

Topic # 12

Natural Climate Processes

A “Primer” on
How the Energy Balance Drives
Atmospheric & Oceanic Circulation,
Natural Climatic Processes

pp 63-68 in Class Notes



RADIATION / ENERGY BALANCE

$$R_{\text{NET}} = \begin{array}{c} \text{SW} \\ \downarrow \end{array} + \begin{array}{c} \text{SW} \\ \vdots \downarrow \end{array} - \begin{array}{c} \text{SW} \\ \nearrow \end{array} - \begin{array}{c} \uparrow \\ \text{LW} \end{array} + \begin{array}{c} \text{LW} \\ \downarrow \end{array} = H + LE + G$$

“Radiation Balance” part

$$R_{\text{NET}} = \begin{array}{c} \text{SW} \\ \downarrow \end{array} + \begin{array}{c} \text{SW} \\ \vdots \downarrow \end{array} - \begin{array}{c} \text{SW} \\ \nearrow \end{array} - \begin{array}{c} \uparrow \\ \text{LW} \end{array} + \begin{array}{c} \text{LW} \\ \downarrow \end{array}$$

All components are referring to electromagnetic radiation

All components are referring to modes of heat energy transfer or heat energy storage involving matter

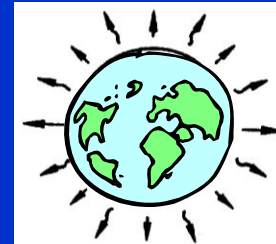
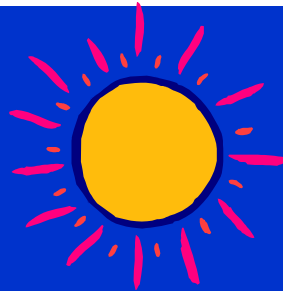
“Energy Balance” part

$$R_{\text{NET}} = H + LE + G$$

Start out here,
with energy
from the SUN
radiated to
Earth and so
forth . . .

“Radiation Balance” part

$$\begin{array}{c} \text{SW} \\ \downarrow \end{array} + \begin{array}{c} \text{SW} \\ \downarrow \end{array} - \begin{array}{c} \text{SW} \\ \nearrow \end{array} - \begin{array}{c} \uparrow \\ \text{LW} \end{array} + \begin{array}{c} \text{LW} \\ \downarrow \end{array} = \boxed{R_{\text{NET}}}$$



The R_{NET} is then able to
be used in thermal
energy “heat transfer”
processes which
manifest themselves
as weather & climate!

“Energy Balance” part

$$R_{\text{NET}} = H + LE + G$$

Thermal Energy Review

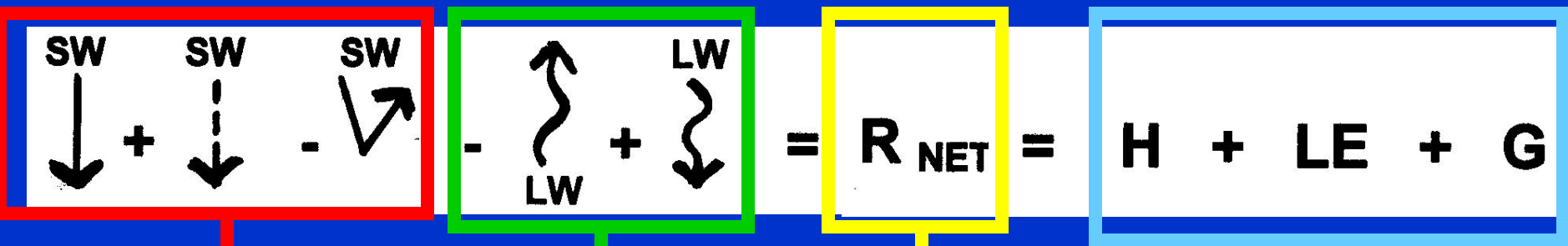
Heat (def) = the thermal energy that is **transferred** from one body to another because of a temperature difference.

- **Sensible Heat transfer (H)**
- **Latent Heat transfer (LE)**

plus (after transfer) thermal energy can be **STORED (G)**

$$\mathbf{H + LE + G}$$

ENERGY IN THE EARTH-ATMOSPHERE SYSTEM



Ultimate source
of energy is the
SUN (SW)

After
absorption of
SW, LW energy
is radiated in &
out by EARTH
& Atmosphere

Any
NET
(leftover)
energy

Goes into
the HEAT
TRANSFER
processes that
drive
**WEATHER &
CLIMATE !**

**The Earth
[as viewed from space]
... has the organized, self-
contained look of a live creature,
full of information, marvelously
skilled in handling the sun.**

~ Lewis Thomas



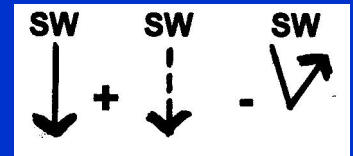
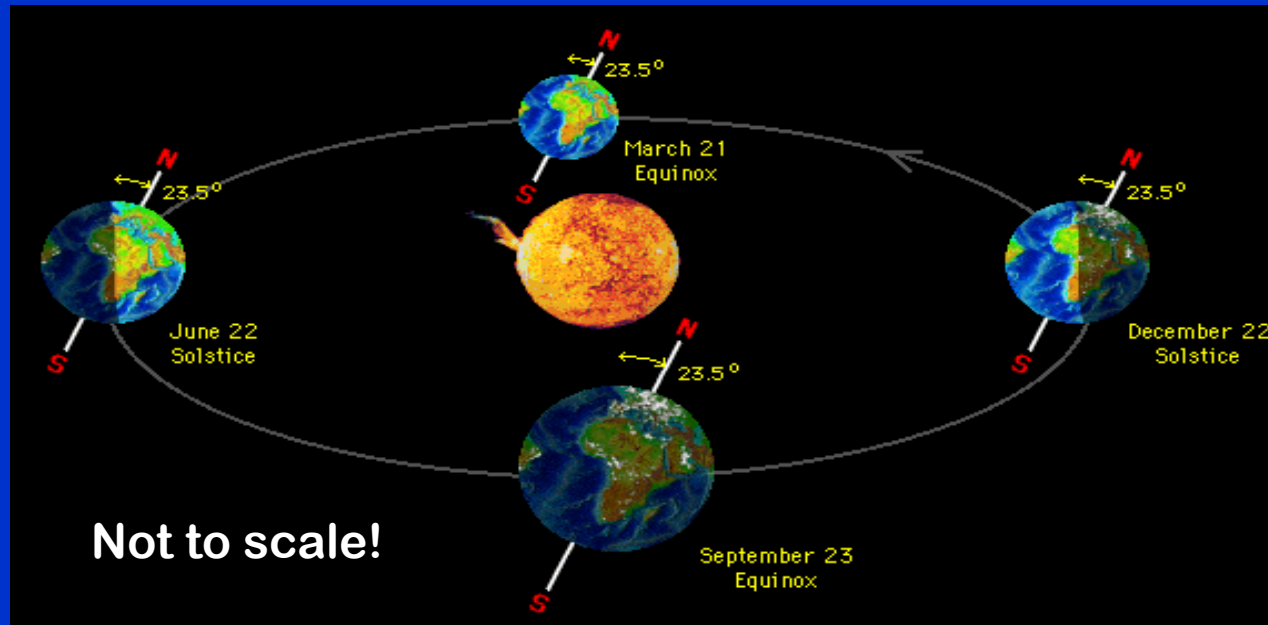
LINKING THE ENERGY BALANCE TO ATMOSPHERIC CIRCULATION . . .

**We'll start with the SUN
(SOLAR INSOLATION)**

IN – SOL- ATION =

**Amount of incoming solar energy
received by a point on Earth's surface**

To drive the circulation, the initial source of energy is from the Sun:



EARTH-SUN Relationships

4 Things to Know about Earth-Sun Relationships:

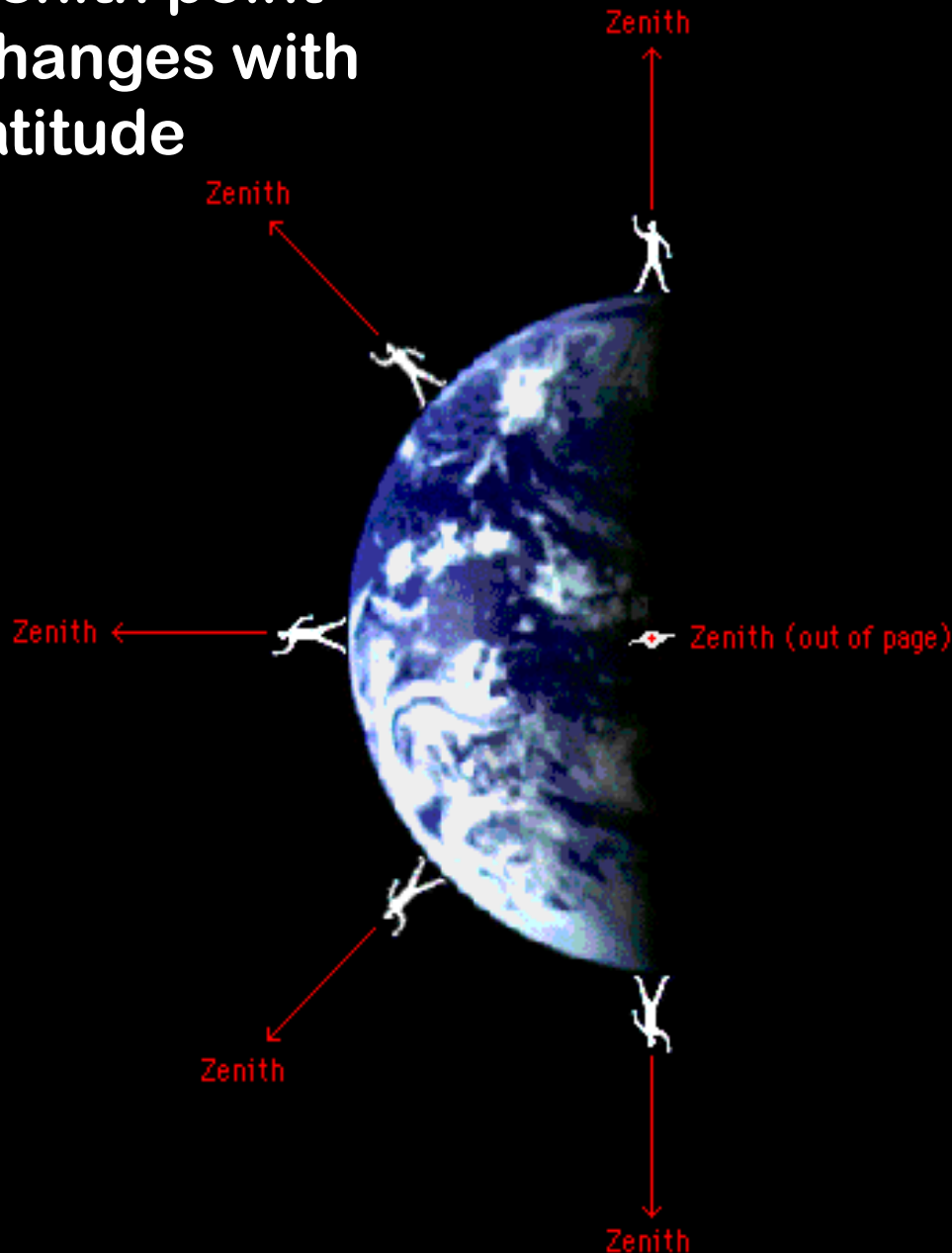
- 1) Earth orbits Sun in one year
- 2) Orbit is not a perfect circle (= an ellipse)
- 3) Earth's orbit around Sun can be "traced" on a plane (called the "**Plane of the Ecliptic**" – plane passes thru the center of Sun & Earth)
- 4) Earth's axis **tilts 23.5°** from a \perp to the "Plane of The Ecliptic"

These 4 Earth-Sun Properties lead to:
Intensity + Duration

2 factors that determine the AMOUNT
OF SOLAR INSOLATION
as seasons progress:

- (1) **INTENSITY** of sun's rays
(perpendicular to surface = more intense)
- (2) **DURATION** of daily insolation
(longer day length = more insolation)

Zenith point
changes with
latitude



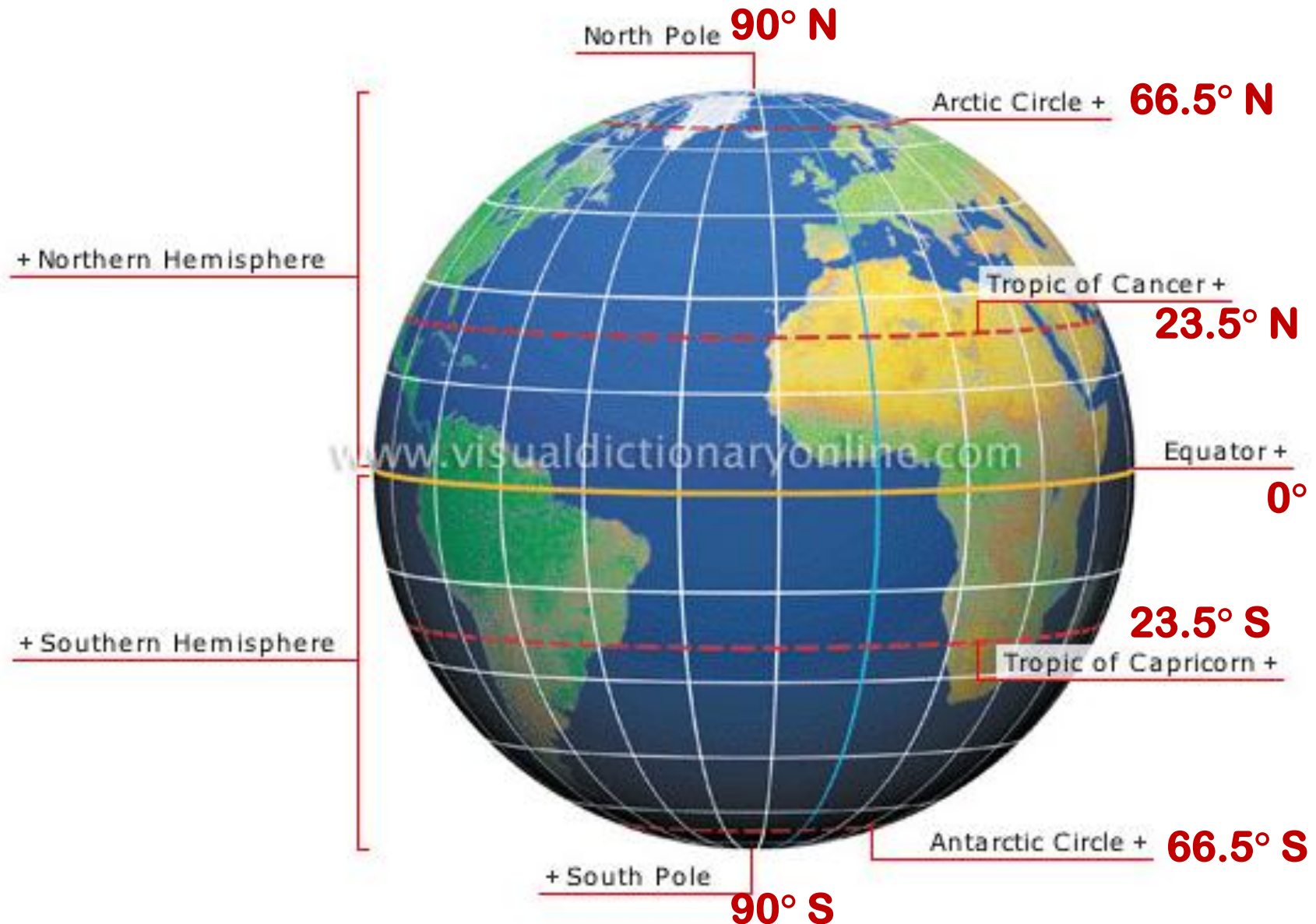
A useful term:

ZENITH =
The point
directly
overhead

INTENSITY is
greatest at any
spot on Earth
when sun is
closest to the
ZENITH!



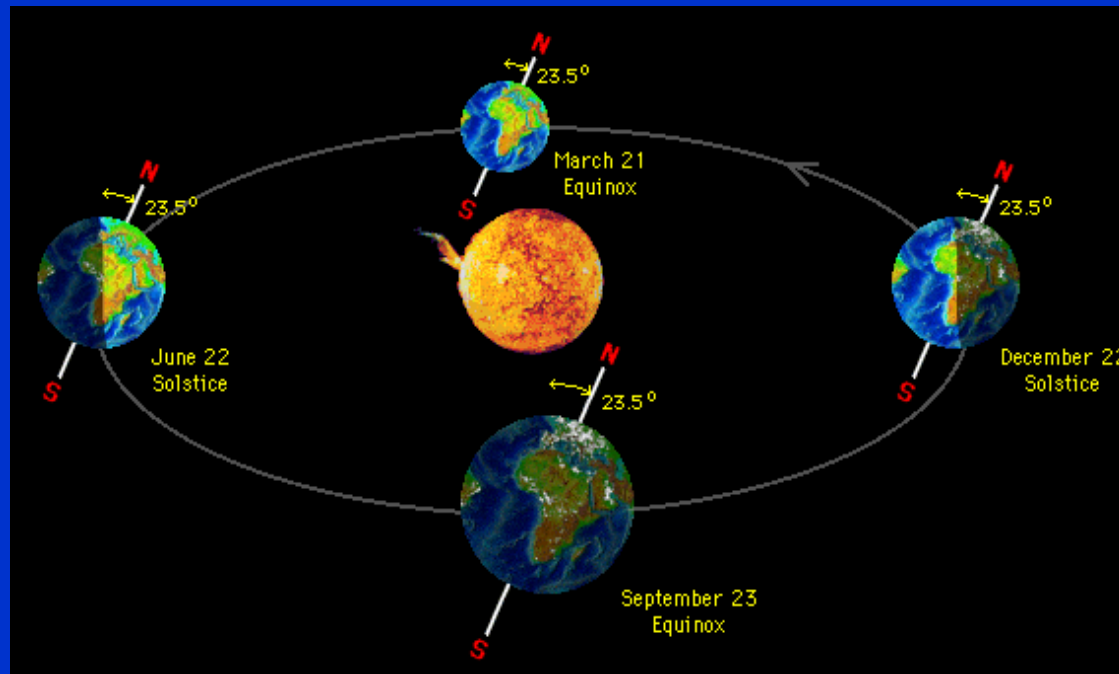
QUICKIE LATITUDE REVIEW:



