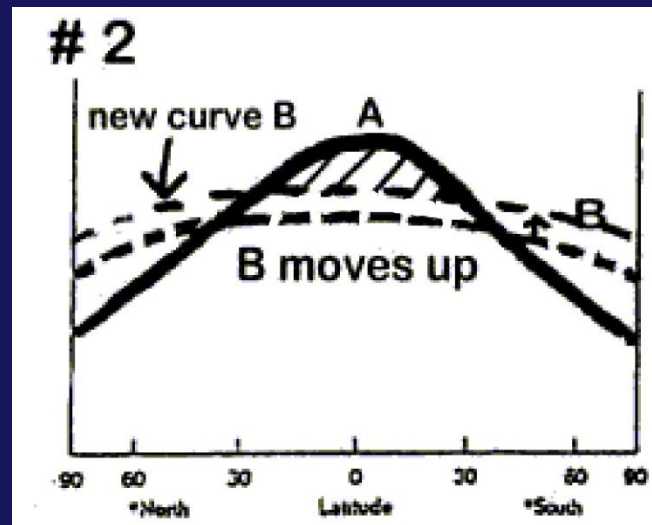
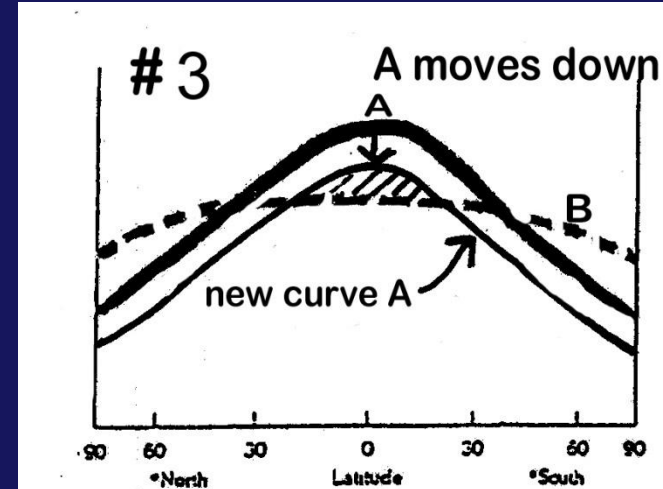
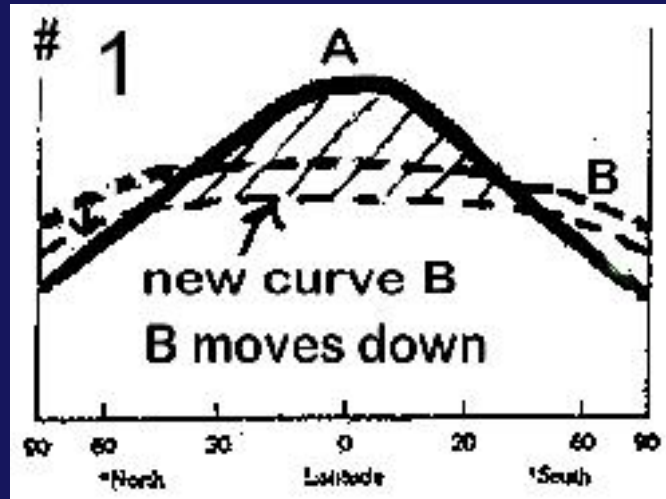
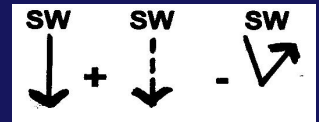
- that the eruption produces a long-lived aerosol veil in the stratosphere over both hemispheres
- that this veil reflects large amounts of incoming solar radiation back to space *before* it enters the troposphere's earth-atmosphere system shown in the graph.
- *Hint: you do not need to worry about stratospheric warming for this question.*

Four  
scenario's  
are  
possible  
for how  
you should  
sketch the  
new graph:

IF CURVE B  
is affected:



If CURVE A  
is affected:



# THE ANSWERS!





# #1. List 4 reasons why Tambora in 1815 resulted in the largest GLOBAL cooling:

- #1 Low latitude eruption → both hemispheres
- #2 Large amount of eruptive material (50 sq km!)
- #3 Aerosol cloud was HUGE and went into both hemispheres equally
- #4 Sulfuric acid ( $\text{H}_2\text{SO}_4$ ) content was very large

**#2. Give at least two reasons why the eruption of Mt St. Helens was NOT a very climatically effective eruption:**


**#1 High latitude – could only affect part of Northern Hemisphere**

**#2 Low sulfur content  
(also, low volume, didn't get to S. Hemisphere, etc.)**



# 3 HOW did the temperature at the 4 levels respond to the Agung and Pinatubo eruptions?

#4 EXPLAIN WHY – referring to Radiation Balance?