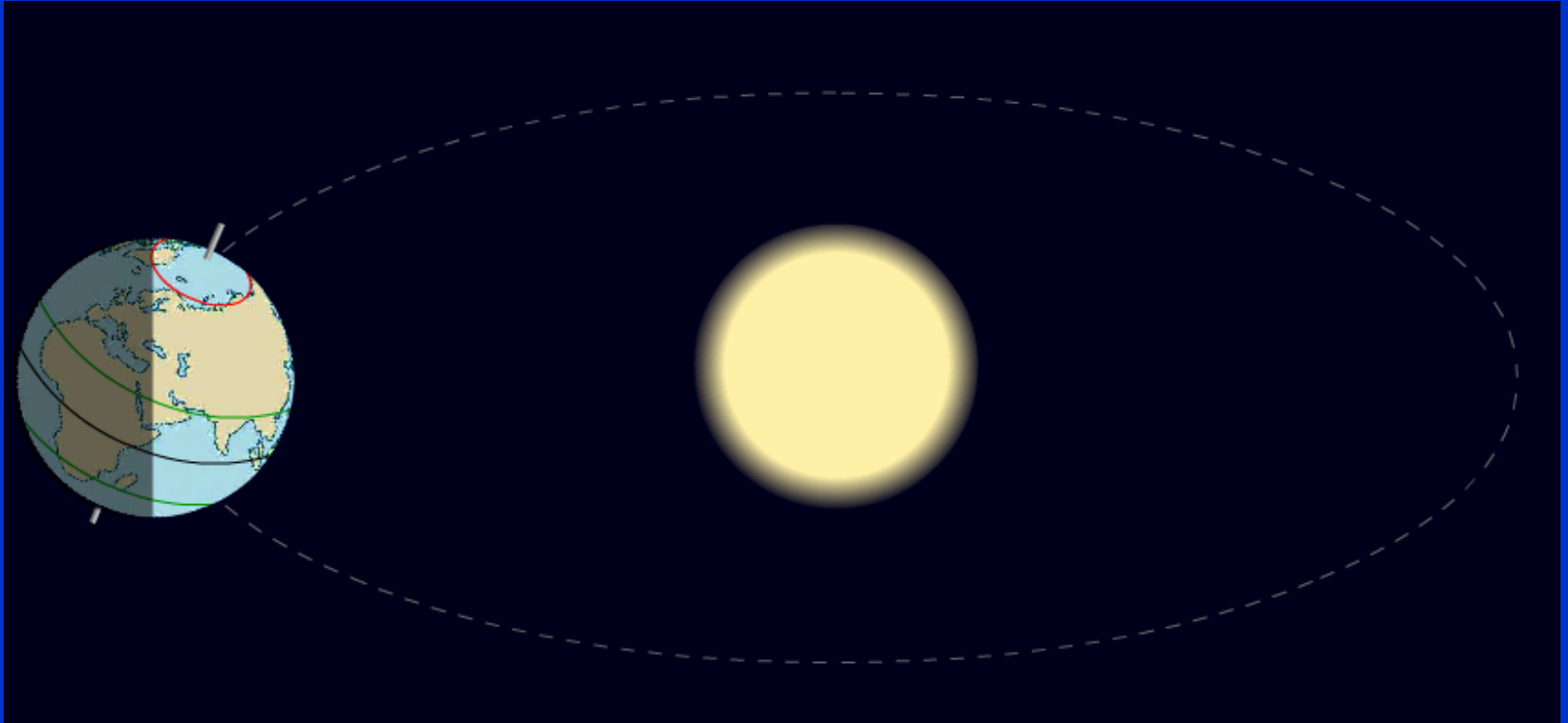
EARTH-SUN RELATIONSHIPS & The SEASONS:

VIEW THE ANIMATION:

http://mesoscale.agron.iastate.edu/agron206/animations/01_EarthSun.html

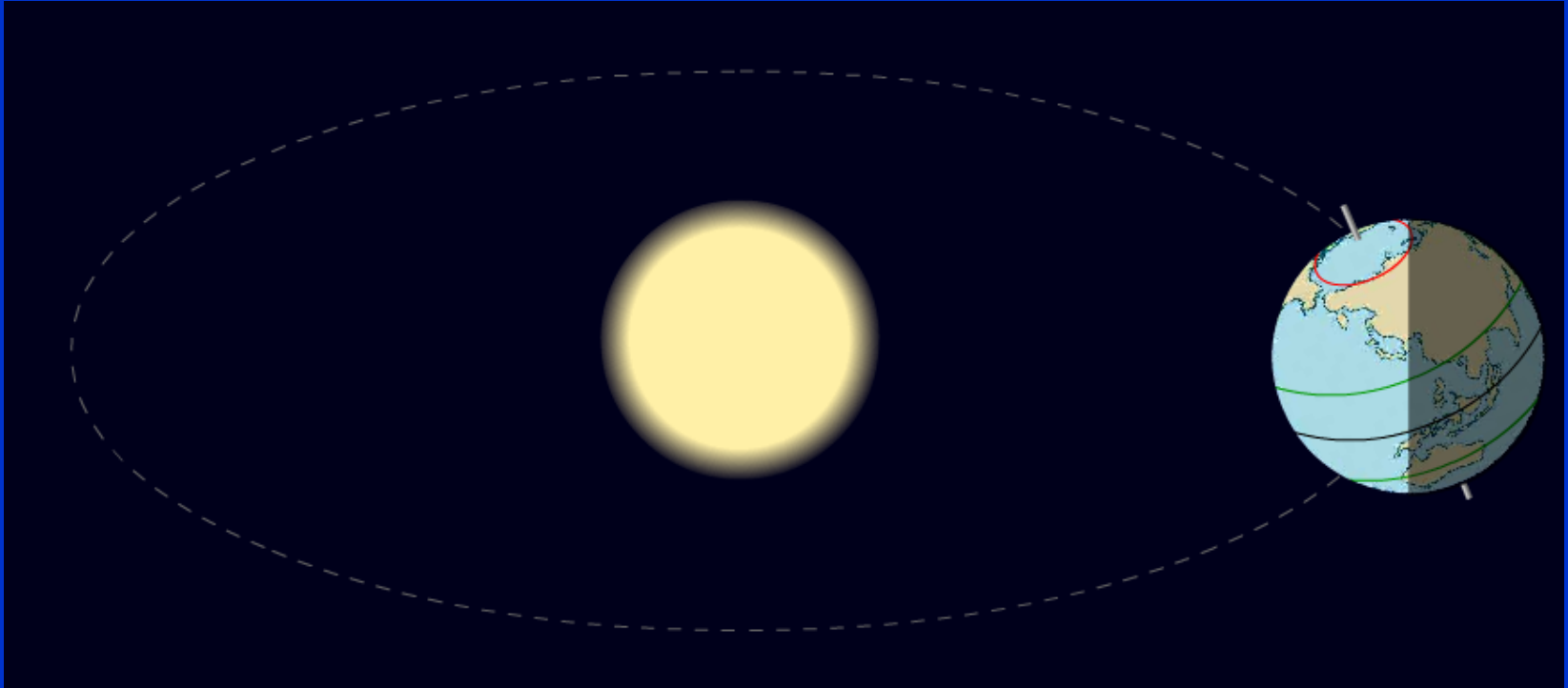


JUNE SOLSTICE



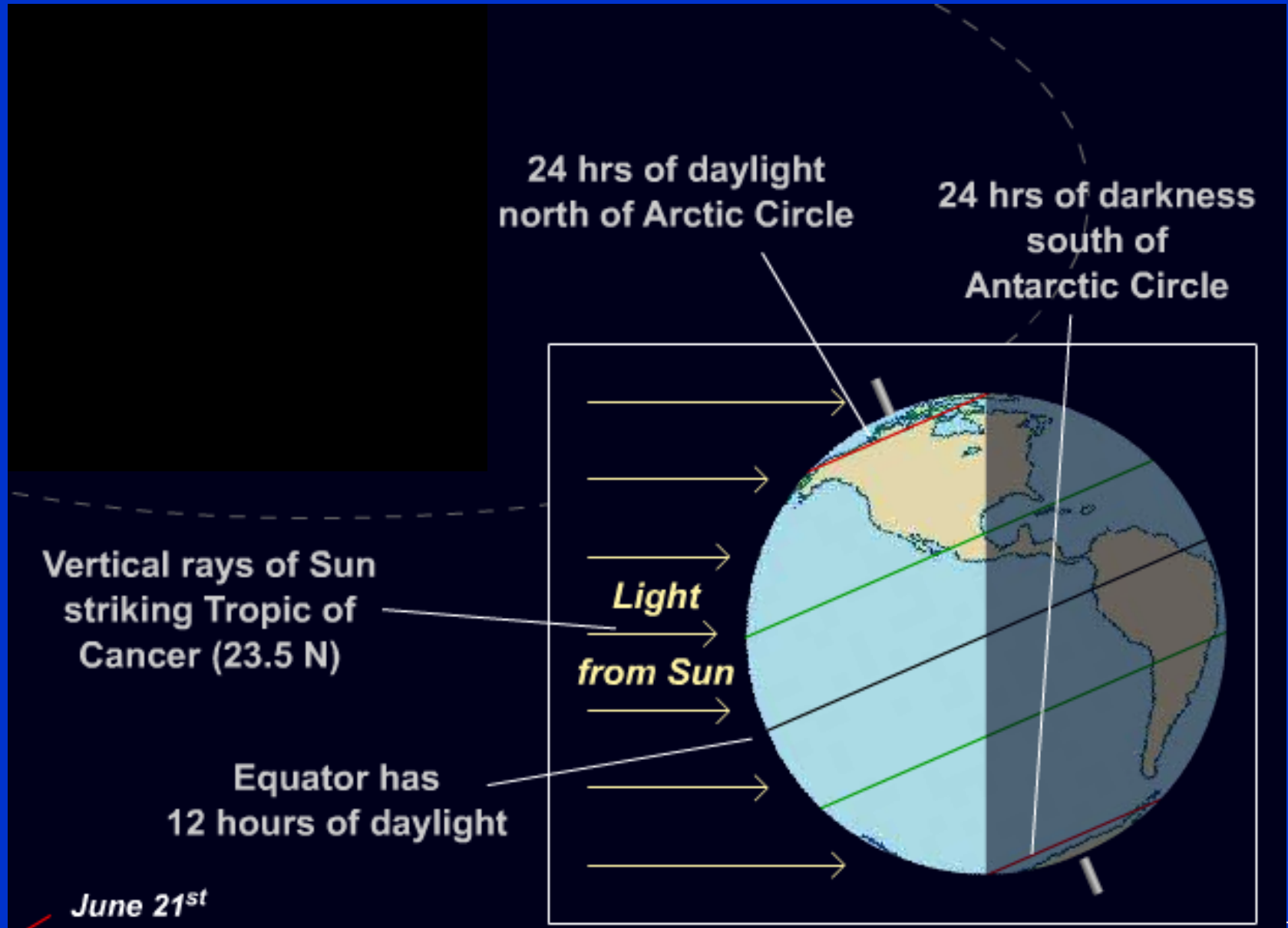
**As viewed from one
side of Sun**

JUNE SOLSTICE



As viewed from the
other side of the Sun

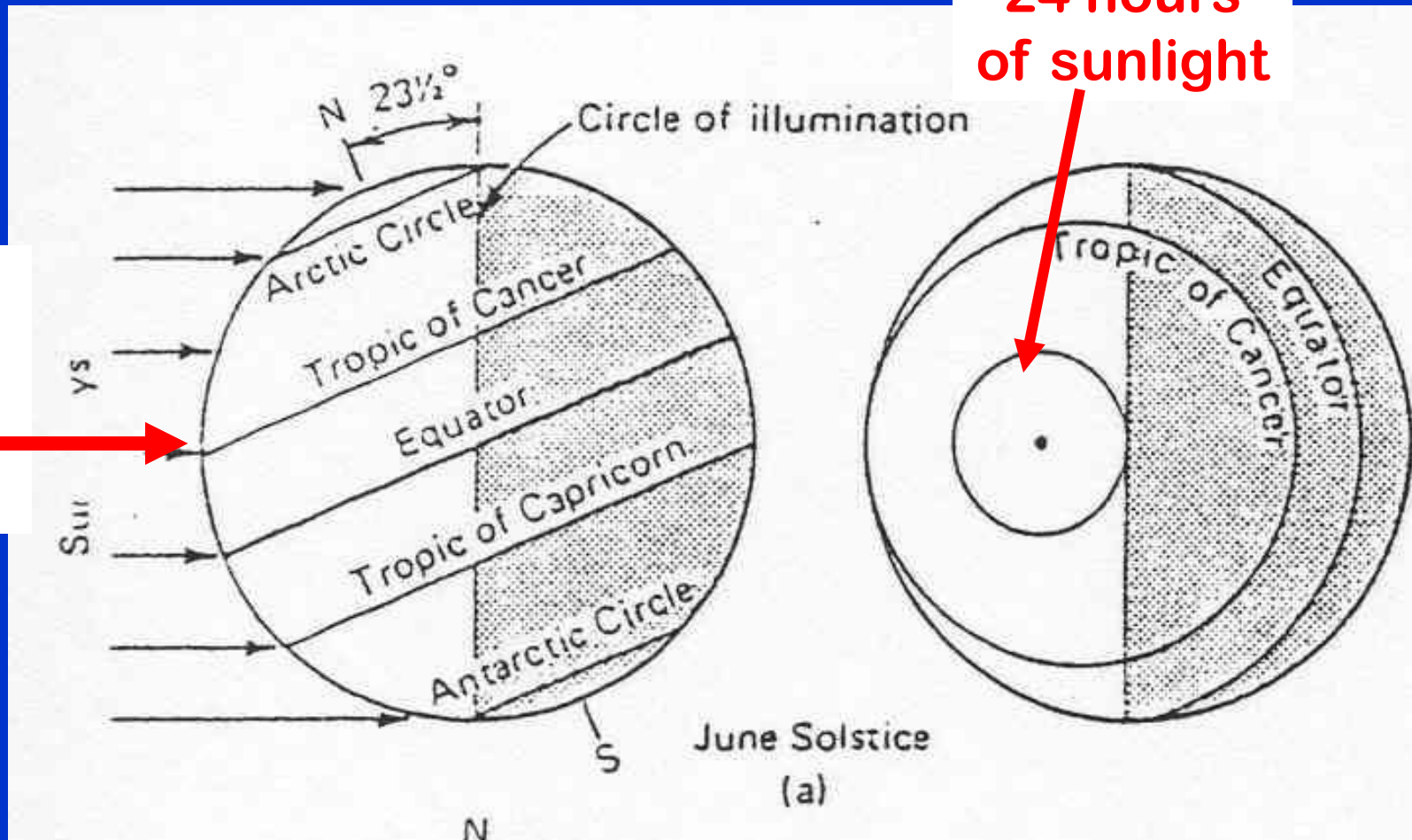
JUNE SOLSTICE



JUNE SOLSTICE

24 hours
of sunlight

Most
intense
solar
radiation



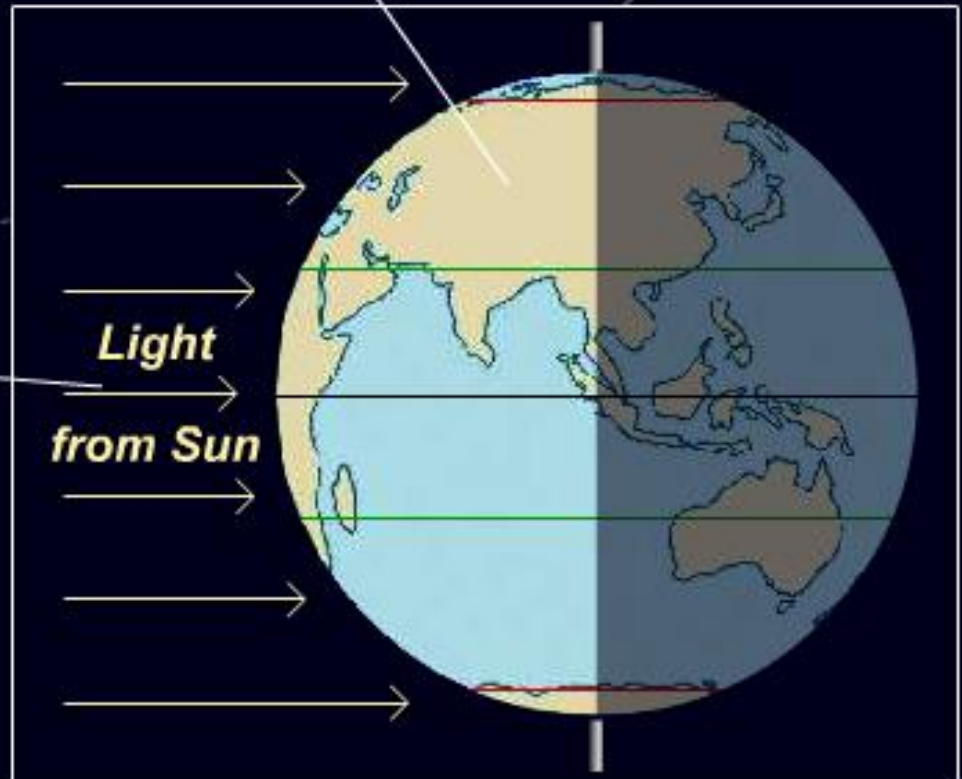
MARCH EQUINOX

**Equinox =
“equal night”**

All locations on
Earth experience
12 hours of daylight

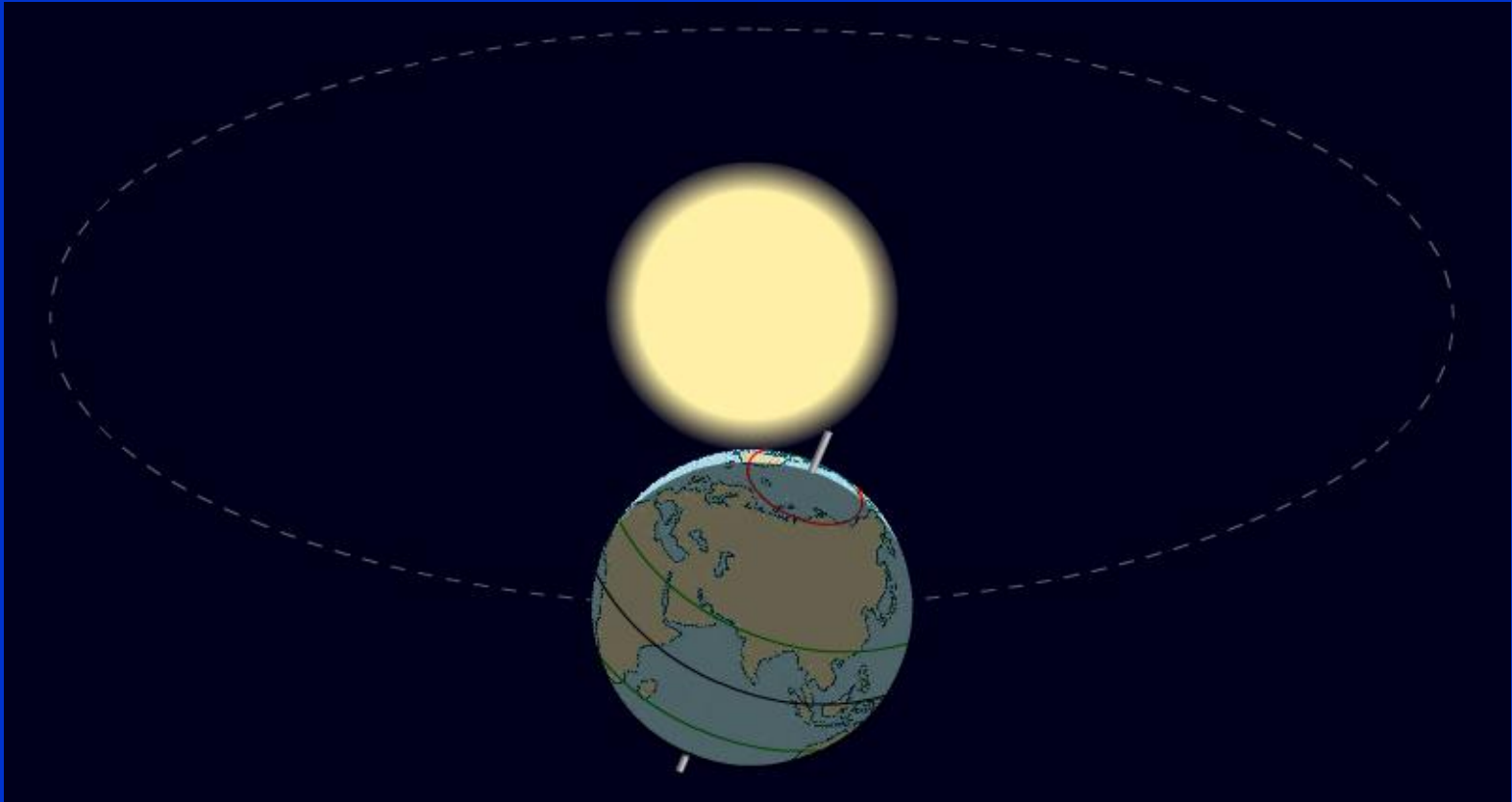
Vertical rays of Sun
striking equator

*Light
from Sun*



SEPTEMBER EQUINOX

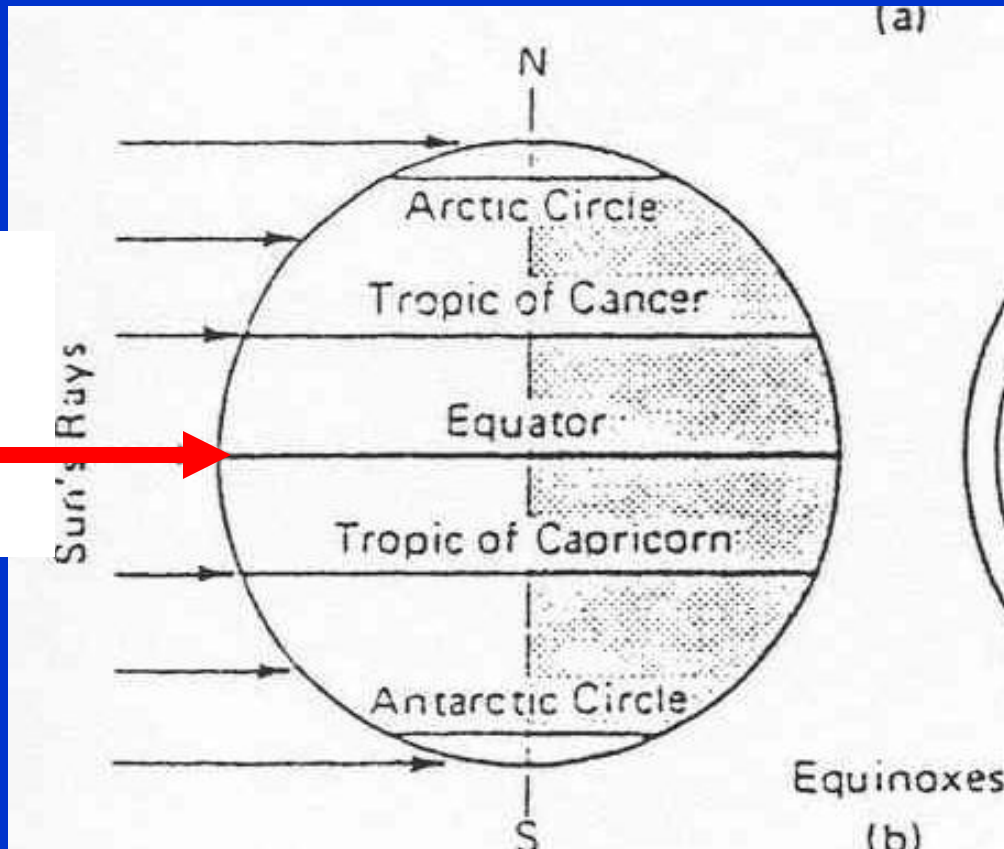
different seasonal position in orbit . . .



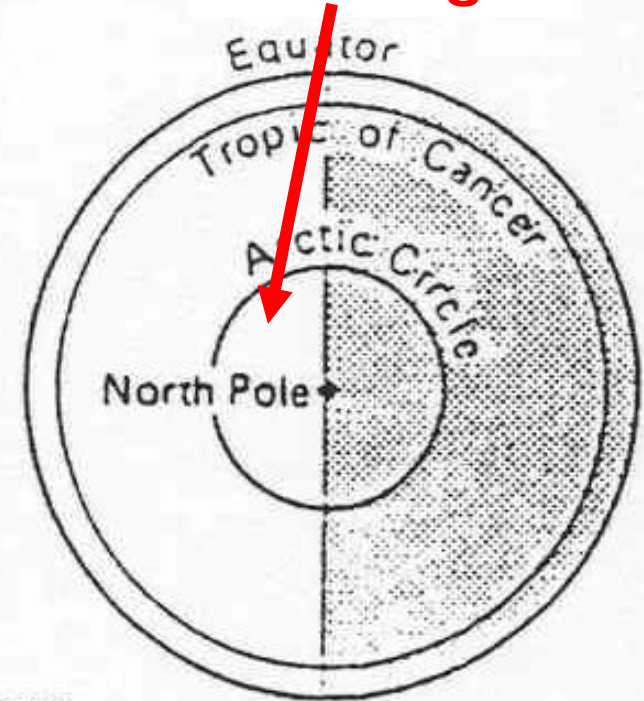
. . . but same latitudinal
insolation as March Equinox

MARCH & SEPTEMBER EQUINOXES

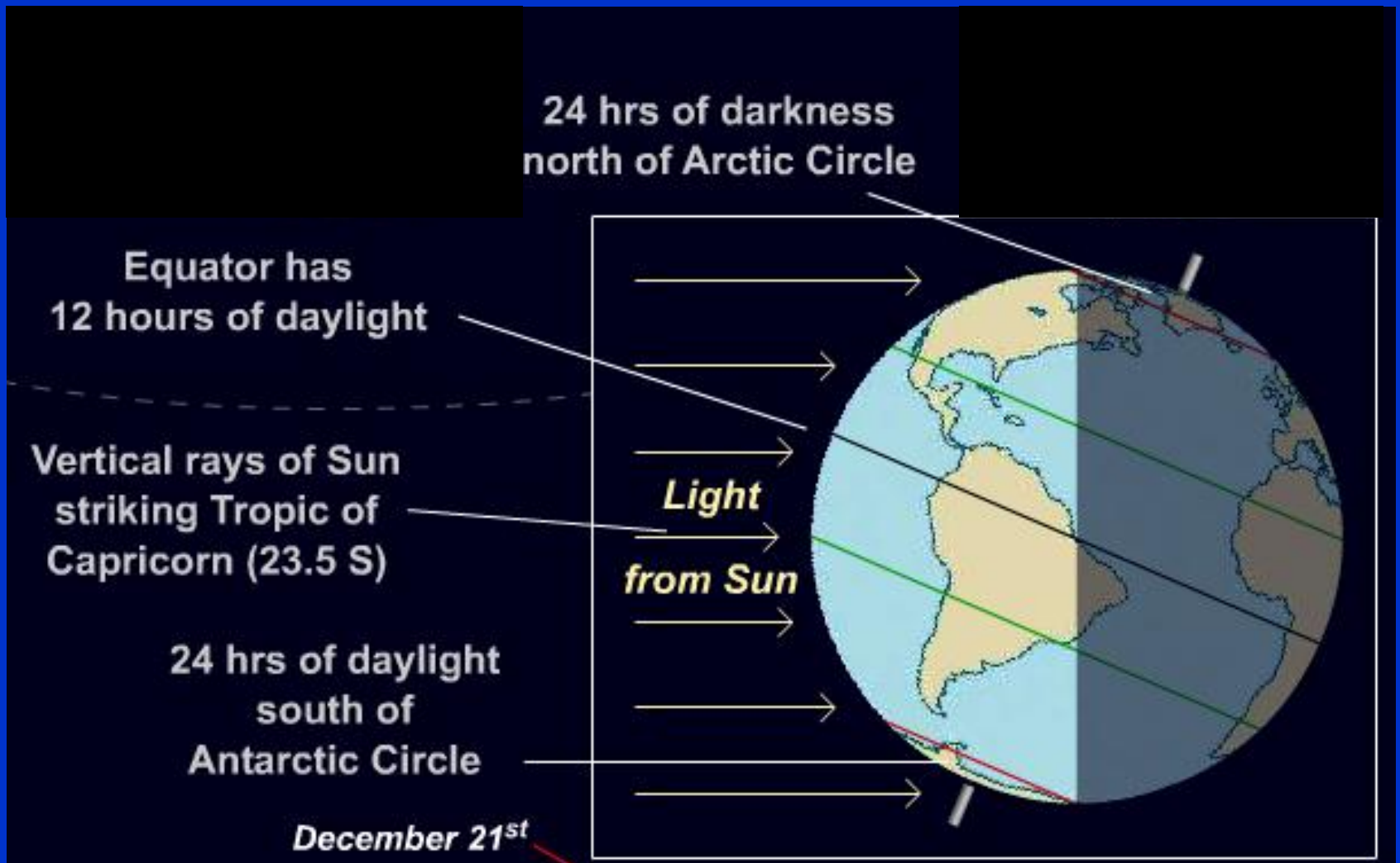
Most
intense
solar
radiation



12 hours
of sunlight



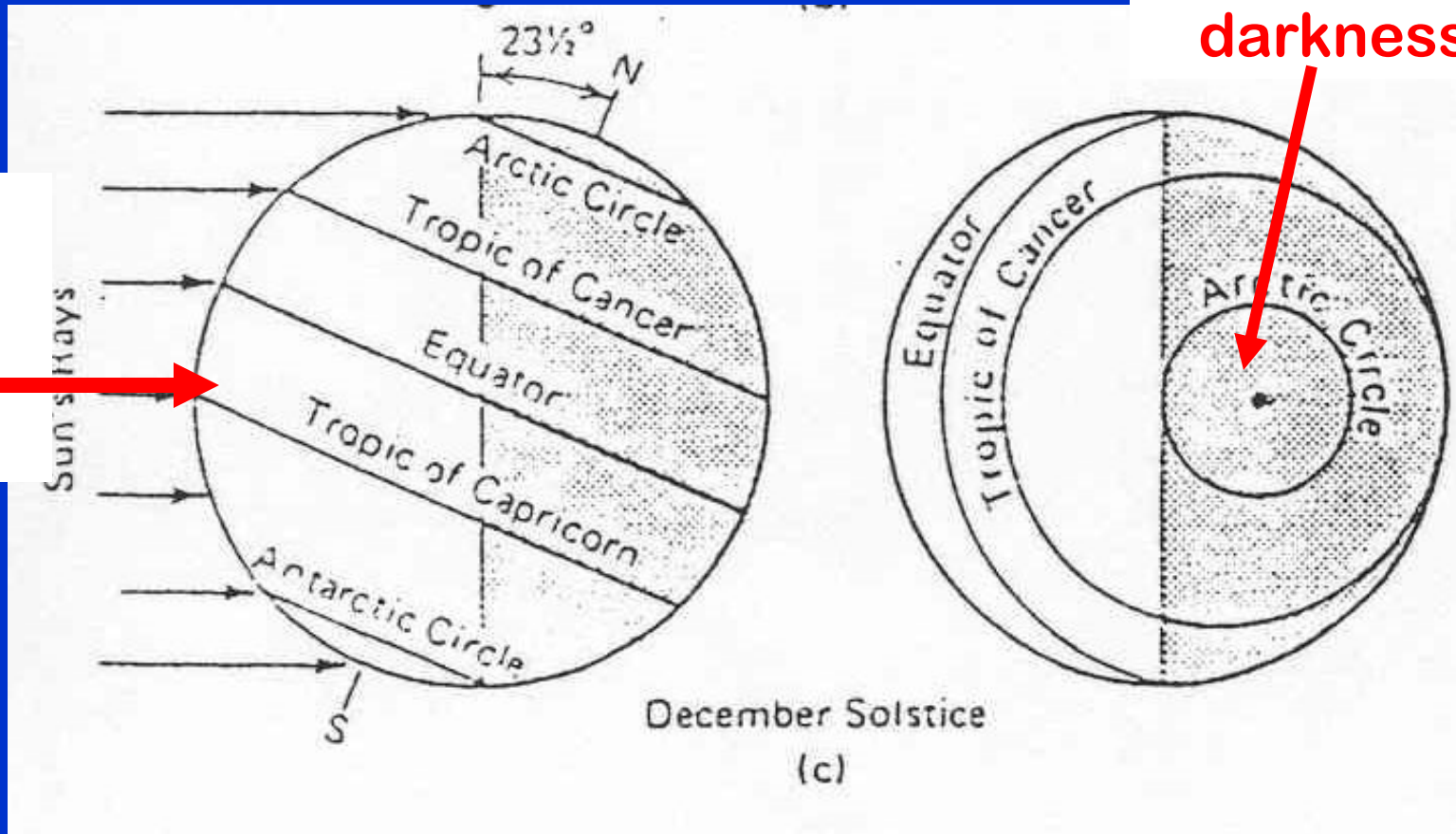
DECEMBER SOLSTICE



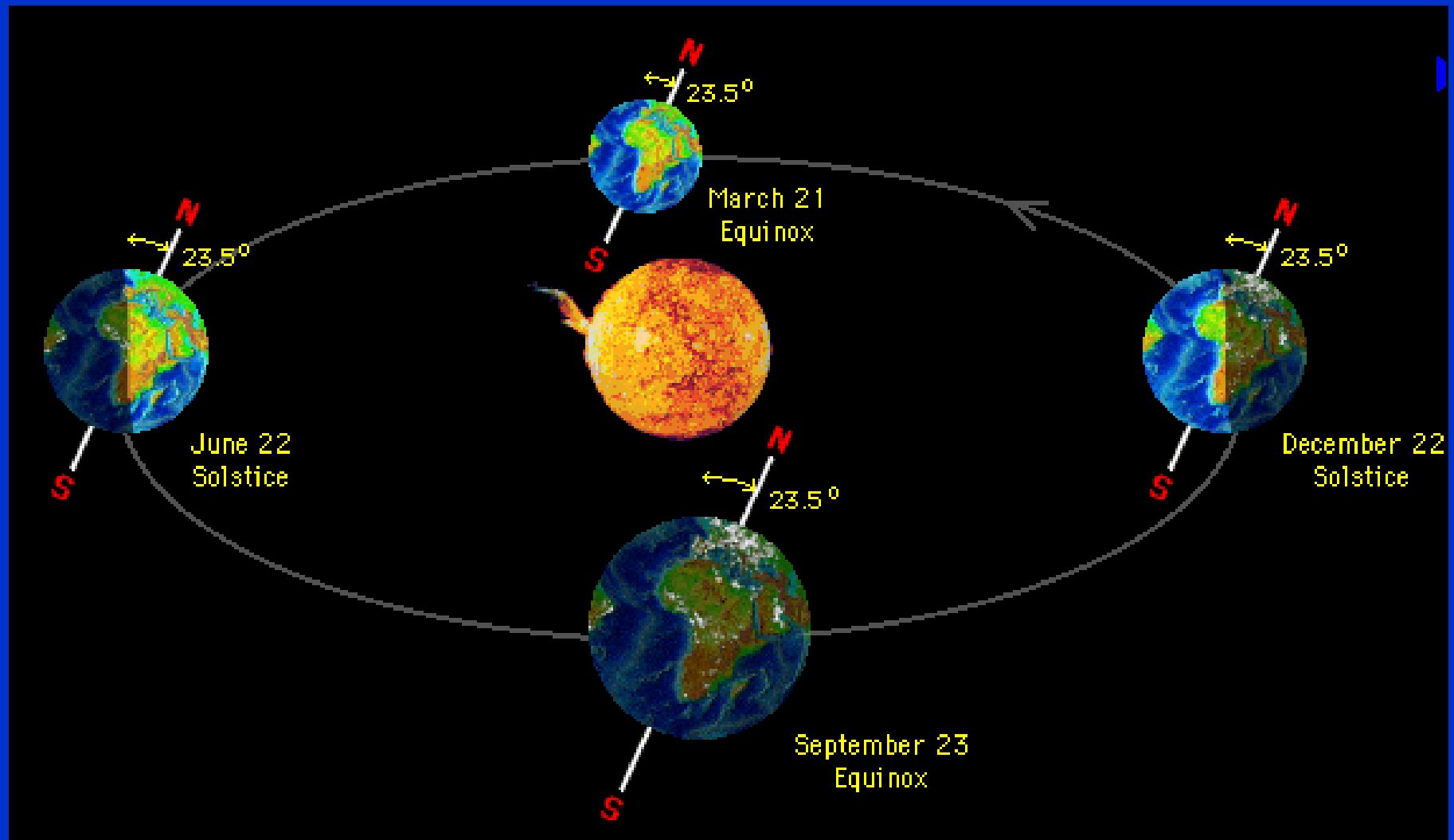
DECEMBER SOLSTICE

24 hours of
darkness

Most
intense
solar
radiation

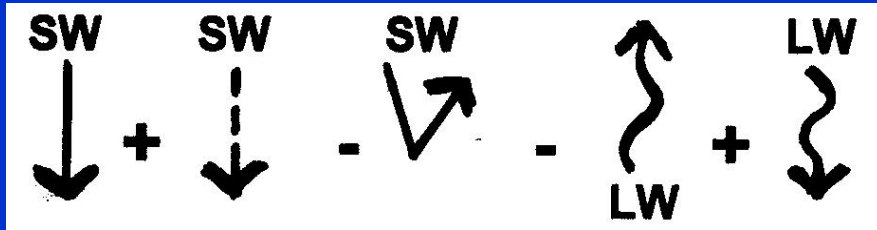


http://mesoscale.agron.iastate.edu/agron206/animations/01_EarthSun.html



Recap

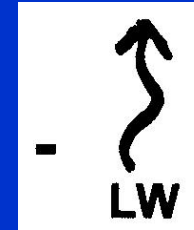
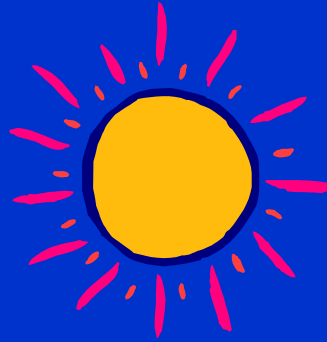
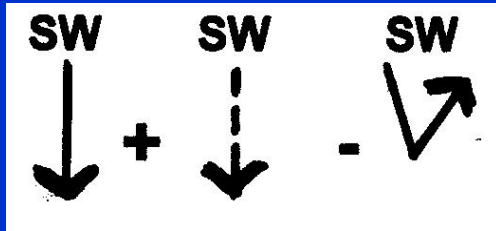
THE RADIATION BALANCE



**& THE GENERAL
CIRCULATION OF THE
ATMOSPHERE**



HOW IT ALL FITS TOGETHER:



Over the course of a year . . .

The amount of **INCOMING SW** (Insolation) absorbed by EARTH **varies by LATITUDE**

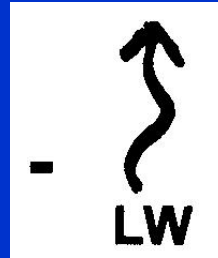
(MORE comes in near the Equator, less near the Poles)

→ **LOW LATITUDES**
absorb more energy
than **HIGH LATITUDES**

The amount of **TERRESTRIAL LW / IR** varies by latitude too -- **MORE LW / IR** is emitted at warmer **TROPICAL LATITUDES**, **LESS** in cooler **HIGH LATITUDES**

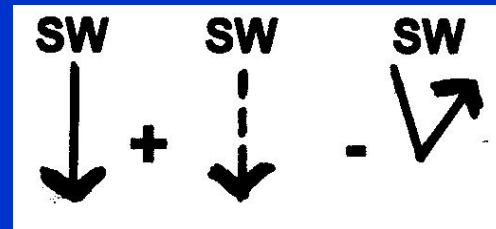
HOWEVER . . .