Level A (Surface) – Cooled

Why?  by sulfate aerosols in stratosphere and therefore less SW got into troposphere to be absorbed by Earth's surface

Level B (Lower Troposphere) – Cooled

Why?  by stratospheric aerosols => less SW absorbed at surface and in troposphere,  
ALSO: less  radiated up into troposphere from the cooler Earth's surface

## Level D (Lower Stratosphere) – Warmed immediately after both eruptions

**Why?** Sulfate aerosols in the stratosphere absorbed some wavelengths of incoming SW  and heated up, they also absorbed some of the Earth's outgoing LW  as it radiated up out of the troposphere

Level C (Mid-Upper Troposphere) –  
After Agung: warmed briefly, then cooled;  
After Pinatubo: cooled

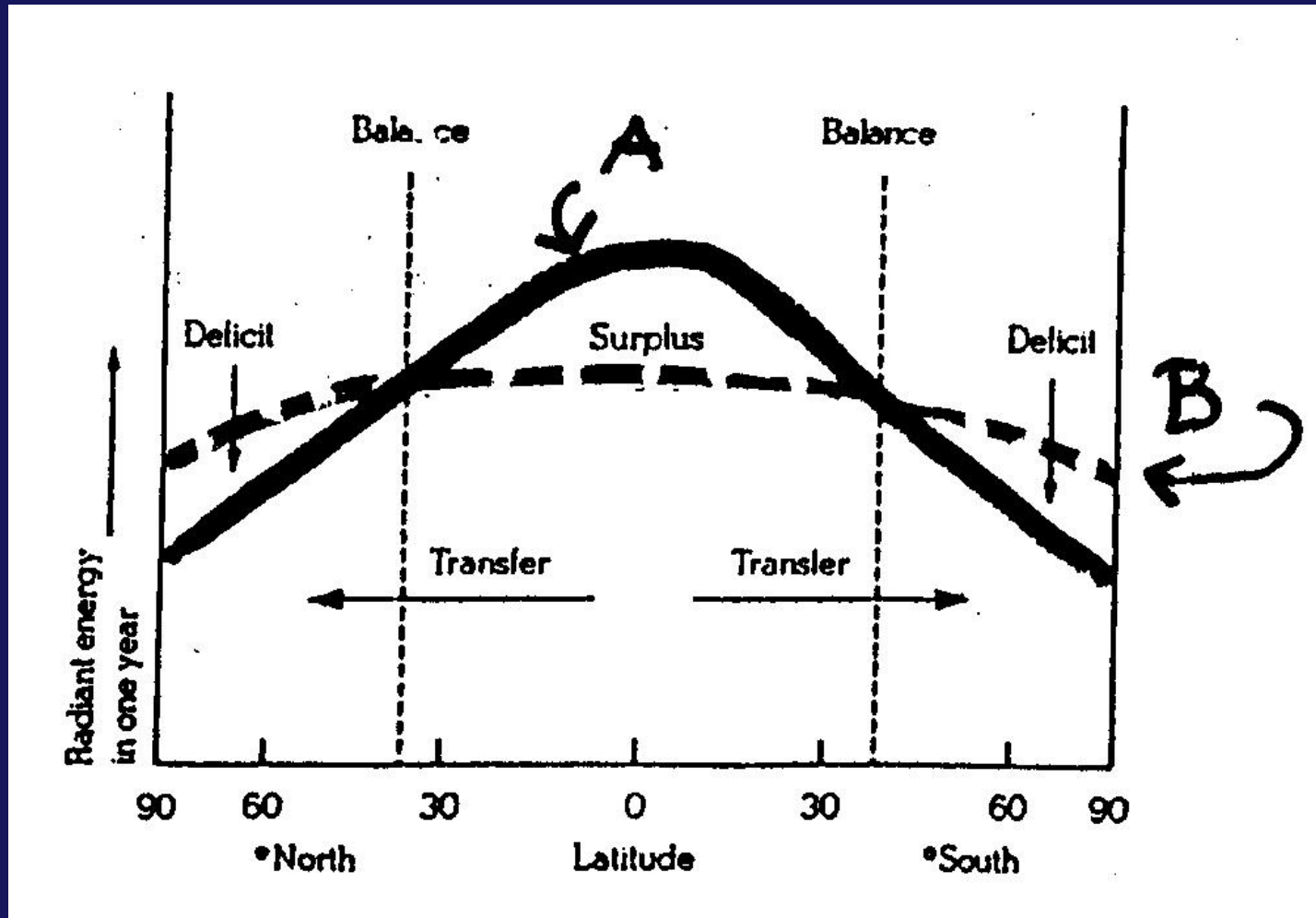
**Why?** A transitional layer with some processes like those in Layer B and Layer D: sulfate aerosols reflected some  and absorbed some  plus outgoing LW 

## TO SUMMARIZE: 2 KEY POINTS

- Major eruptions with a long-lived sulfate aerosol veil REFLECT incoming solar radiation back to space BEFORE it enters the mid- & lower troposphere or gets to the Earth's surface, hence the **troposphere & surface get COOLER** after an eruption.
- The aerosols in the stratosphere can also ABSORB some wavelengths of incoming SW and outgoing LW, so that the **stratosphere WARMS** slightly after an eruption.