**EQUATOR-POLE
DIFFERENCES in
what goes OUT**

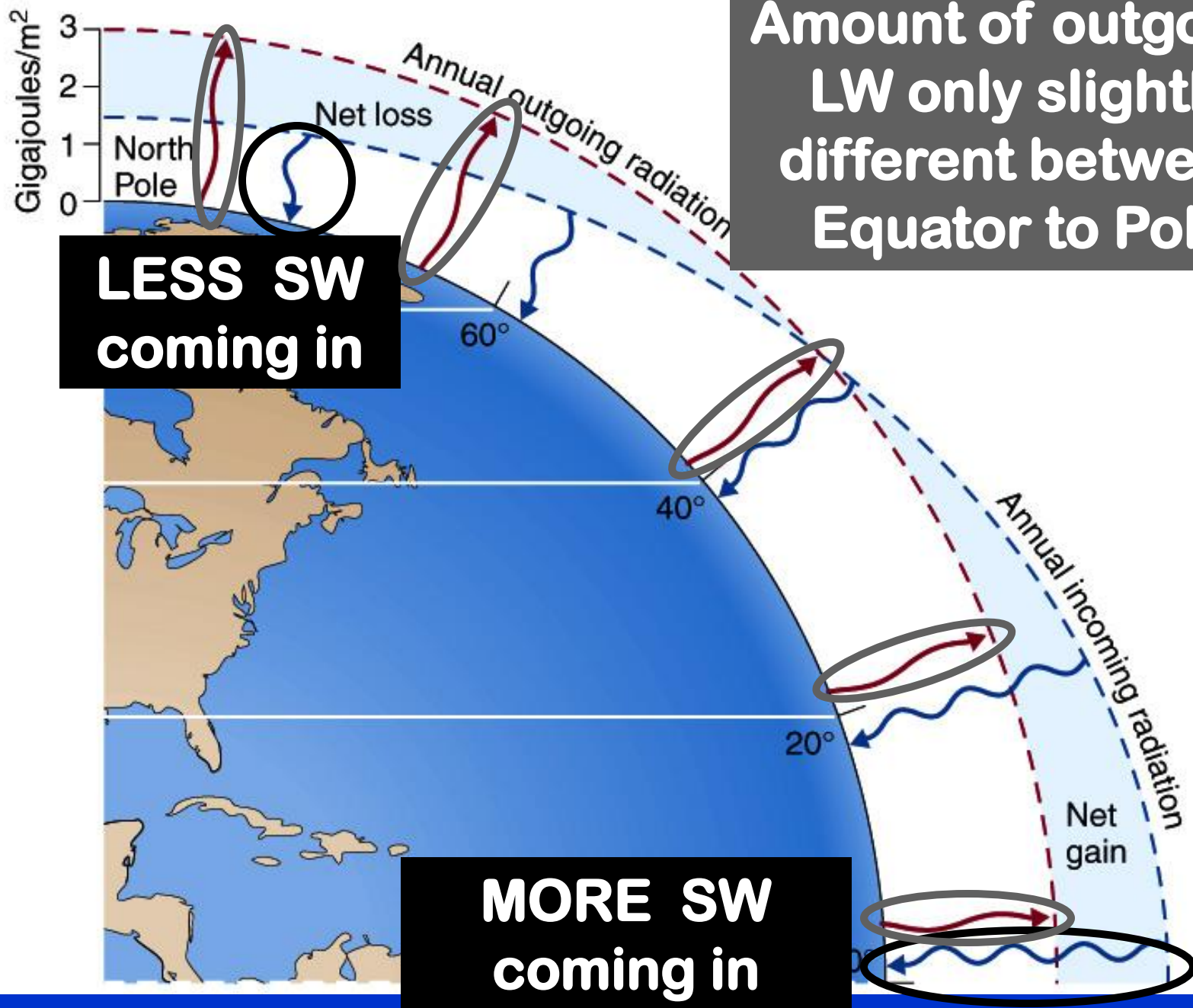


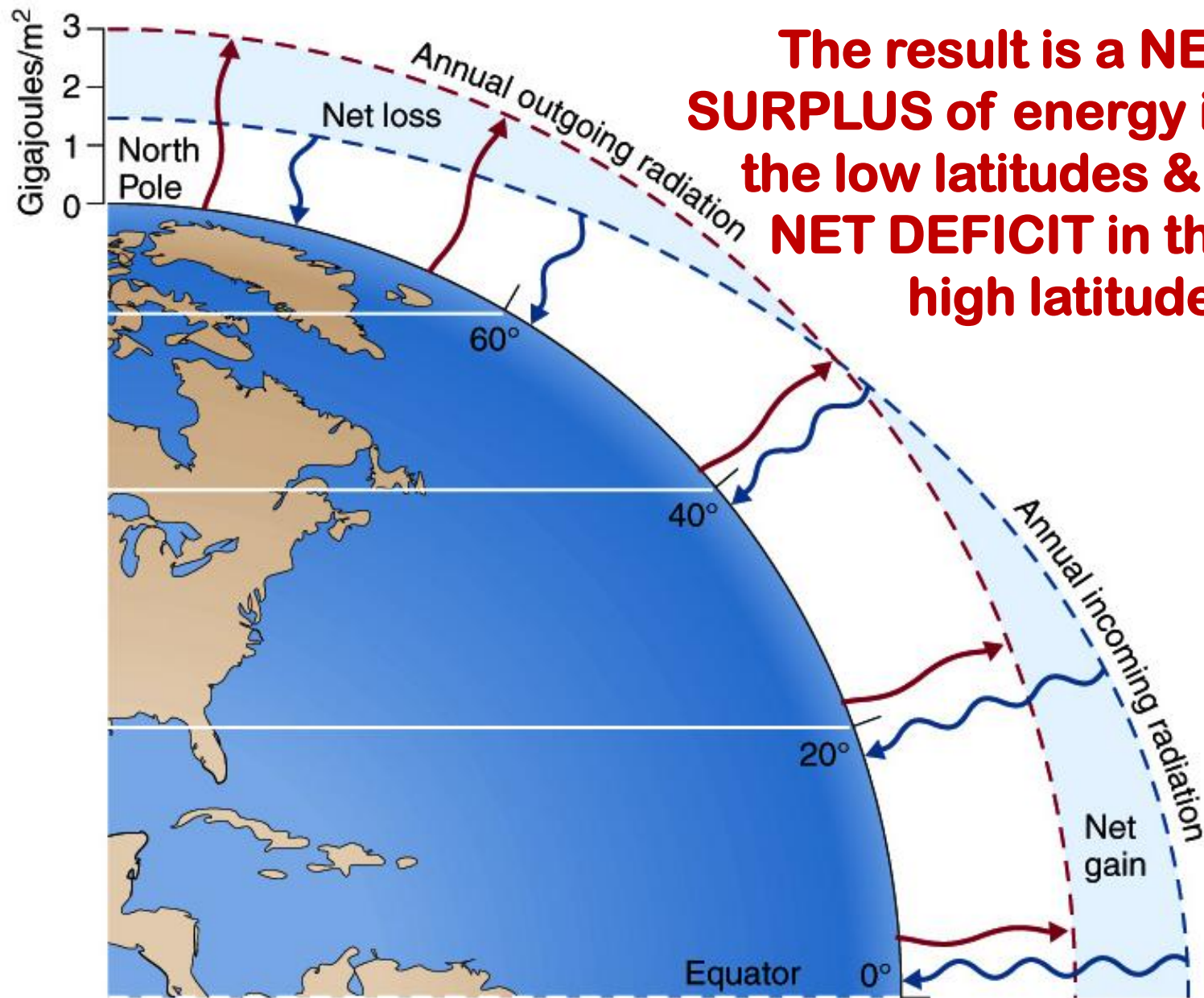
are less than the

**EQUATOR-POLE
DIFFERENCES in
what comes in**

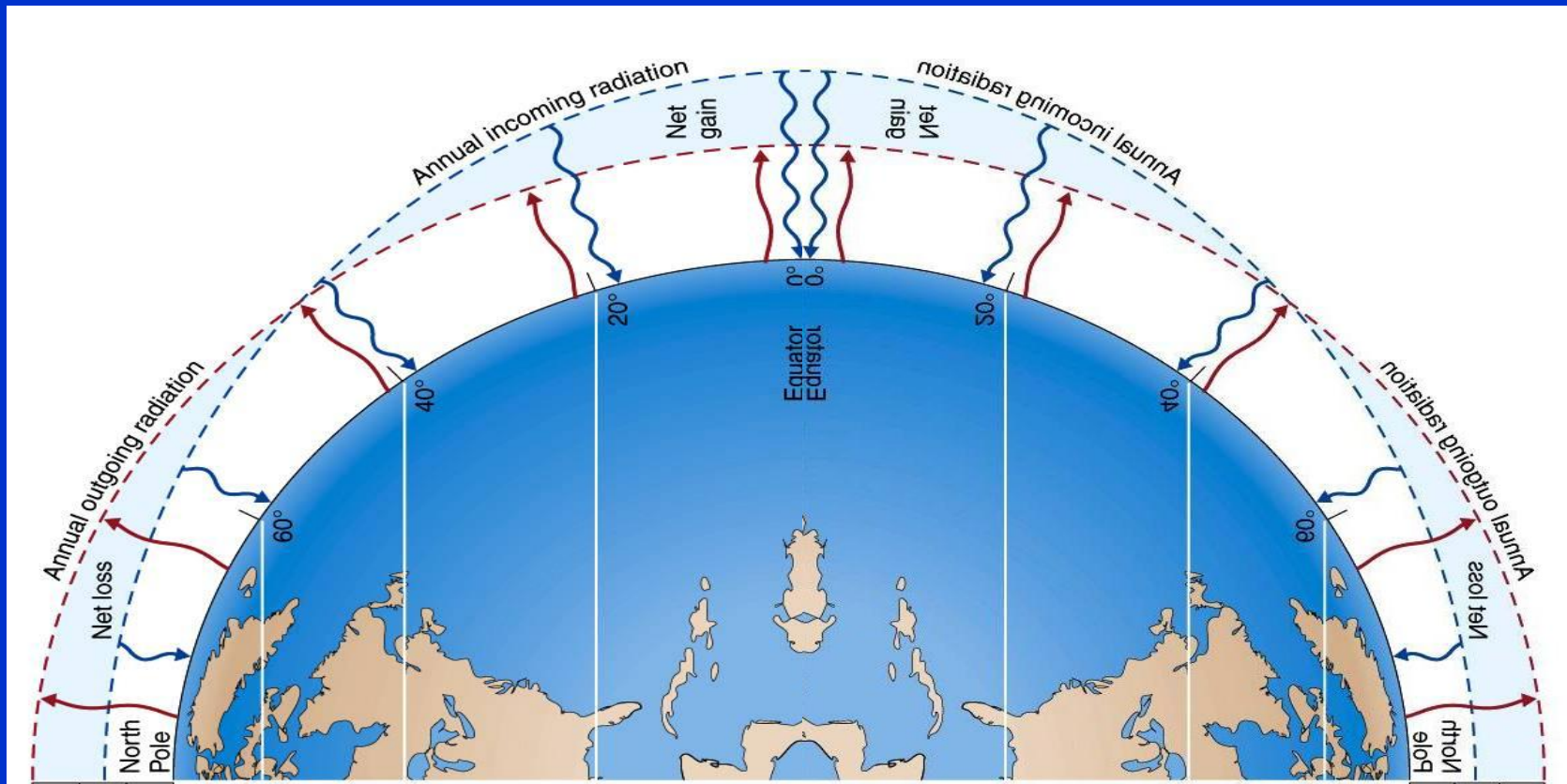


Amount of outgoing
LW only slightly
different between
Equator to Pole





The result is a NET SURPLUS of energy in the low latitudes & a NET DEFICIT in the high latitudes

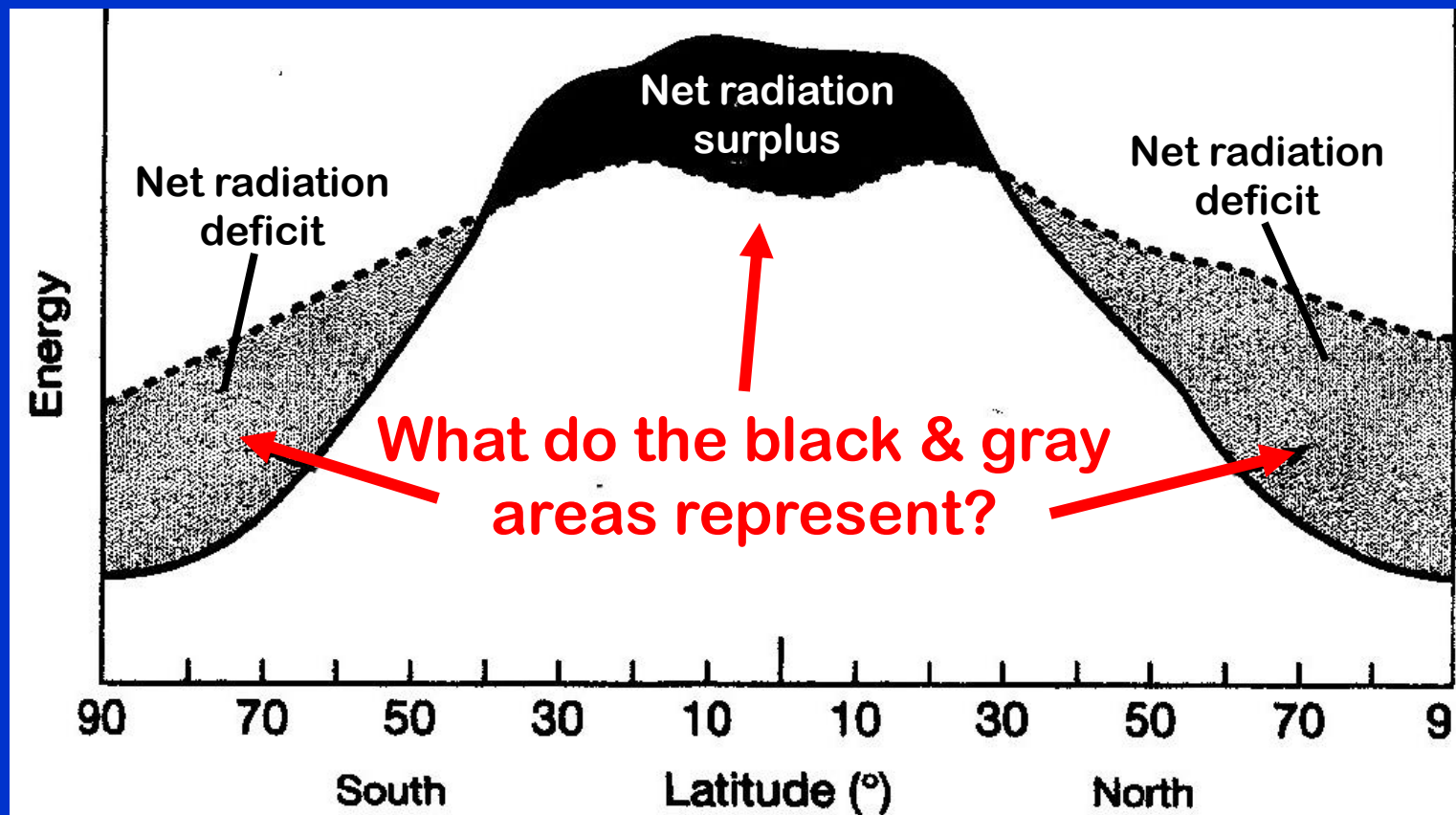


POLE

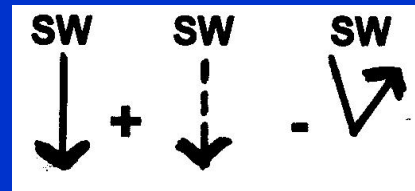
EQUATOR

POLE

Now lets look at a
Pole to Pole Transect



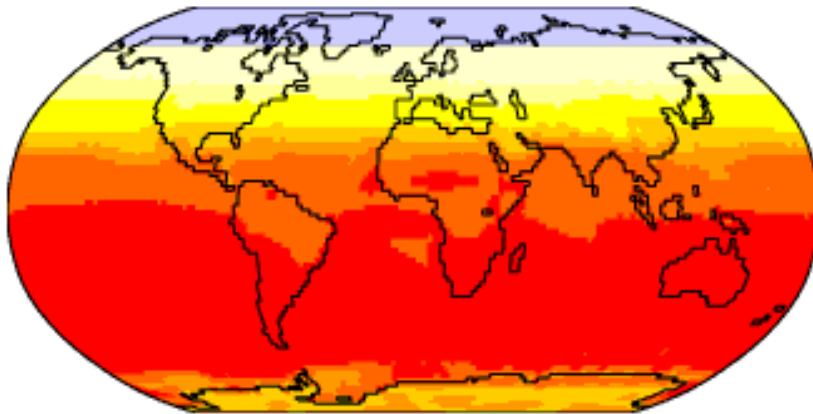
———— Absorbed solar energy



----- Emitted infrared energy
(at top of atmosphere)

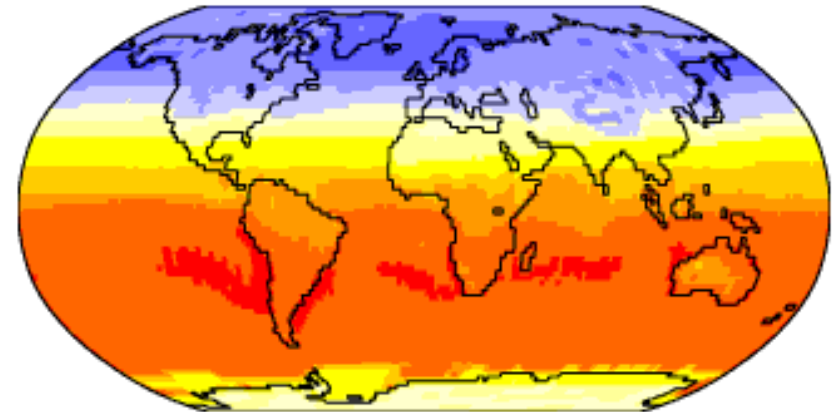


Short-Wave Radiation

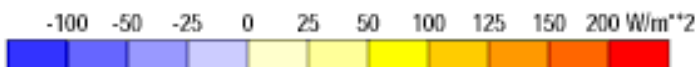
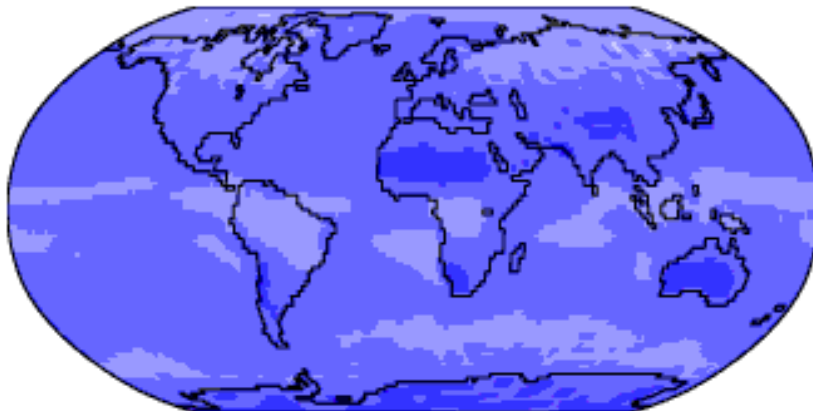


Dec

Net Radiation



Long-Wave Radiation

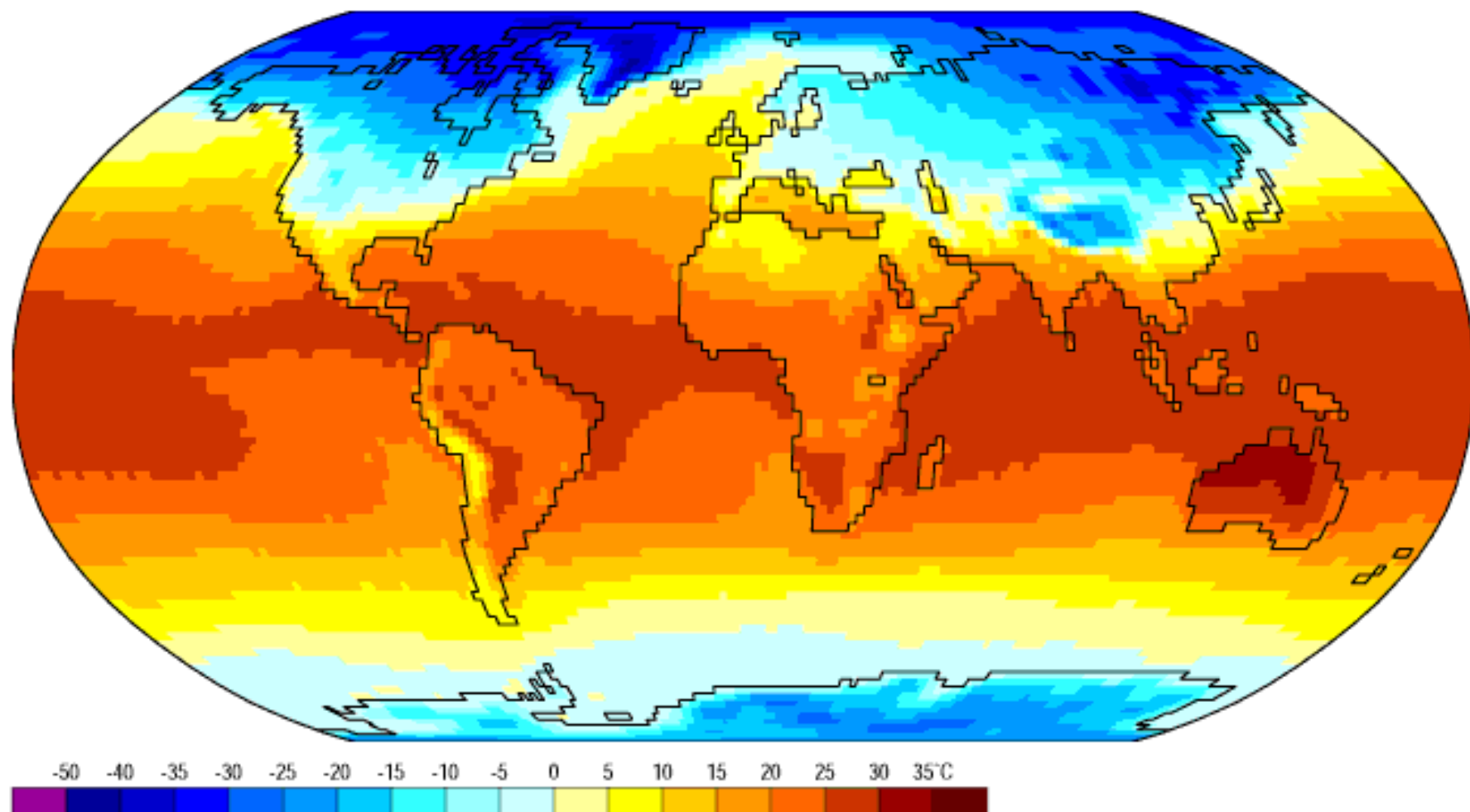


Data: NCEP/NCAR Reanalysis Project, 1959-1997 Climatologies
Animation: Department of Geography, University of Oregon, March 2000

http://geography.uoregon.edu/envchange/clim_animations/

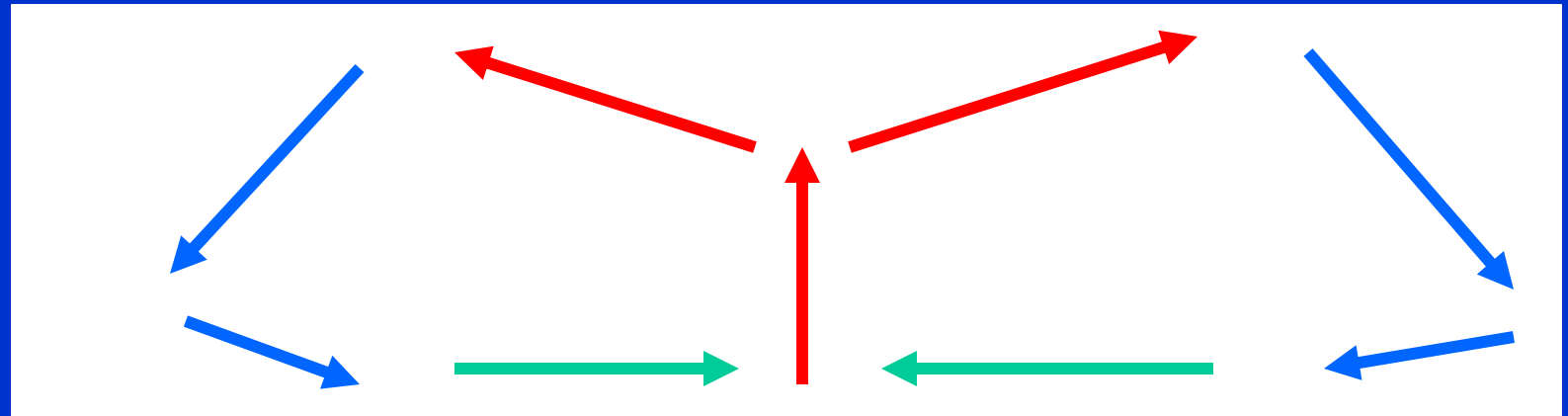
Air Temperature

Dec



Data: NCEP/NCAR Reanalysis Project, 1959-1997 Climatologies
Animation: Department of Geography, University of Oregon, March 2000

Global-scale motions driven by thermal differences:



Northern Hemisphere

EQUATOR

Southern Hemisphere

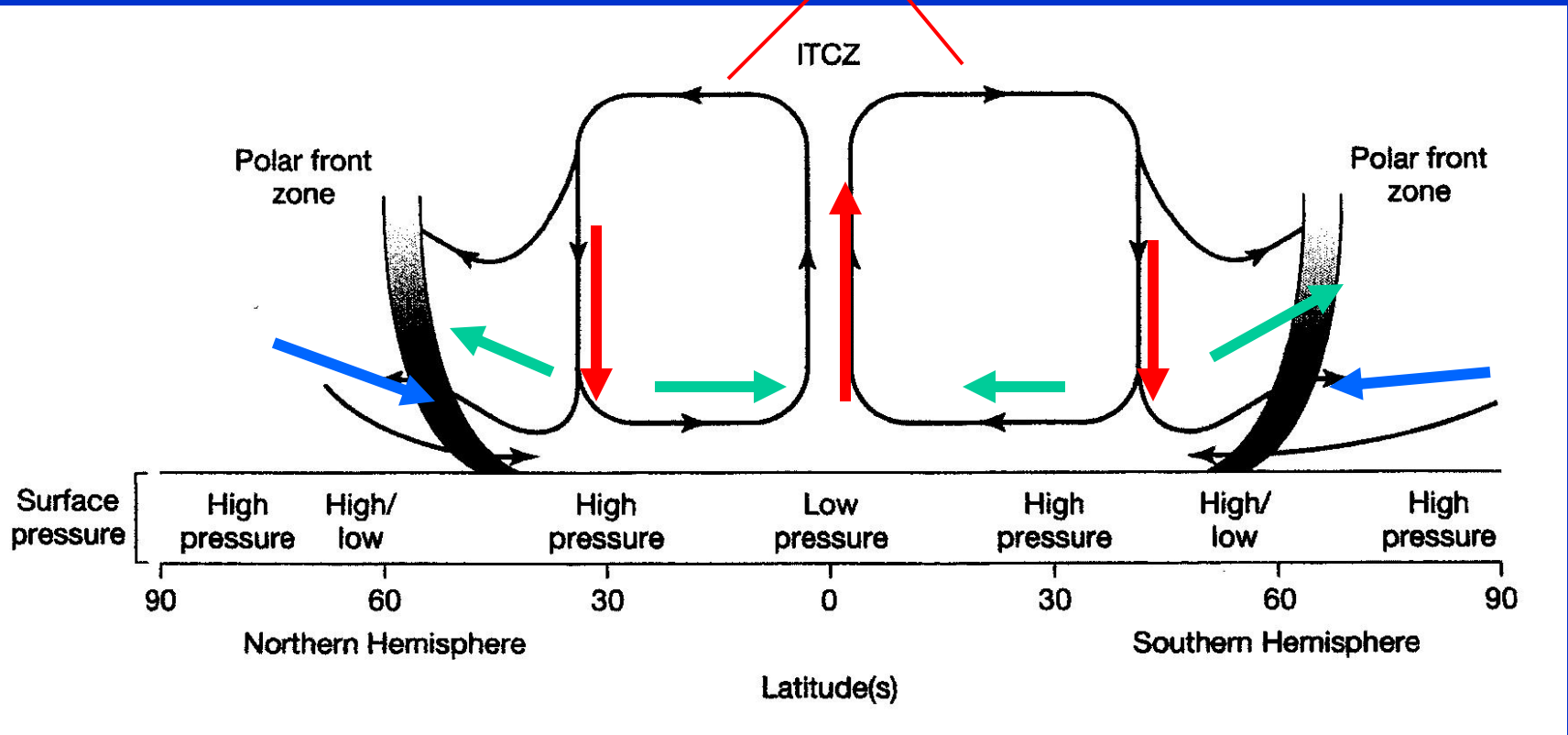
COLD
POLAR
REGIONS

HOT
TROPICS

COLD
POLAR
REGIONS



HADLEY CELLS



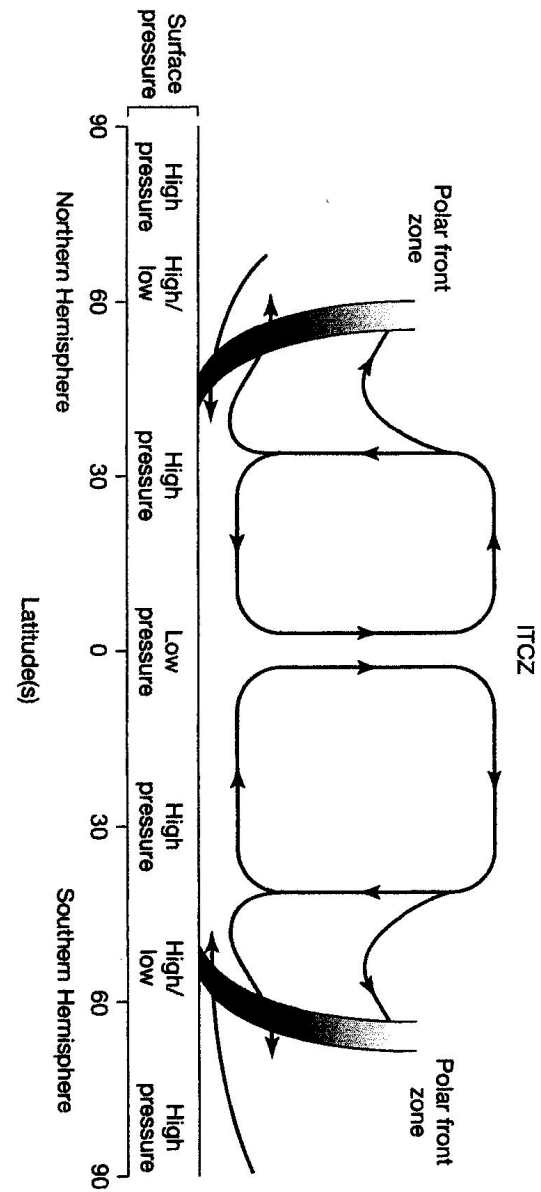
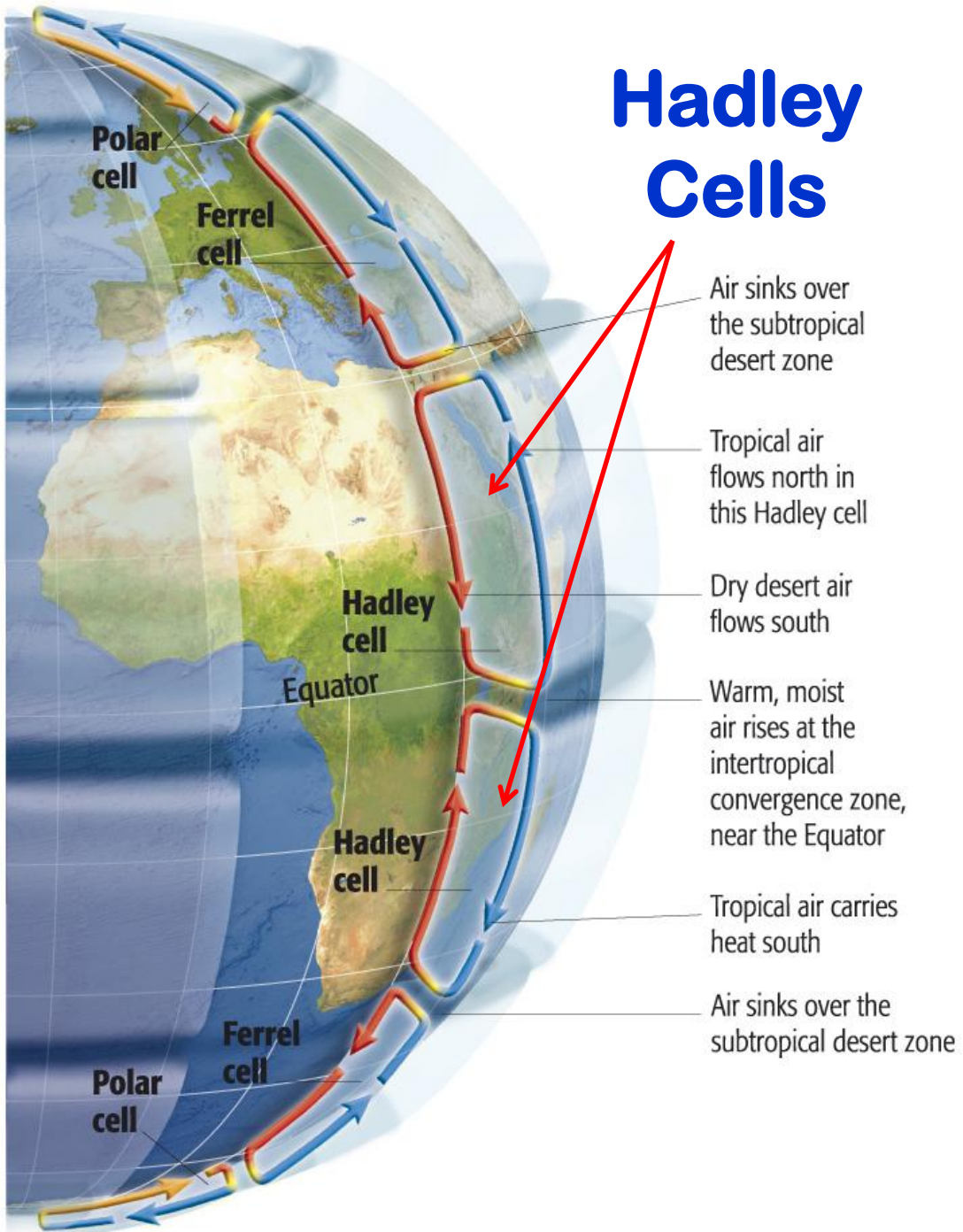
**COLD
POLAR
REGIONS**

**HOT
TROPICS**

**COLD
POLAR
REGIONS**

From SGC-I Chapter 4

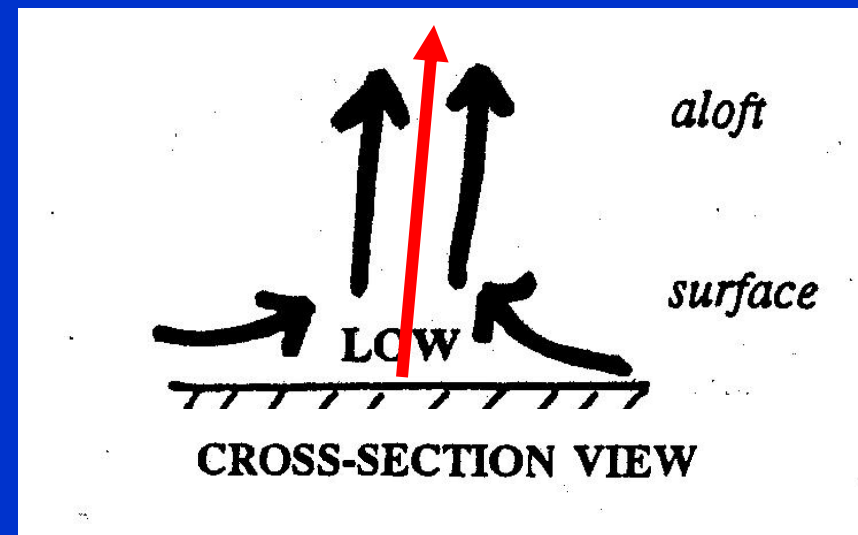
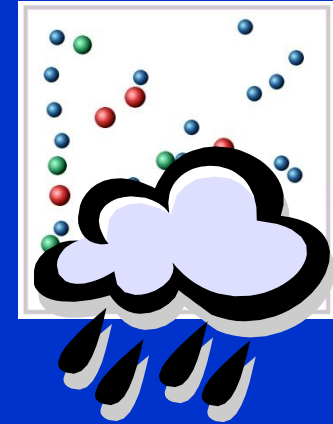
Hadley Cells



LOW PRESSURE AREAS:

Hot surface → Rising
air → **expansion and
cooling** of air, and
condensation
of water vapor →
**clouds and possibly
precipitation . . .**

HUMID REGIONS



DANCE YOUR PH.D!

“Precipitation Initiation in Warm Clouds”



This dance shows **how a rain drop can form** when one slightly larger rain drop is present among a large population of smaller drops. The large drop only forms after mixing occurs.

http://www.youtube.com/watch?v=4O7G7F_e7I0

Condensation
nucleus

H₂O droplet



Here the **women** represent **water molecules** while the **men** represent **cloud condensation nuclei**.

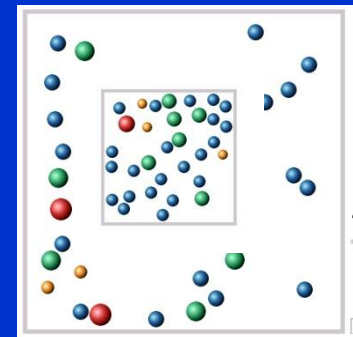
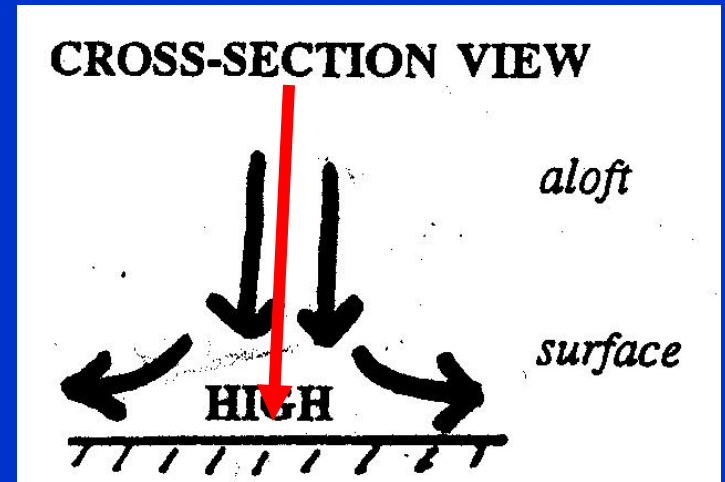


Additional women dance in a manner suggestive of mixing processes in order to **create the slightly larger drop**, called a collision coalescence initiator.