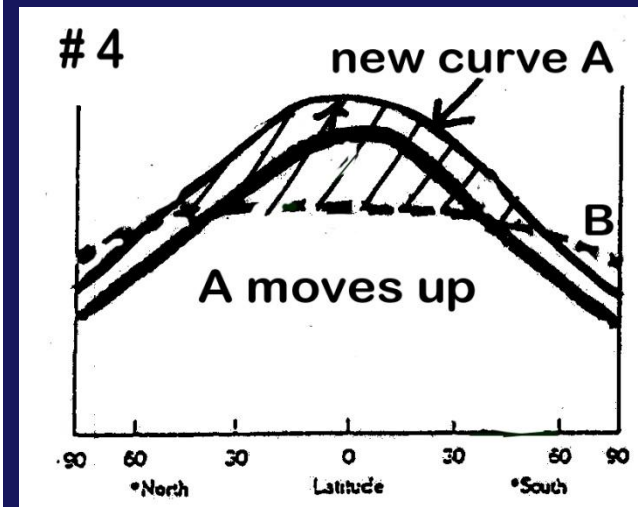
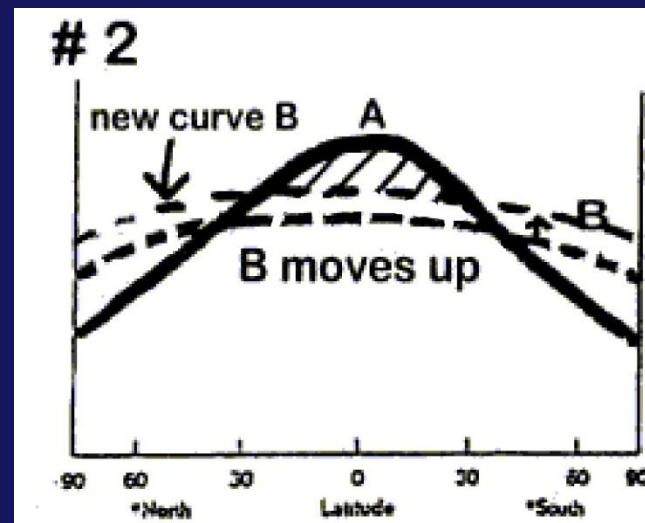
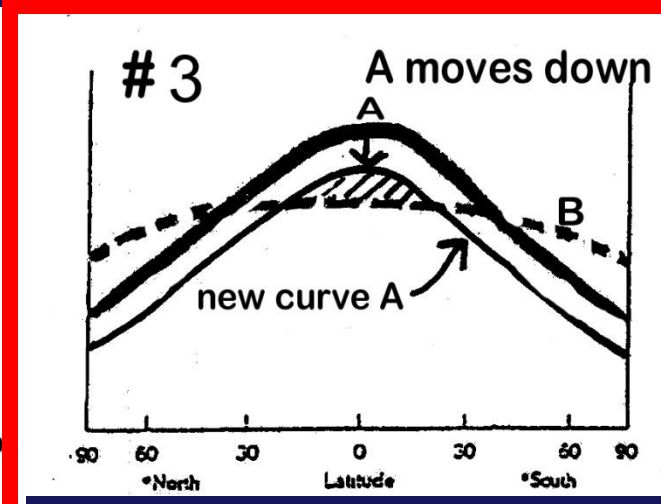
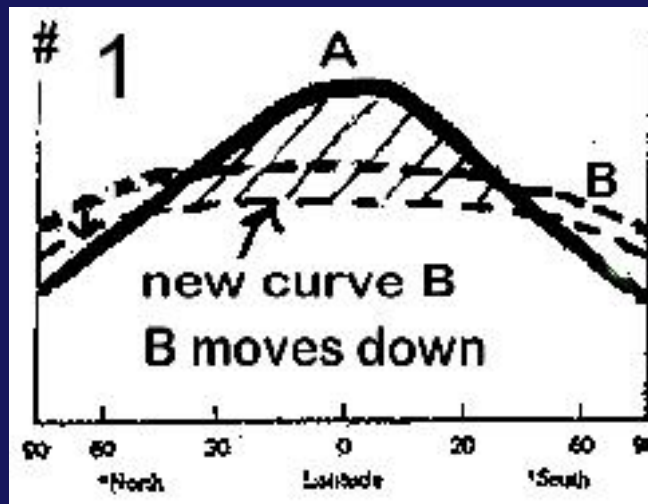
**SKETCH A NEW CURVE A OR NEW CURVE B to show how the energy balance would change if a major volcanic eruption occurred .**



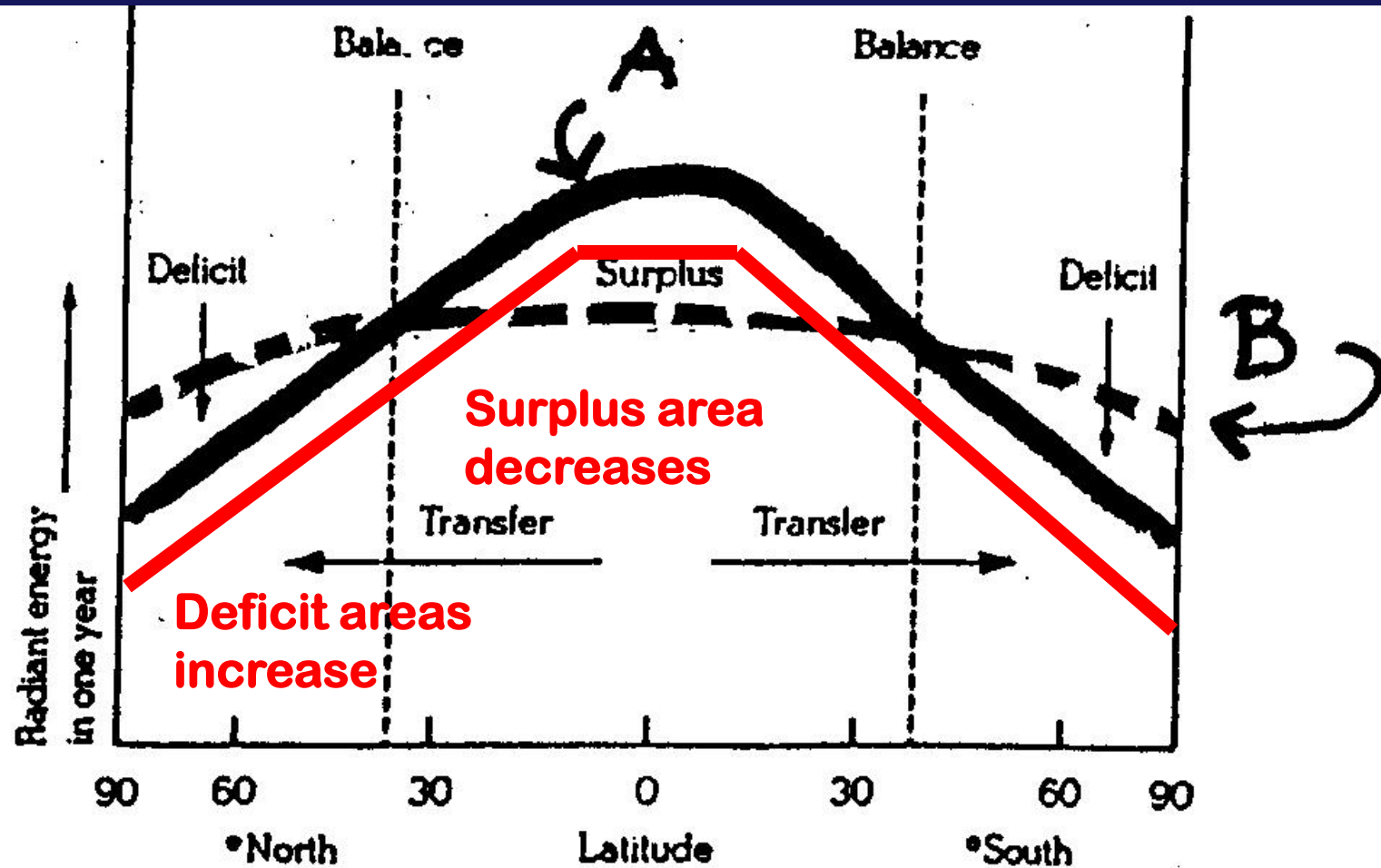
**WHICH ONE IS  
RIGHT?**

**Does the  
change affect  
CURVE A or  
CURVE B?**

**Show how the  
energy  
balance would  
change if a  
major volcanic  
eruption  
occurred .**



A moves down, & B stays ~ same; eventually  
B will also move down a bit due to cooler  
Earth temps and less outgoing LW



# The End . . . but



## Why a red sunset after an eruption???