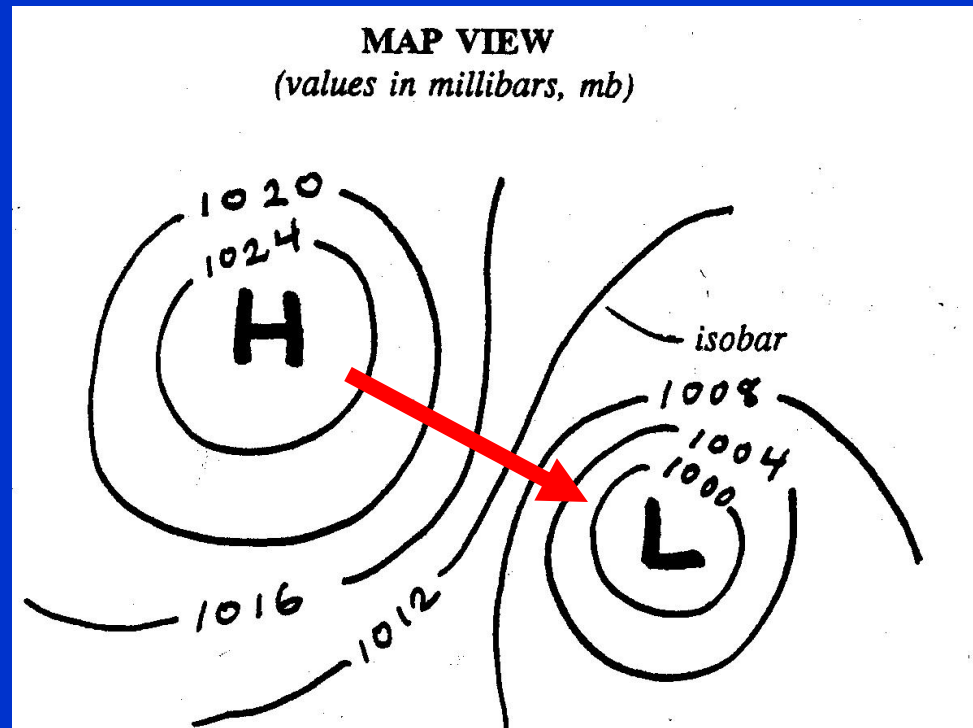
Through “coalescence” a single nucleus attracts all the other water droplets -- when large enough **RAIN FALLS!**

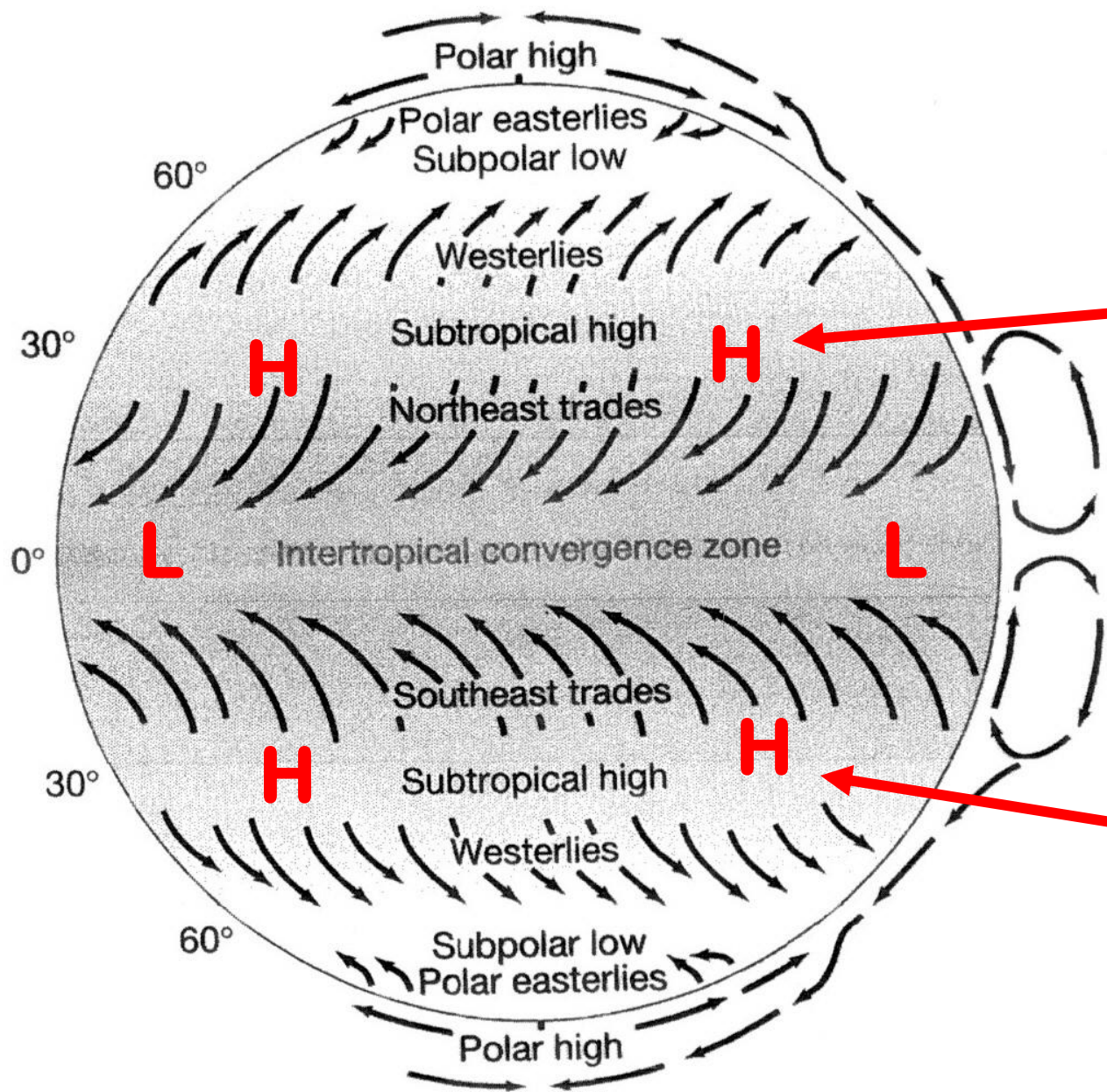
HIGH PRESSURE AREAS:

Forced sinking
(e.g. in HADLEY CELL)
leads to **contraction**
and warming of air, and
increased water vapor
holding capacity →
clear skies, dry air and
ARID REGIONS.



In general: Winds tend to flow from
HIGH → LOW Pressure areas

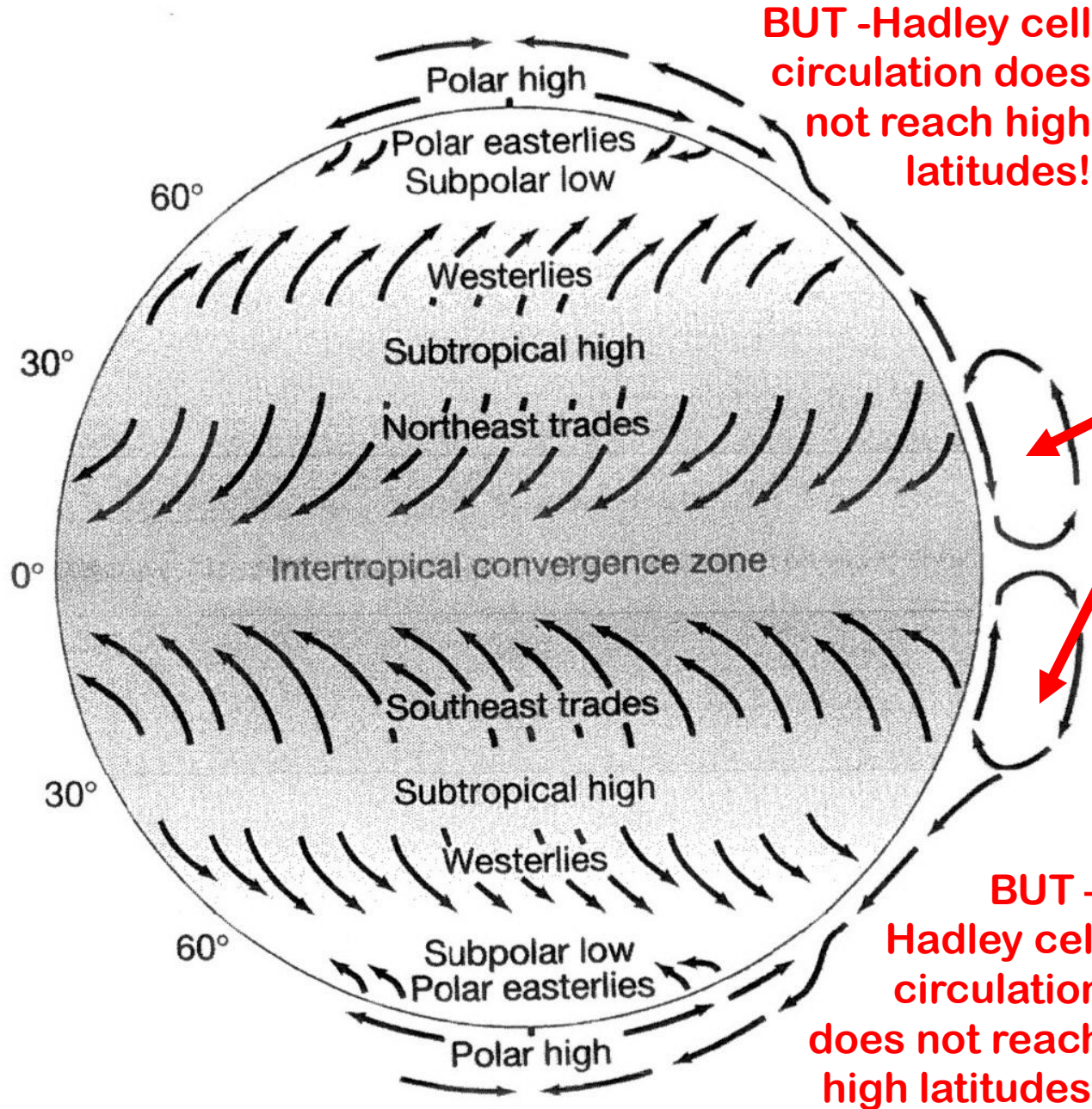




**Sub-
tropical
HIGH
PRESSURE**

**Intertropical
Convergence
ITCZ**

**Sub-
tropical
HIGH
PRESSURE**



**BUT -Hadley cell
circulation does
not reach high
latitudes!**

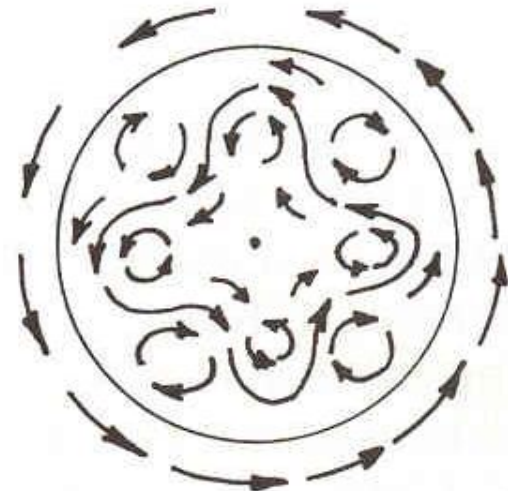
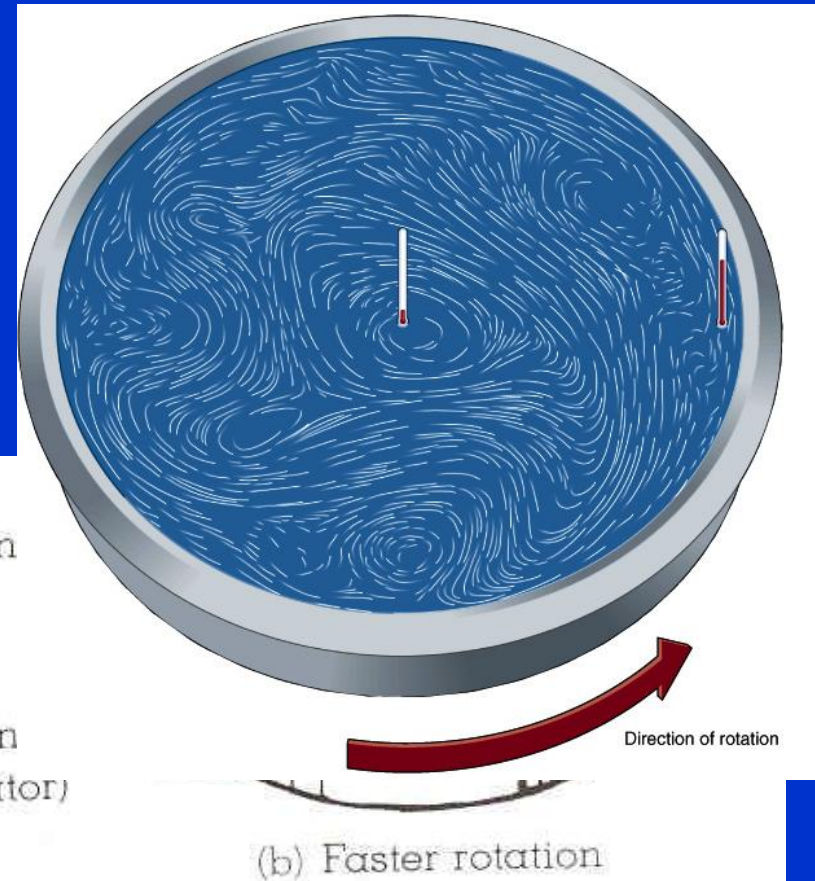
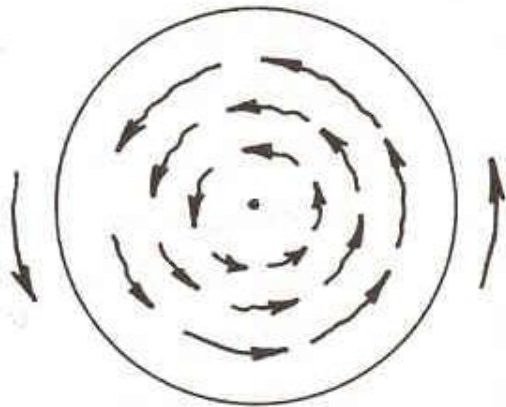
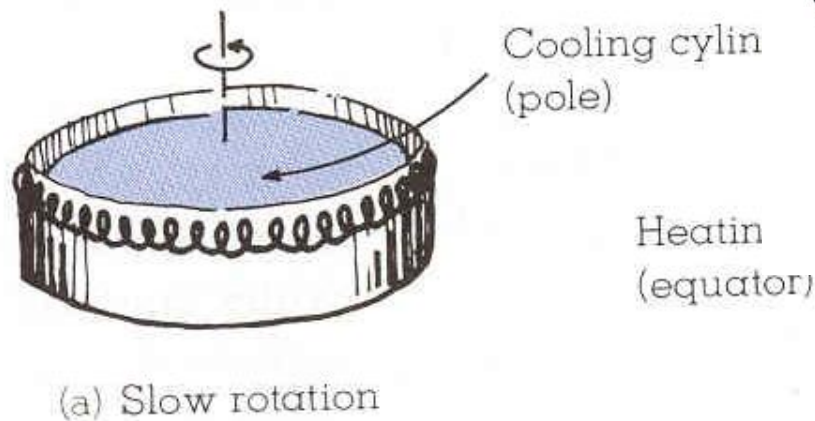
H

**HADLEY
CELLS =
key drivers!**

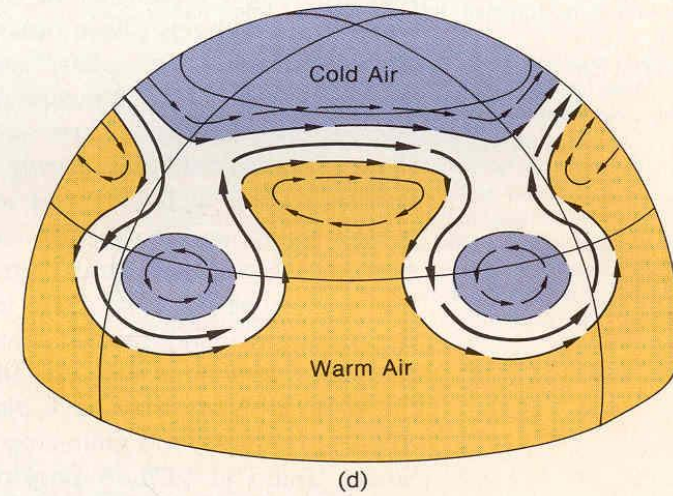
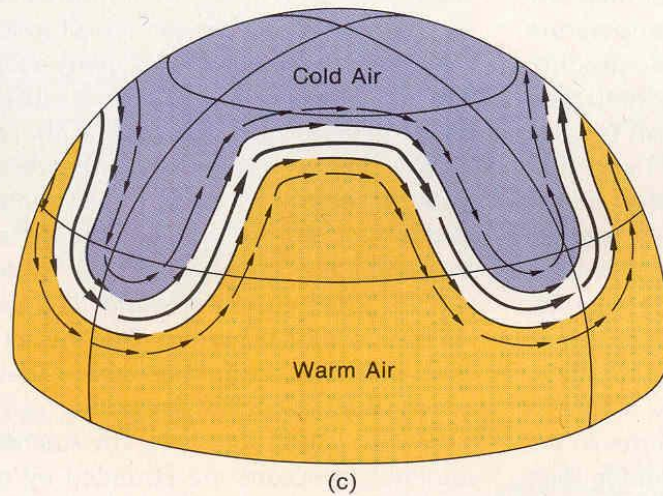
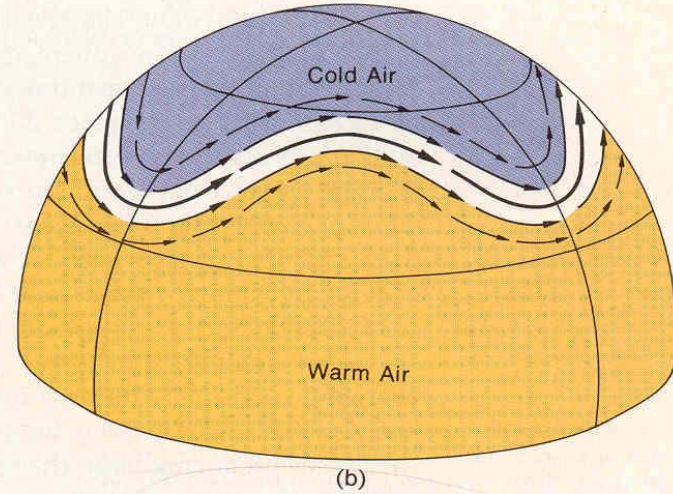
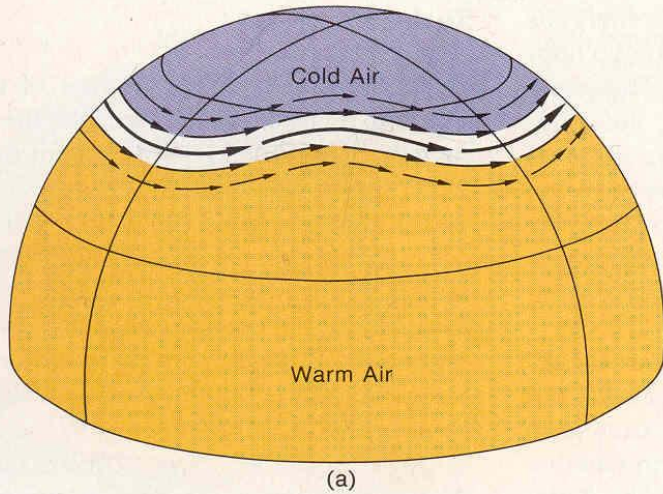
Convection cell
transfer of
thermal energy
from low
latitude area of
energy
SURPLUS to
higher latitude
area of energy
DEFICIT

**BUT –
Hadley cell
circulation
does not reach
high latitudes!**

Why Hadley convective cell transport breaks down at higher latitudes:



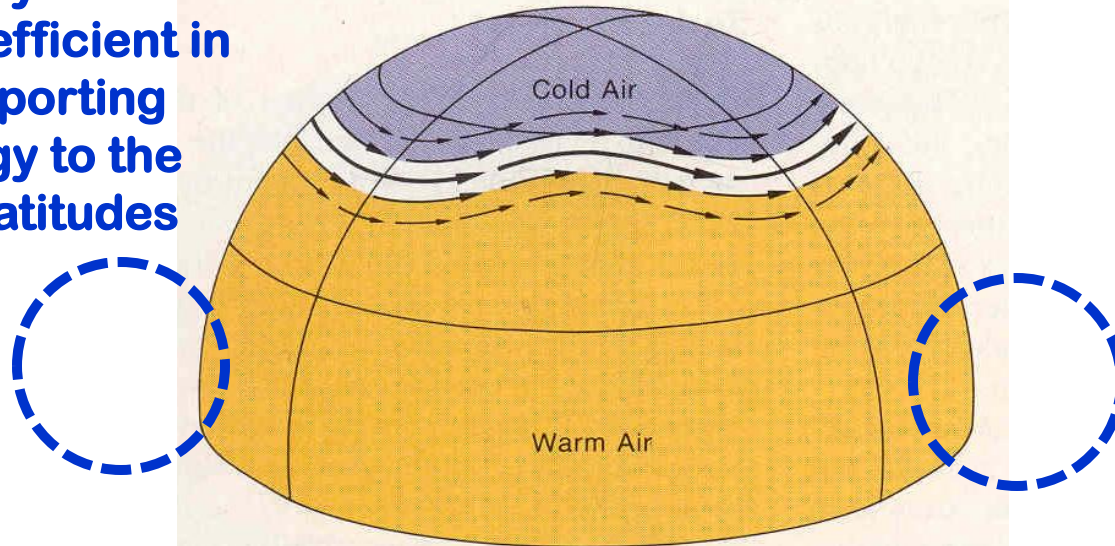
UPPER LEVEL CIRCUMPOLAR WINDS !



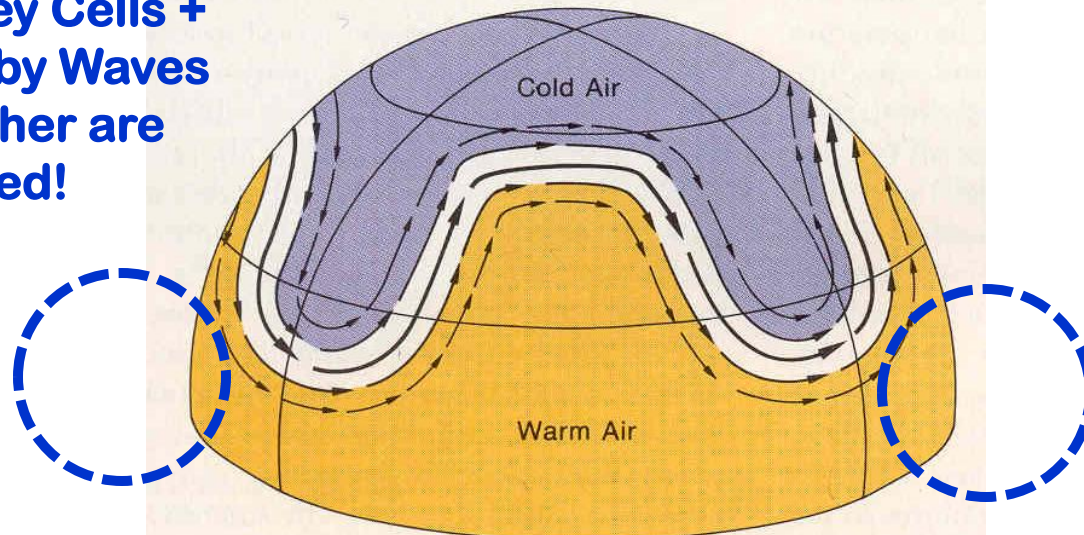
H

**Wave transport of thermal energy
instead of Hadley cell transport!**

Hadley Cells are only efficient in transporting energy to the mid-latitudes

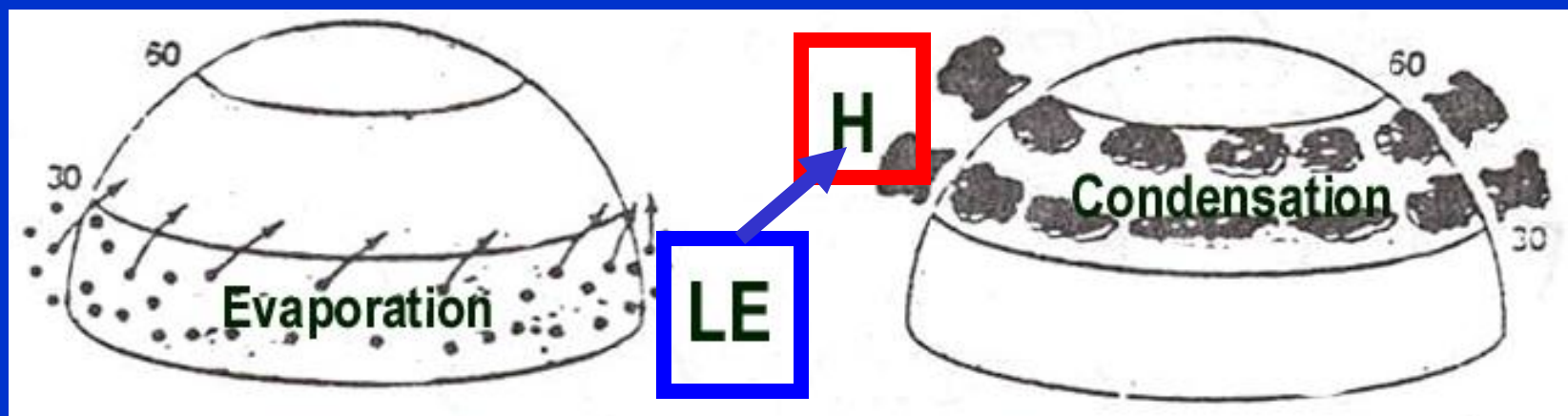
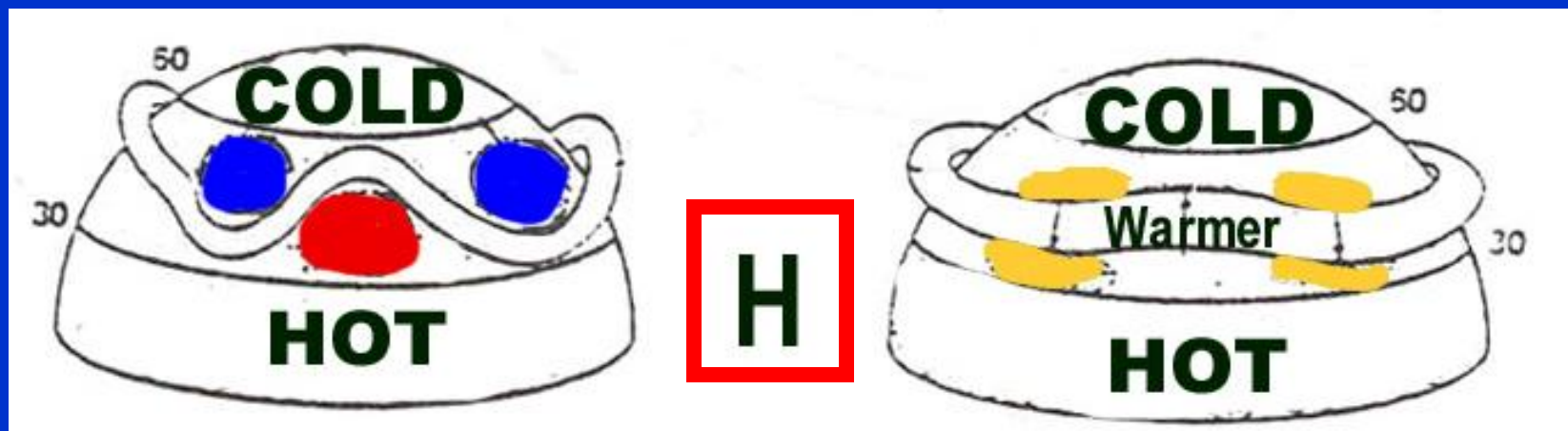


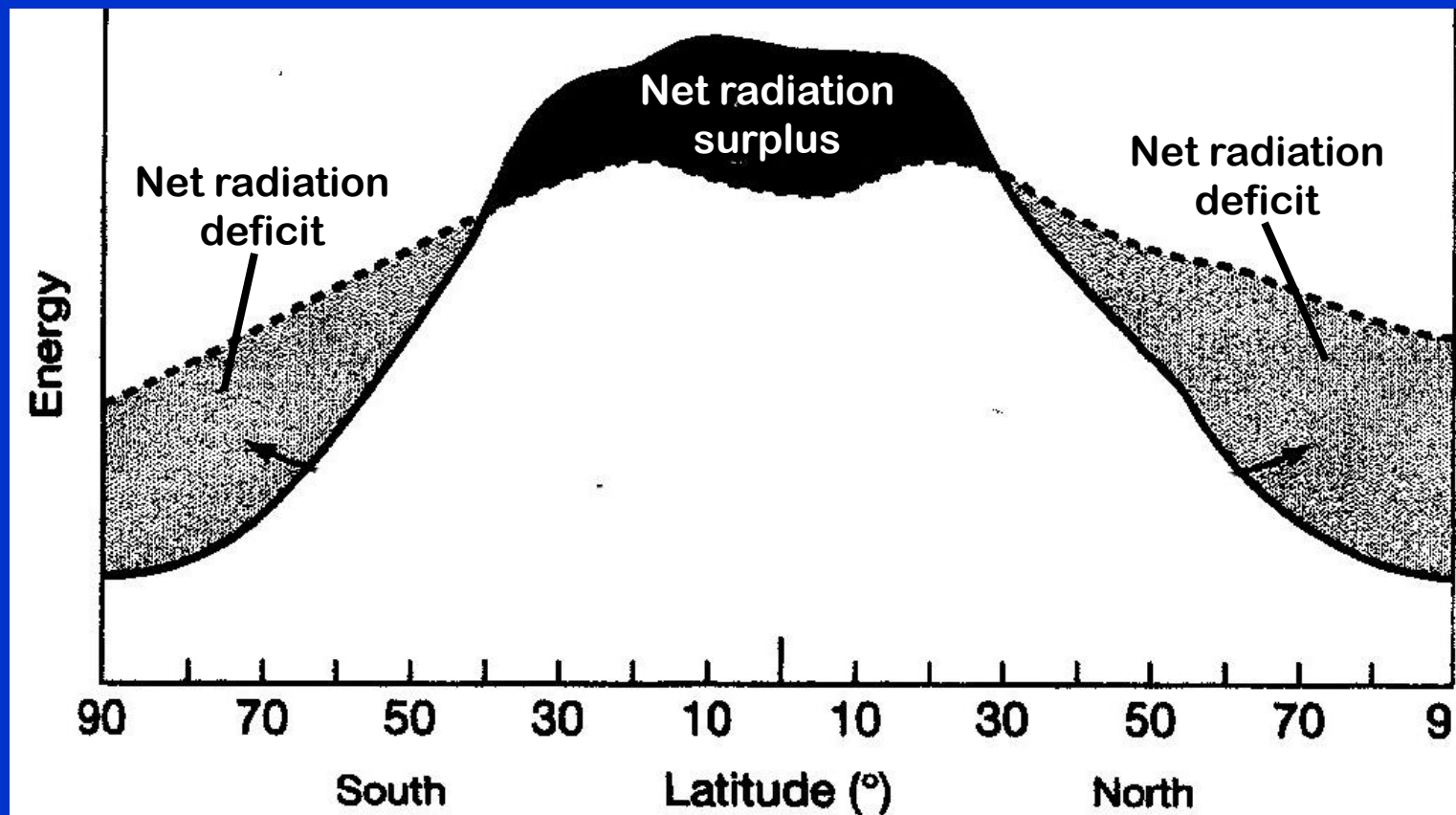
Hadley Cells + Rossby Waves together are needed!



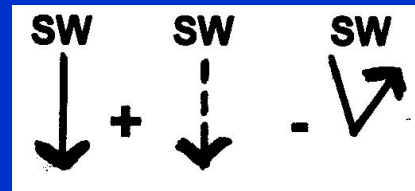
Energy is transported from areas of surplus to deficit in form of: **H (sensible heat) & LE (Latent Energy)**

H + LE



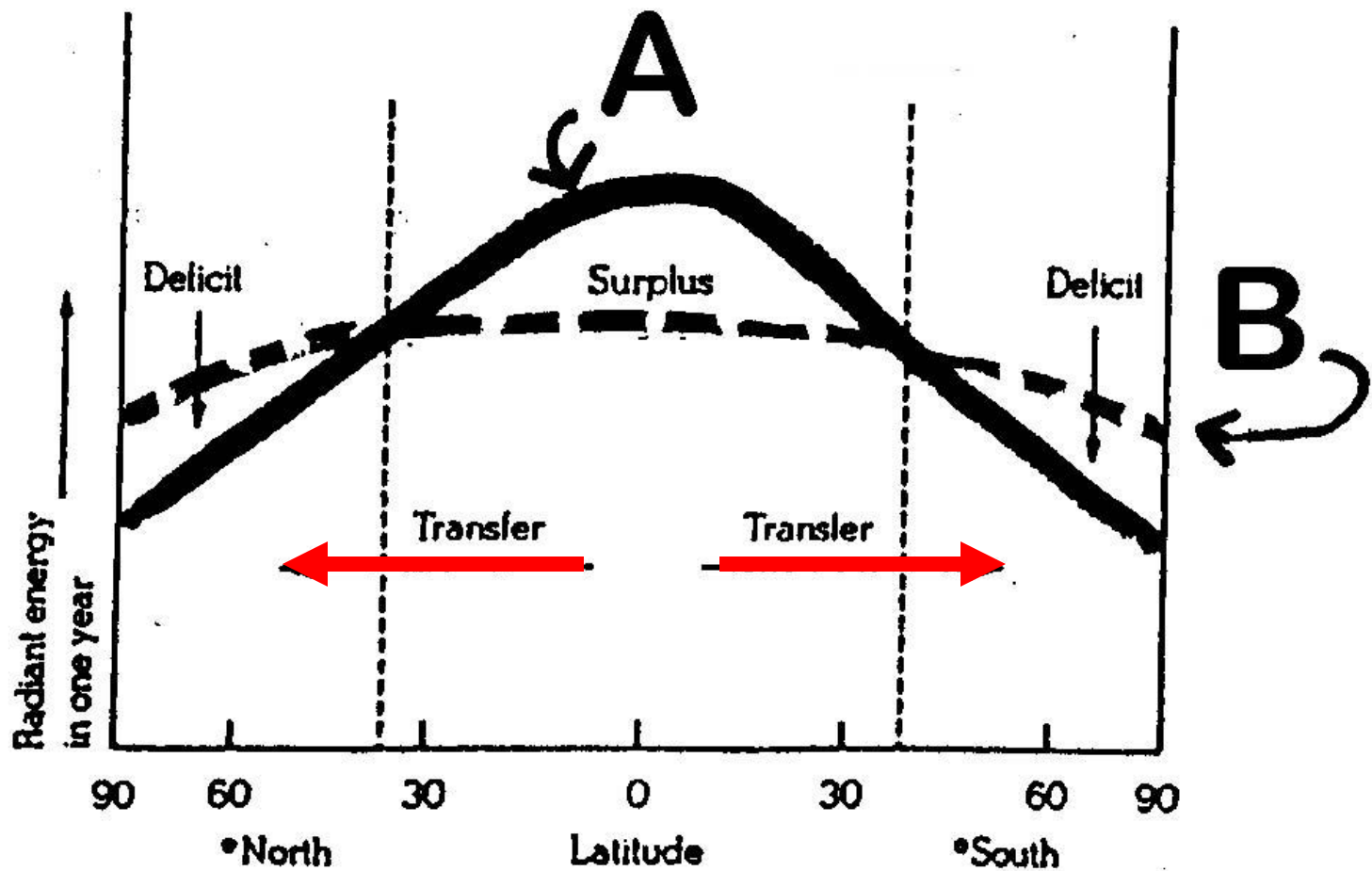


———— Absorbed solar energy



----- Emitted infrared energy
(at top of atmosphere)



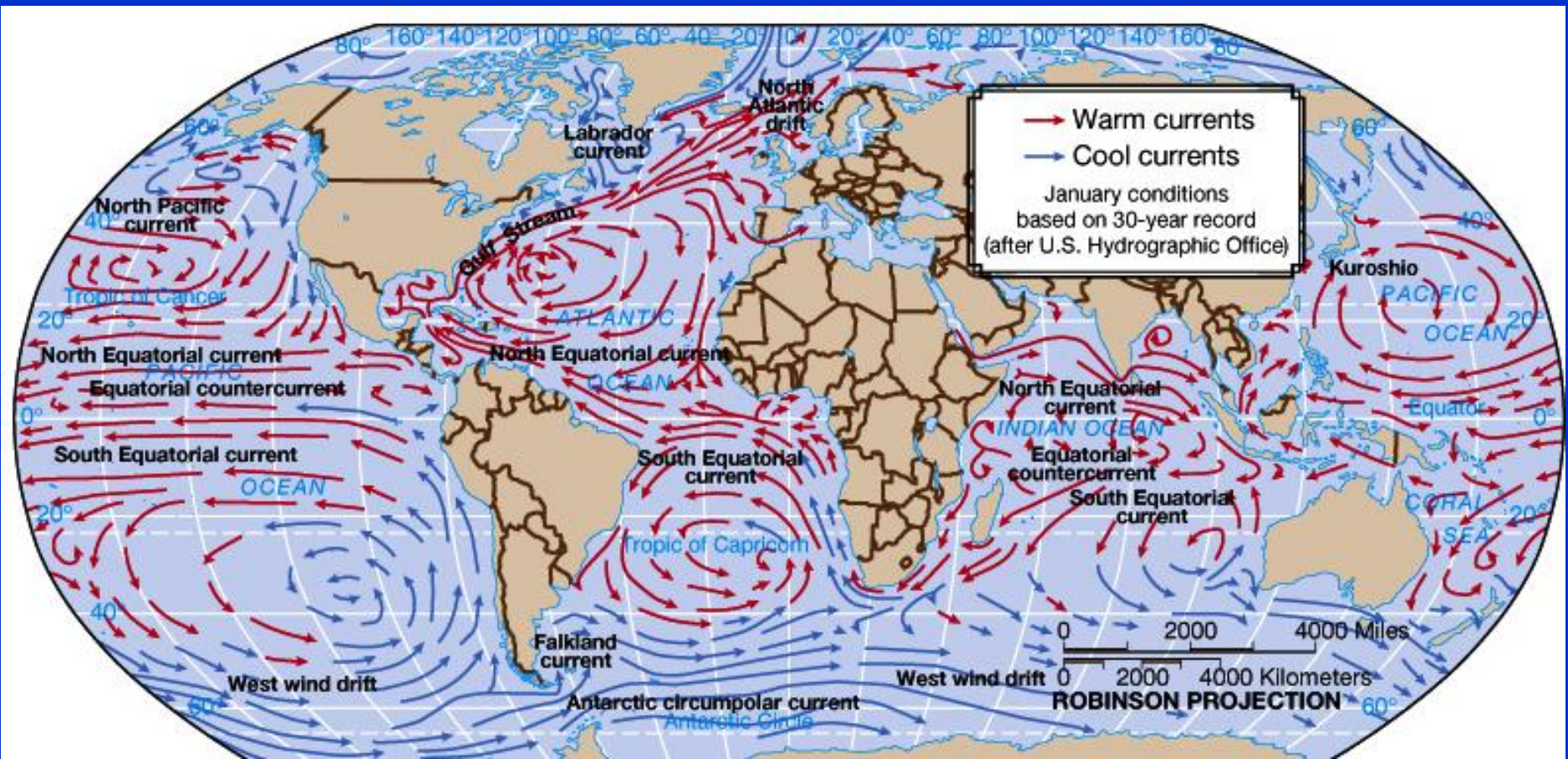


**THERMAL ENERGY TRANSPORTED
FROM LOW TO HIGH LATITUDES TO
BALANCE OUT DEFICIT!**

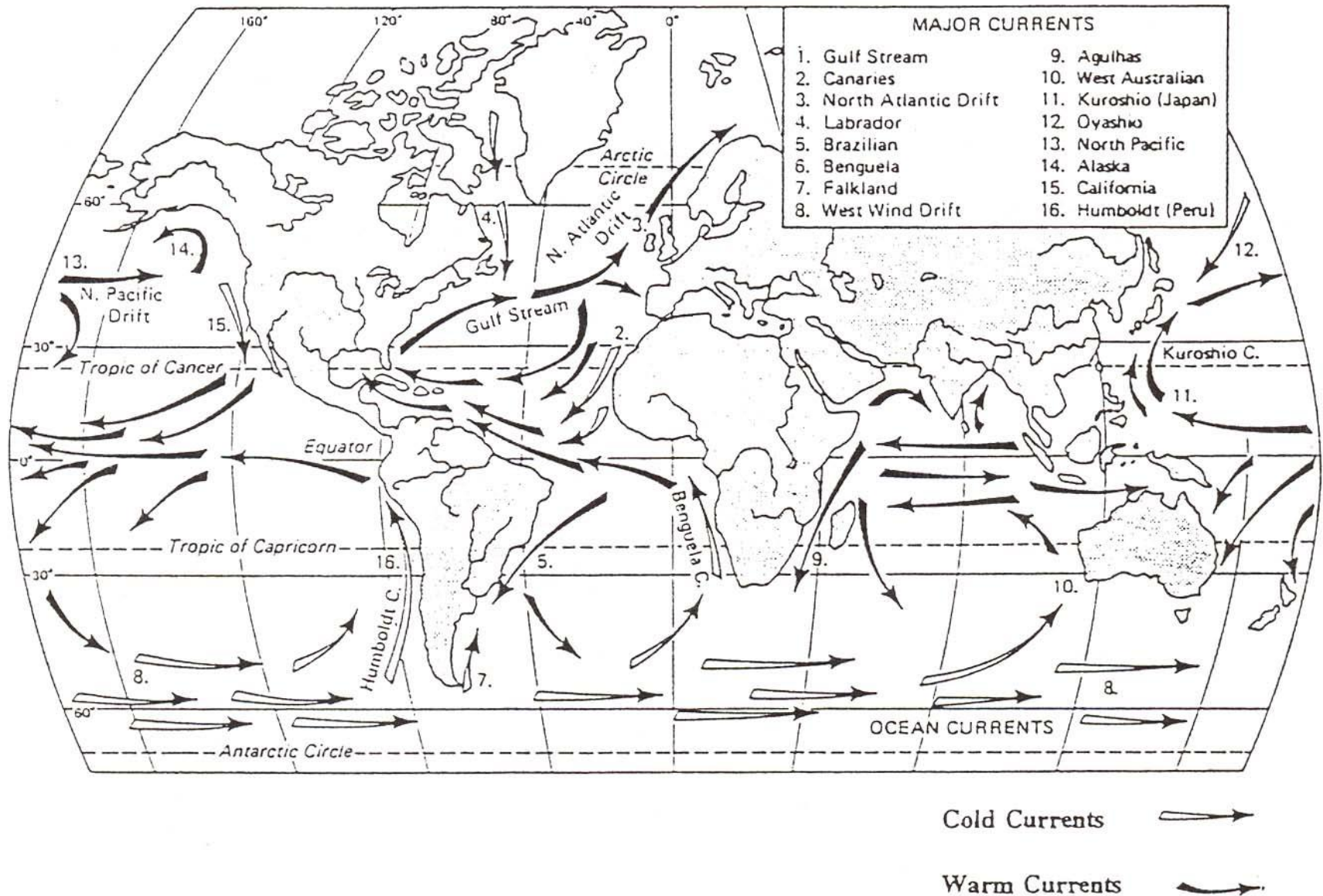
H + LE + G

BUT WHAT ABOUT G ?

G is a STORAGE component, not a transfer component BUT energy stored in the OCEAN, can later be transported via ocean currents as **H** !

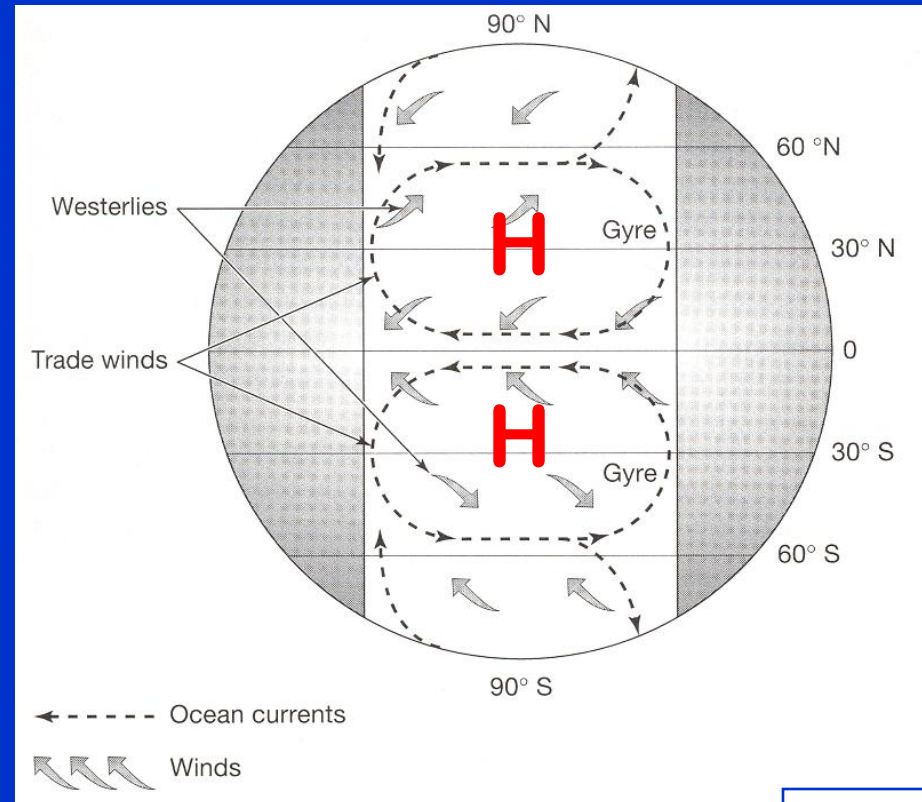
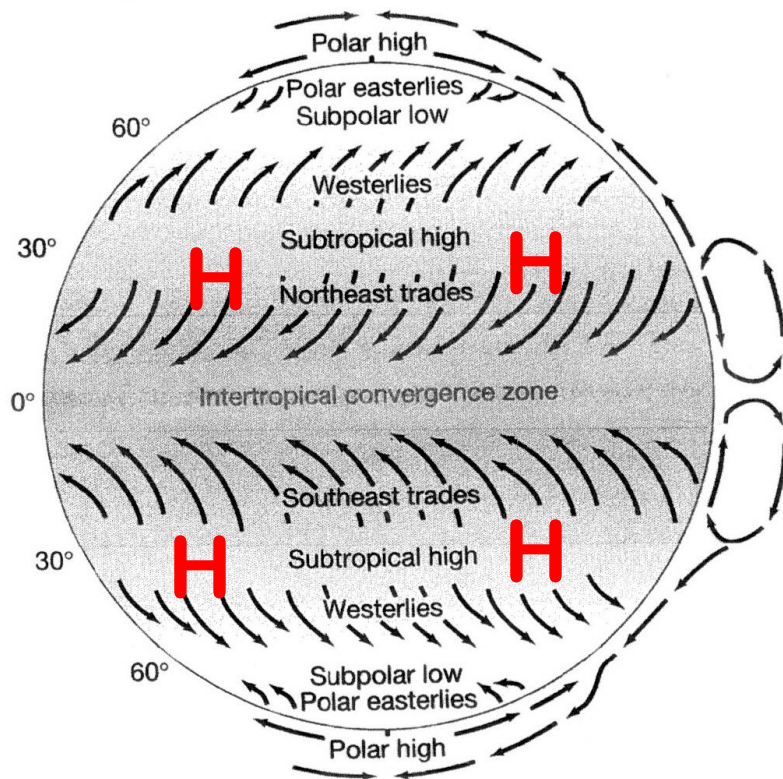


WARM & COLD SURFACE OCEAN CURRENTS:



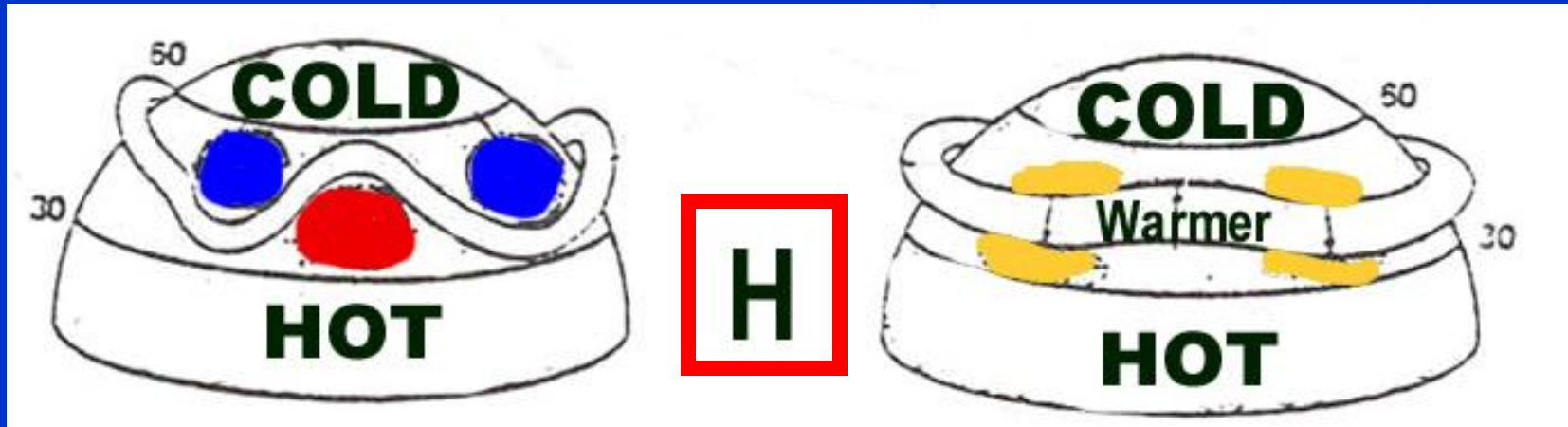
→ Large OCEAN GYRES -- driven by Trade Winds & Westerlies in Oceanic Subtropical HIGH PRESSURE CELLS (STH)

Leads to SURFACE ocean currents

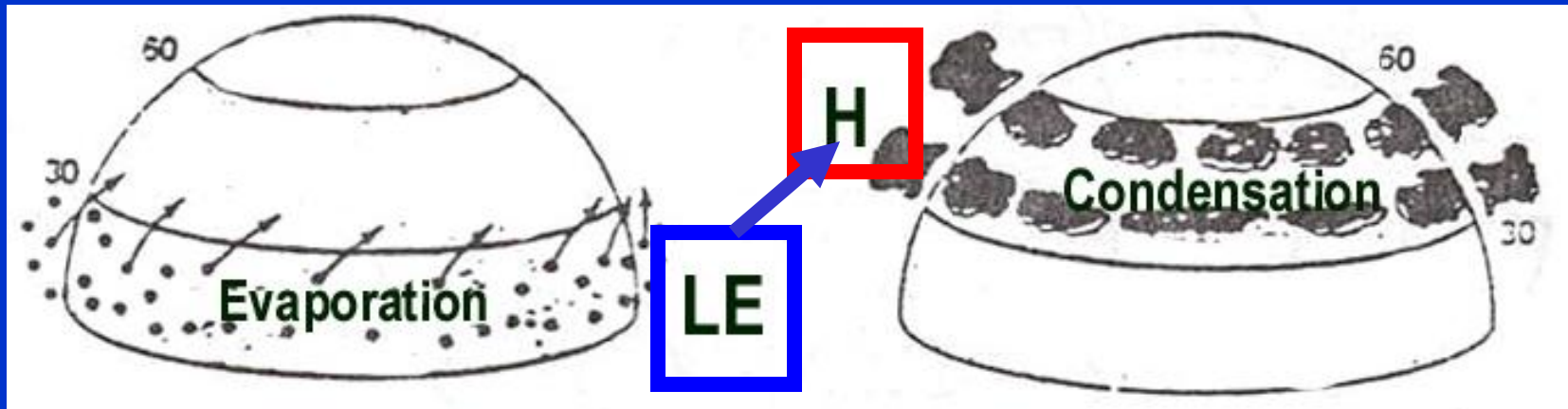


Energy is transported from areas of surplus to deficit via:

H (sensible heat)



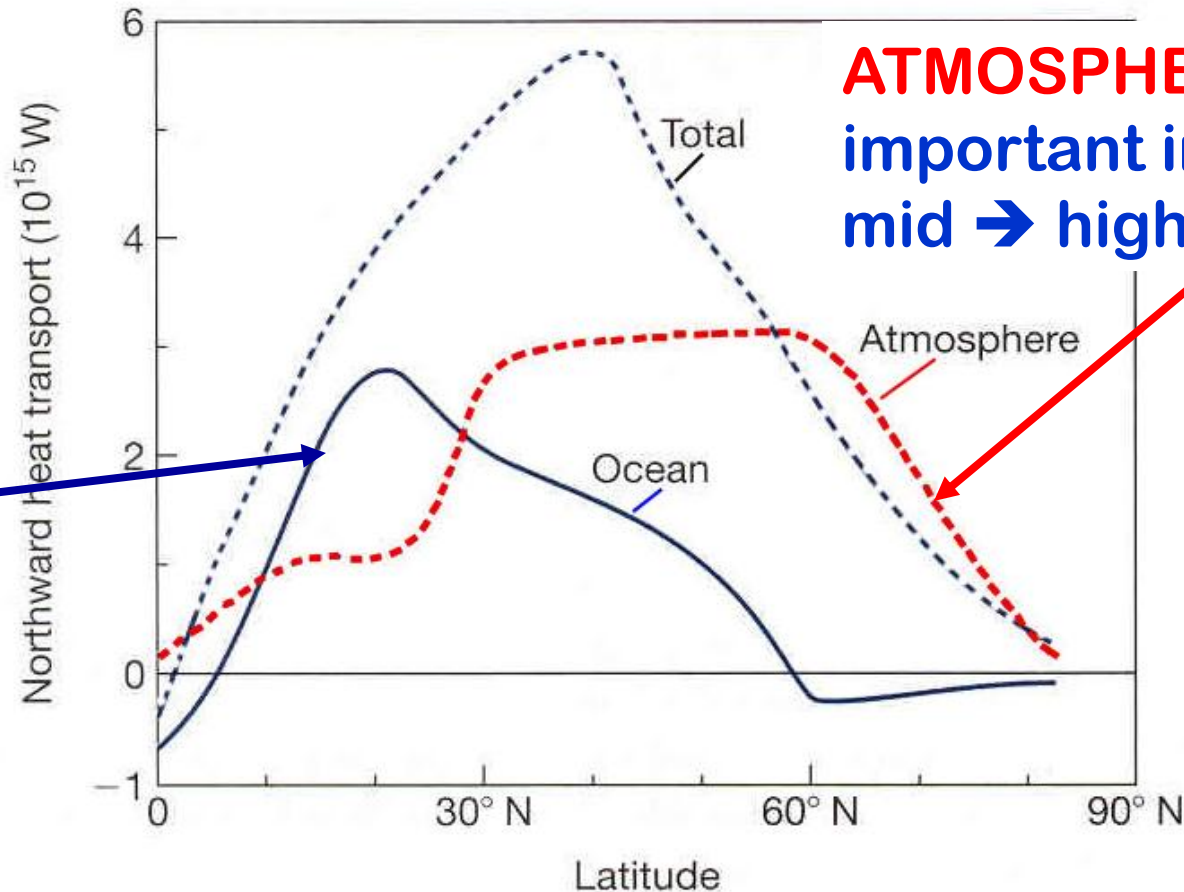
& LE (Latent Energy)



H + LE

Both **ATMOSPHERE** & **OCEAN** play important roles in **BALANCING OUT ENERGY SURPLUS & DEFICIT AREAS**:

OCEAN transports **MOST** of the energy in **LOW** → subtropical latitudes



ATMOSPHERE more important in mid → high latitudes

Poleward transport of energy in N.H.



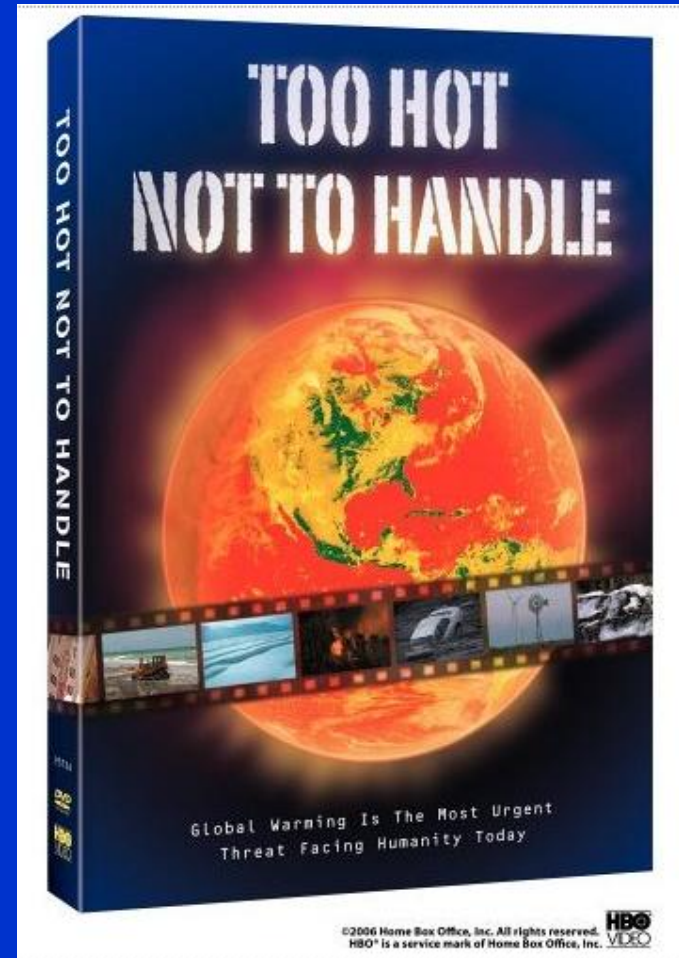
MOVIE TIME!!

GROUP BONUS POINT CHALLENGE :

Watch the video carefully – at some point a feedback loop process is described:

On an index card, state which feedback loop is described and sketch the diagram for it.

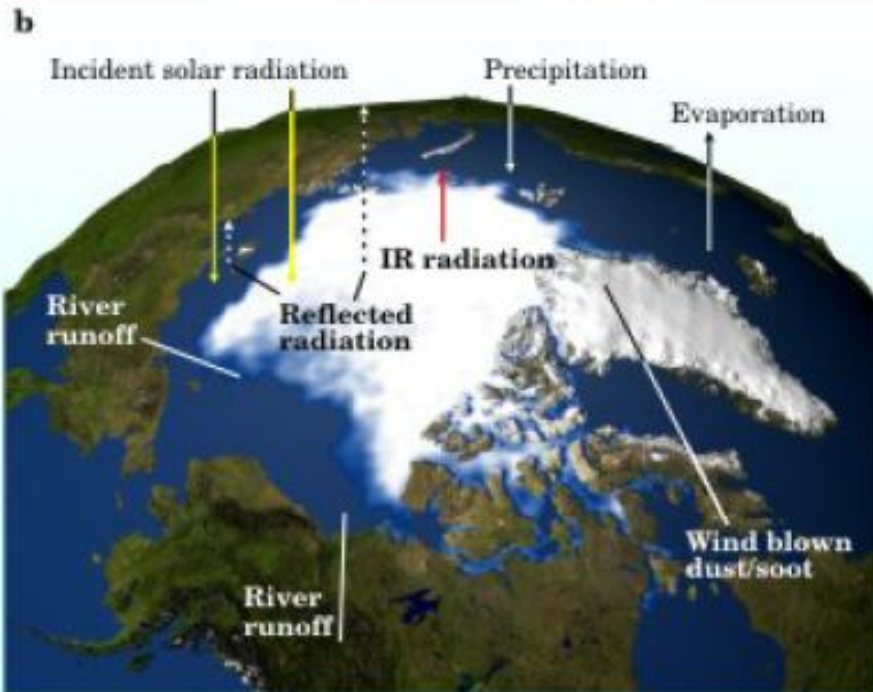
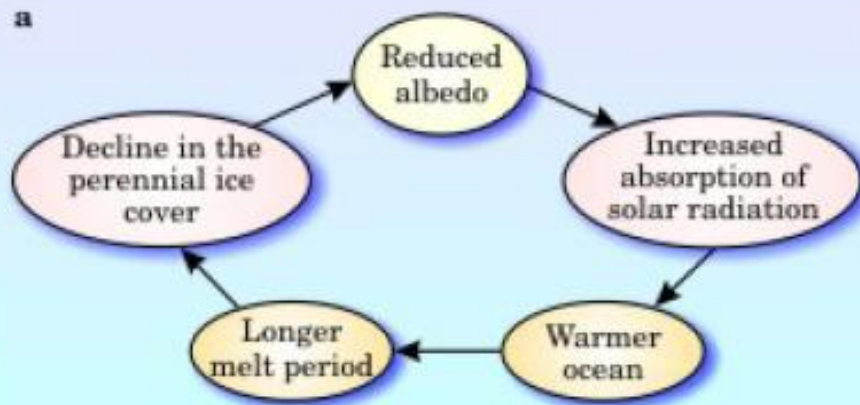
(HINT: it is one of the loops shown on p 56 in Class Notes)



NOTE: This video is posted in D2L. The part we watched was “Melting Alaska” – which starts at **19:57 min**

REMEMBER FEEDBACK LOOPS:

Is this one positive or negative?



THEN – on the back of the index card, as a group complete the feedback loop on the bottom of p 58 page by linking the components with the proper coupling arrow symbols as used in the SGC text

Here are the
components
from p 58:

albedo

Extent of
ice cover

SW
radiation
absorbed

Amount of
melting

Ocean
temperature

The first
coupling has
been done to get
you started!

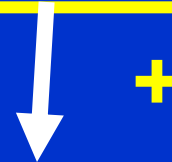
albedo

Extent of
ice cover

SW
radiation
absorbed

Amount of
melting

Ocean
temperature



**We'll finish this and the G-3
Assignment in class on
Thursday.**